

A Theory To Explain Tropical Cyclone Kinetic Energy Spectra

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1. Objective

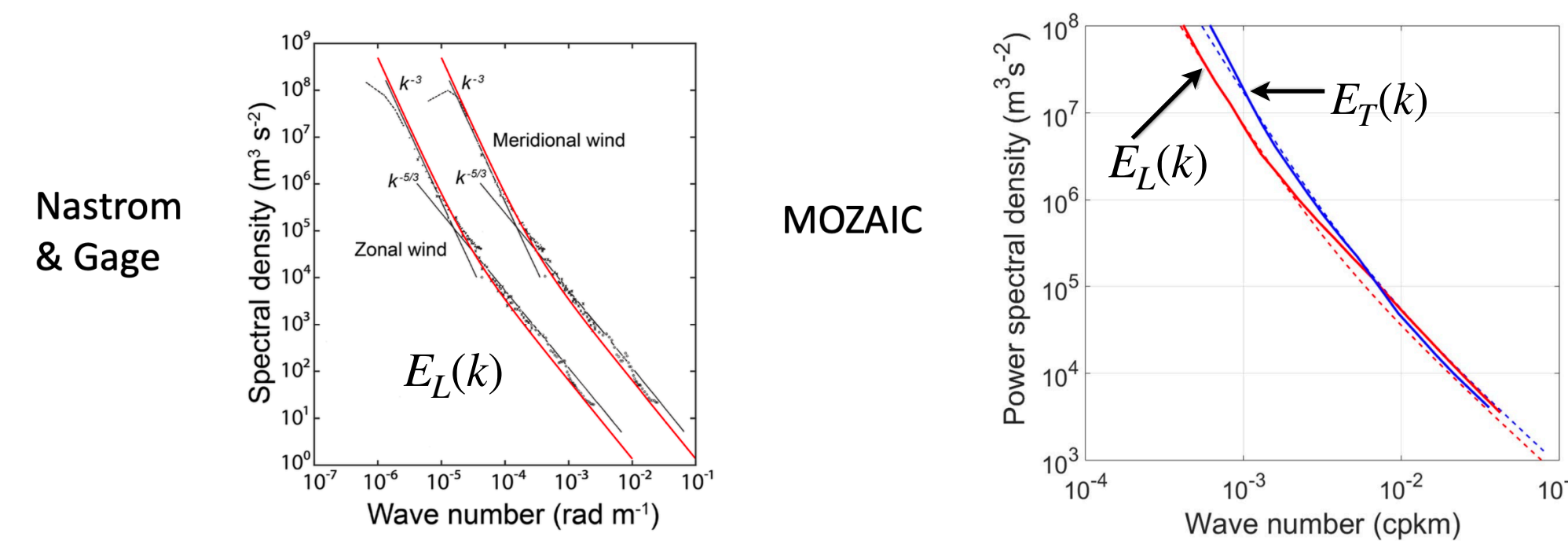
To develop a theory of TC turbulence that successfully explains observations of **1D KE energy spectral slopes and amplitudes**.

2. Flight Data

- Dataset:** Consisted of 3000 reconnaissance aircraft missions into 320 tropical storms and tropical cyclones in the North Atlantic and Eastern Pacific Oceans during the years 1977-2022. Also, NOAA Best Track (NBT) and Hurricane Research Division (HRD) storm parameters and track information.
- Samples:** Only flights along *radial legs* were used. The flight legs spanned the full TC. From these, *sublegs* were extracted that sampled either the TC inner core (IR) or TC outer region (OR).
- Storm classification:** This was carried out by classifying the maximum 60-second sustained wind speed of each flight leg and grouping them according to the Saffir-Simpson scale.
- Wind field regions:** Inner region IR: $0 < r < 2R_{mw}$ and outer region OR: $r > 2R_{mw}$. The IR includes the *eyewall* (green) and the *near-core* (pink). See Fig 1.

3. Previous Work

Observed atmospheric KE spectra (midlatitudes)



Quasi-Normal Scale Elimination (QNSE) theory
See [1]

Navier-Stokes
rotating frame
Coriolis parameter, f

Langevin Eqns
(stochastic forcing)

Apply successive
coarse-graining

Predicted isotropic 1D KE spectra

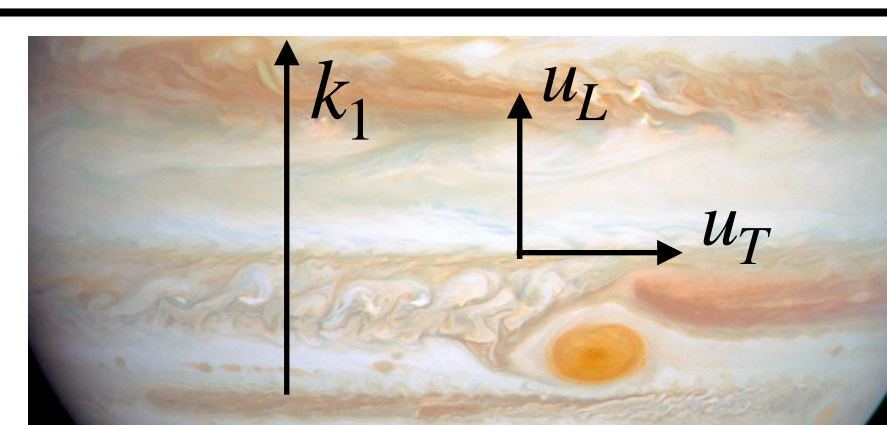
$$E_L(k_1) = C_{KL} \Pi_e^{2/3} k_1^{-5/3} + 0.0926 f^2 k_1^{-3}$$

$$E_T(k_1) = C_{KT} \Pi_e^{2/3} k_1^{-5/3} + 0.24 f^2 k_1^{-3}$$

Note...
dependence on f
peristrophic regime
(peristrophy = rotation)

Simulations of 2D turbulence
 β -plane; surface of rotating sphere ...

- Rossby waves
- Zonal jets



zonal jets

Simulations produced
anisotropic 1D KE spectra

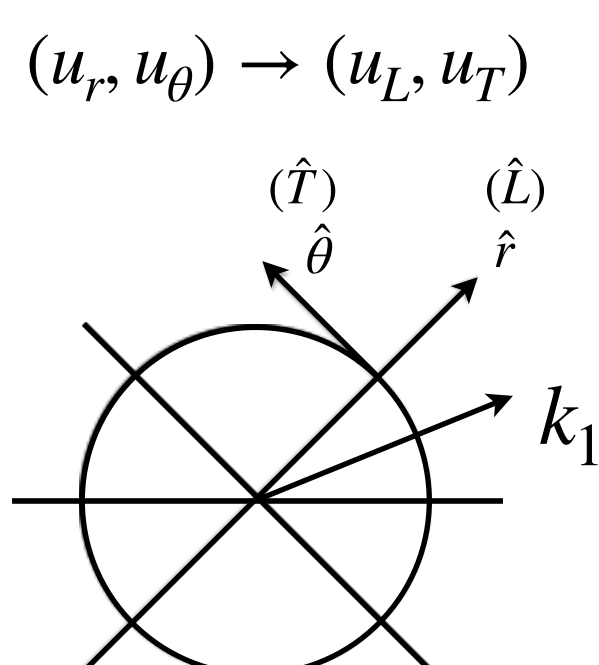
$$E_L(k_1) = C_{KL} \Pi_e^{2/3} k_1^{-5/3}$$

$$E_T(k_1) = C_{KT} \Pi_e^{2/3} k_1^{-5/3} + C_z \beta^2 k_1^{-5}$$

($k_1 \perp$ to axis of zonal jet)

Note...
dependence on β
 k_1^{-5} spectra
zonostrophic regime

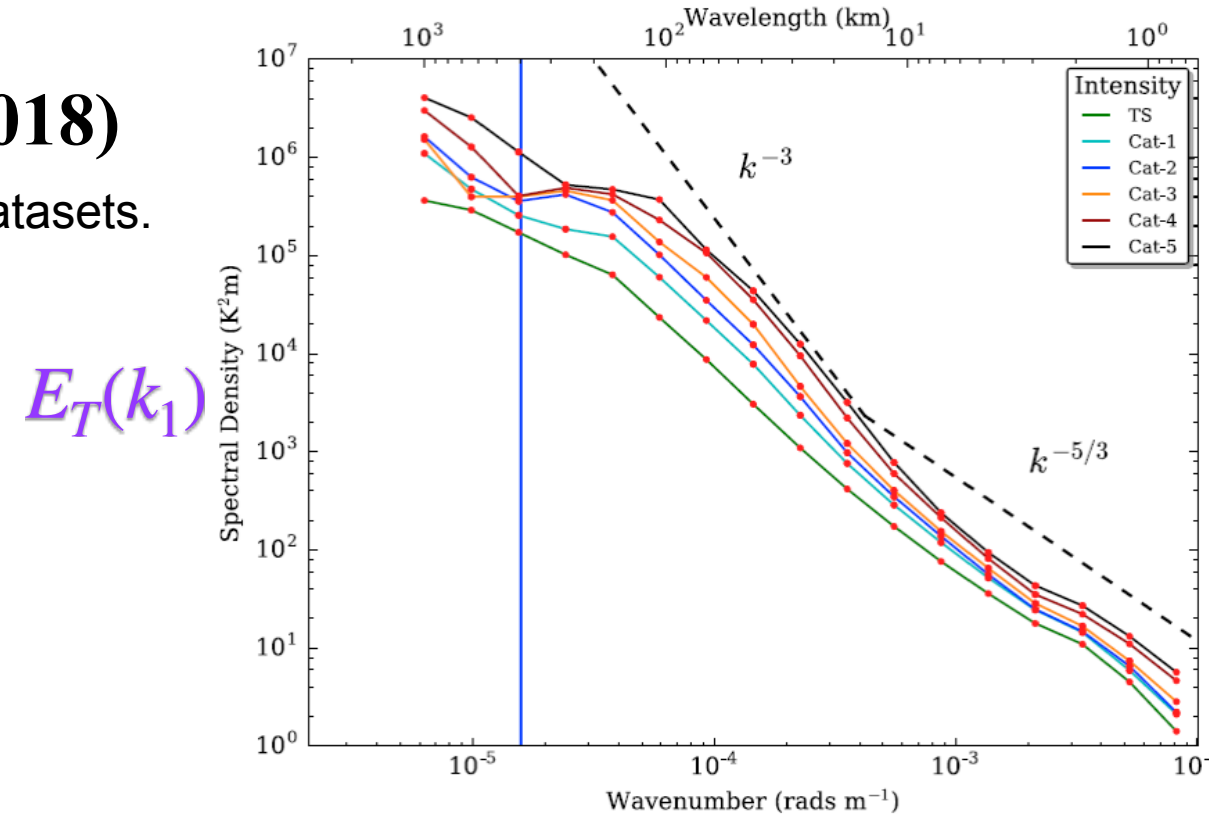
4. Tropical Cyclones — Previously observed spectra



Vonich & Hakim (2018)

Spectral composites of three datasets.
Radial flight legs span TC.
Taken from Ref [2].

- flight legs grouped into classes defined by TC category.
- spectra ensemble averaged for each class.



Results ...

- As TC intensity increases ...
- amplitude increases
- slopes steepen
- resemble NASTROM & GAGE spectra

5. TC Model

- Barotropic framework
- Rapidly spinning vortex ($\hat{f}$) in a planetary flow (f) ...

$\Rightarrow \hat{f}$: **cyclostrophic** Coriolis parameter

$\Rightarrow \hat{f}$: **effective** Coriolis parameter

$$\hat{f} = f + \hat{f}$$

Principle of Least Action
(Landau & Lifshitz: *Mechanics*)

If $\hat{f} = \text{constant}$ in the TC core, theory predicts ...

$$E_L(k_1) = C_{KL} \Pi_e^{2/3} k_1^{-5/3} + 0.0926 \hat{f}^2 k_1^{-3}$$

$$E_T(k_1) = C_{KT} \Pi_e^{2/3} k_1^{-5/3} + 0.24 \hat{f}^2 k_1^{-3}$$

If $\hat{f} \neq \text{constant}$, there is a **cyclostrophic β -effect** — which generates **Vortex Rossby Waves (VRWs)** in the TC eyewall. ($\hat{\beta} \approx \hat{f}/R_{mw}$)

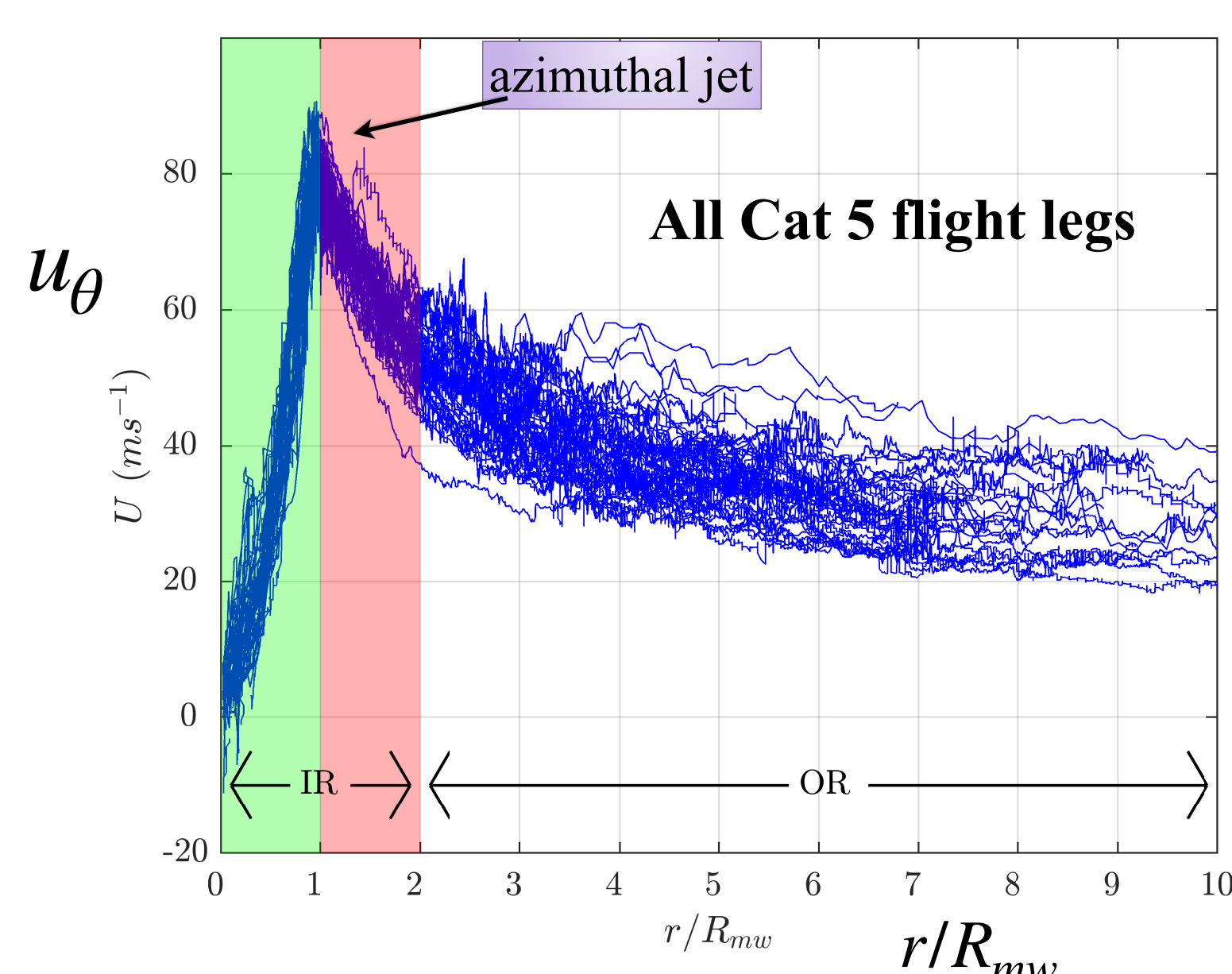
VRWs interact with turbulence and produce **anisotropic 1D KE spectra** ...

$$E_L(k_1) = C_{KL} \Pi_e^{2/3} k_1^{-5/3} + C_{fL} \hat{f}^2 k_1^{-3}$$

$$E_T(k_1) = C_{KT} \Pi_e^{2/3} k_1^{-5/3} + C_{fT} \hat{f}^2 k_1^{-3} + C_z \hat{\beta}^2 k_1^{-5}$$

6. Comparison with Observations

Structure of the TC wind field



IR
Inner Region -
includes eyewall and
near-core (pink)

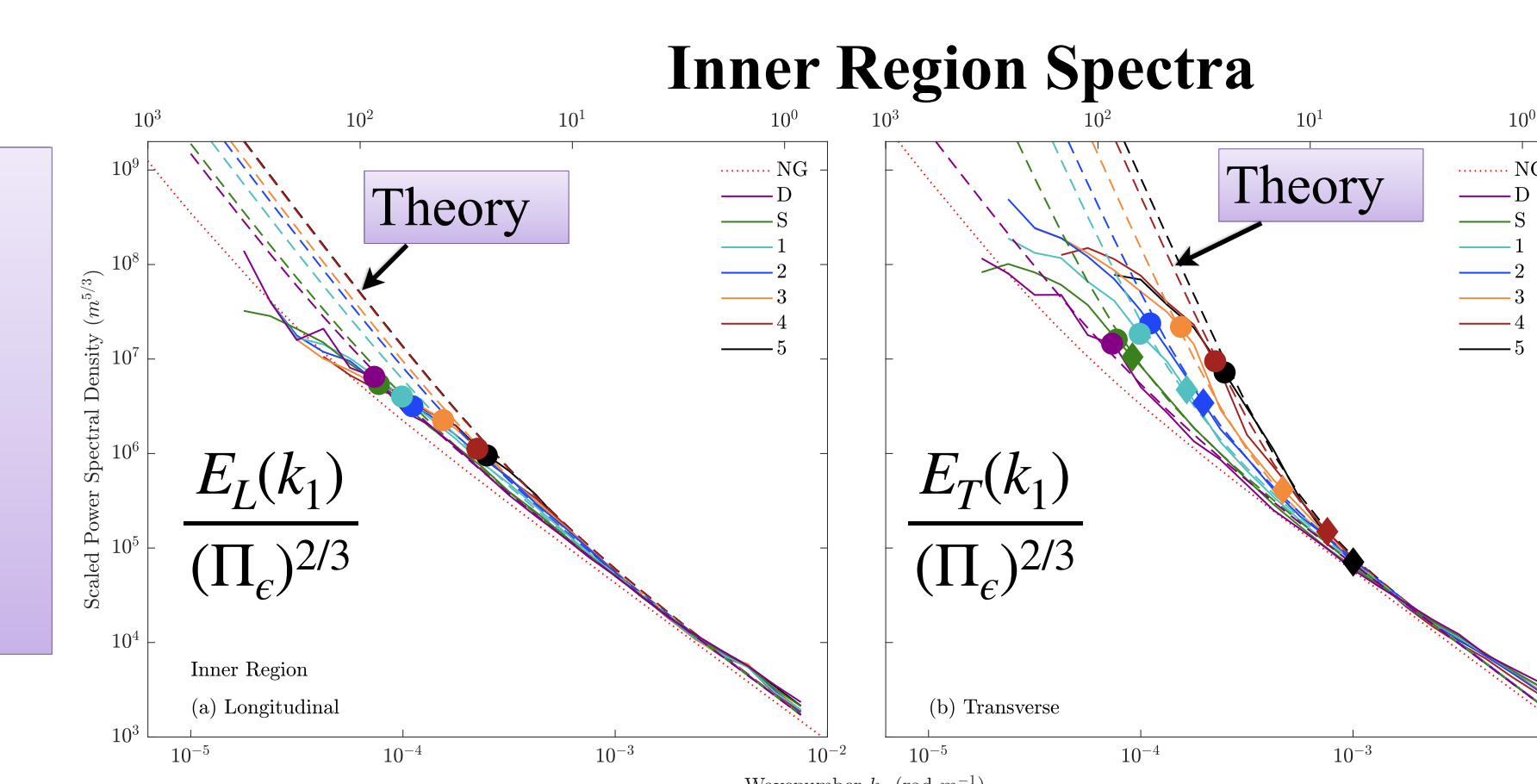
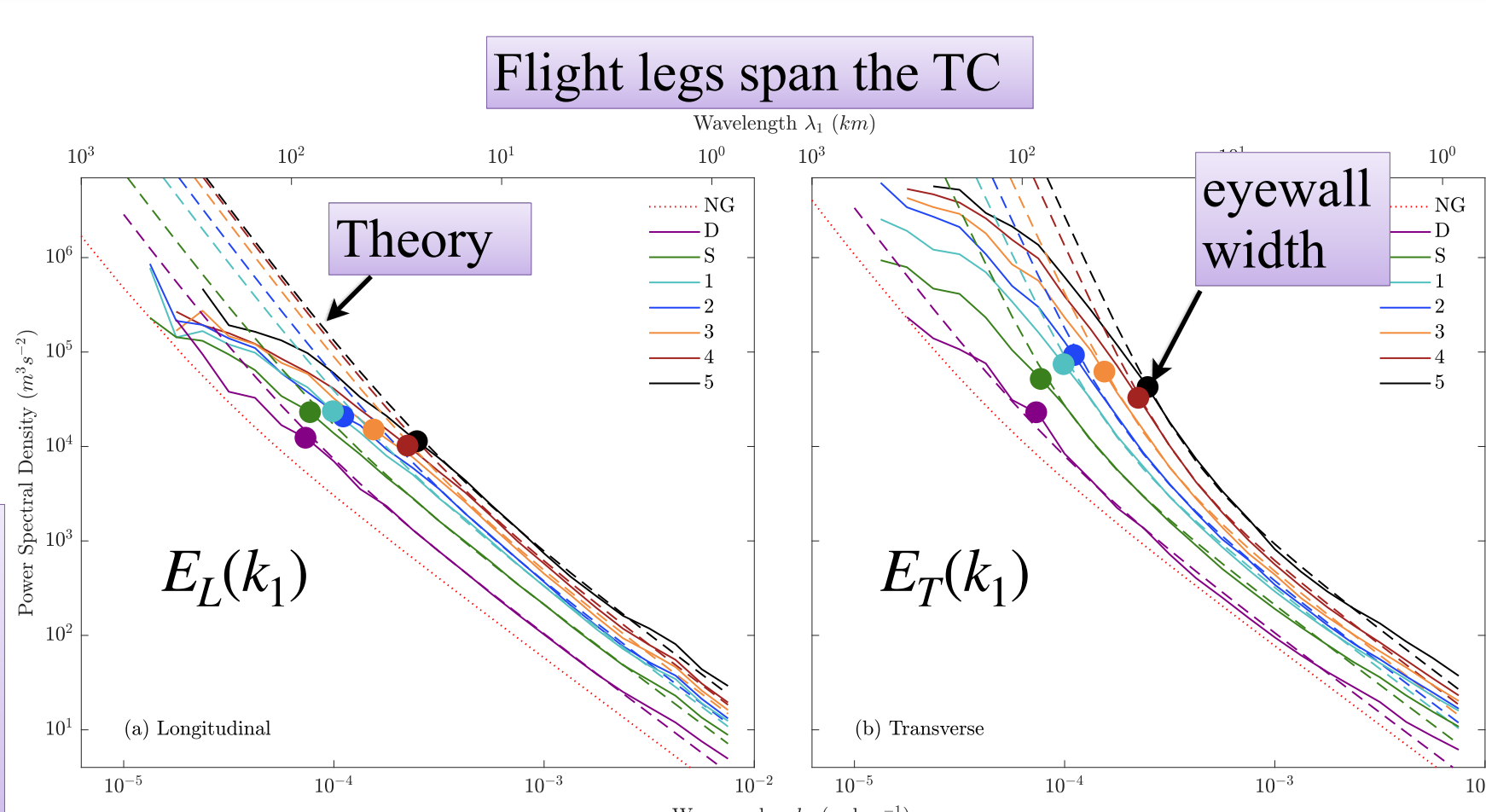
OR
Outer Region

Eyewall azimuthal velocity profile resembles an eastward zonal jet found on Saturn and Jupiter!

Amplitudes
increase with
TC category

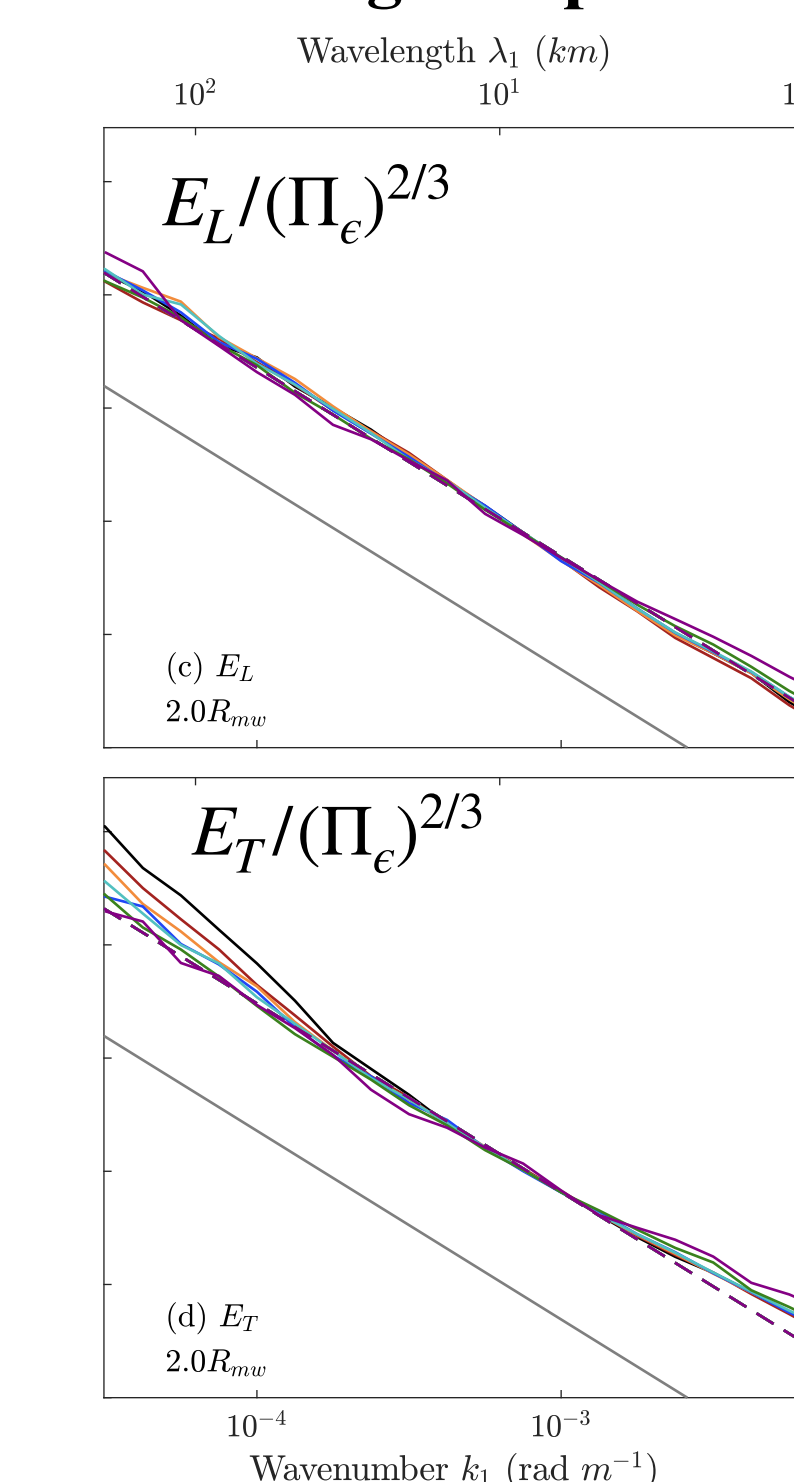
Fig shows that
both E_L and E_T
are needed to
account for
spectral
anisotropy

Normalization
collapses all spectra in
the Kolmogorov
subrange.
The subrange shrinks
with increasing TC
intensity.

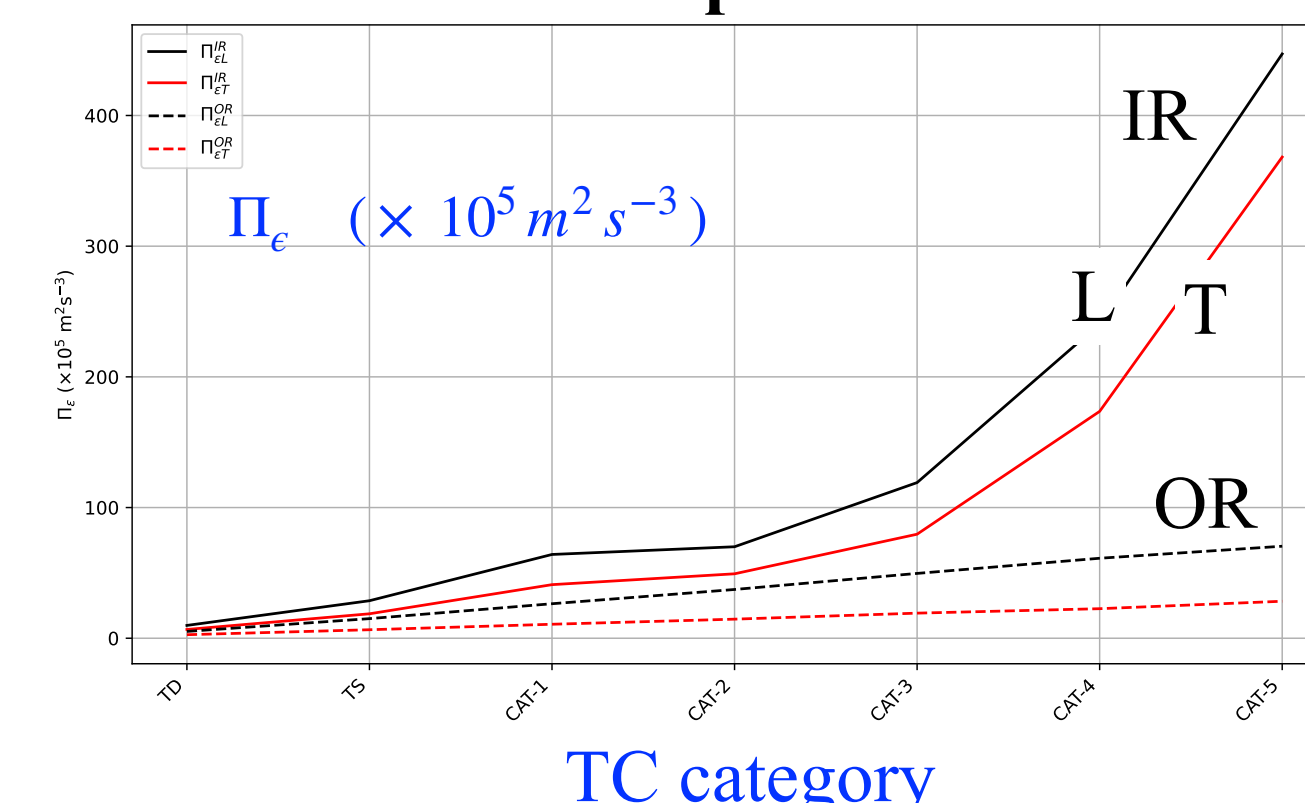


The diamonds mark the boundary between peristrophic and zonostrophic subranges (see Ref [3]).

Outer Region Spectra



TC Dissipation Rates



6. Conclusion:

Observed TC spectra shows that cyclostrophic turbulence evolves from purely peristrophic to mixed peristrophic-zonostrophic turbulence. See Ref [3] for more results and conclusions.

References

- [1] Sukoriansky, S., & Galperin, B. (2016). QNSE theory of turbulence anisotropization and onset of the inverse energy cascade by solid body rotation. *J. Fluid Mech.*, **805**, 384–421.
- [2] Vonich, P. T., & Hakim, G. J. (2018). Hurricane kinetic energy spectra from in situ aircraft observations. *J. Atmos. Sci.*, **75**, 2523–2532.
- [3] Galperin, Nickerson, King and Zhang (2025).