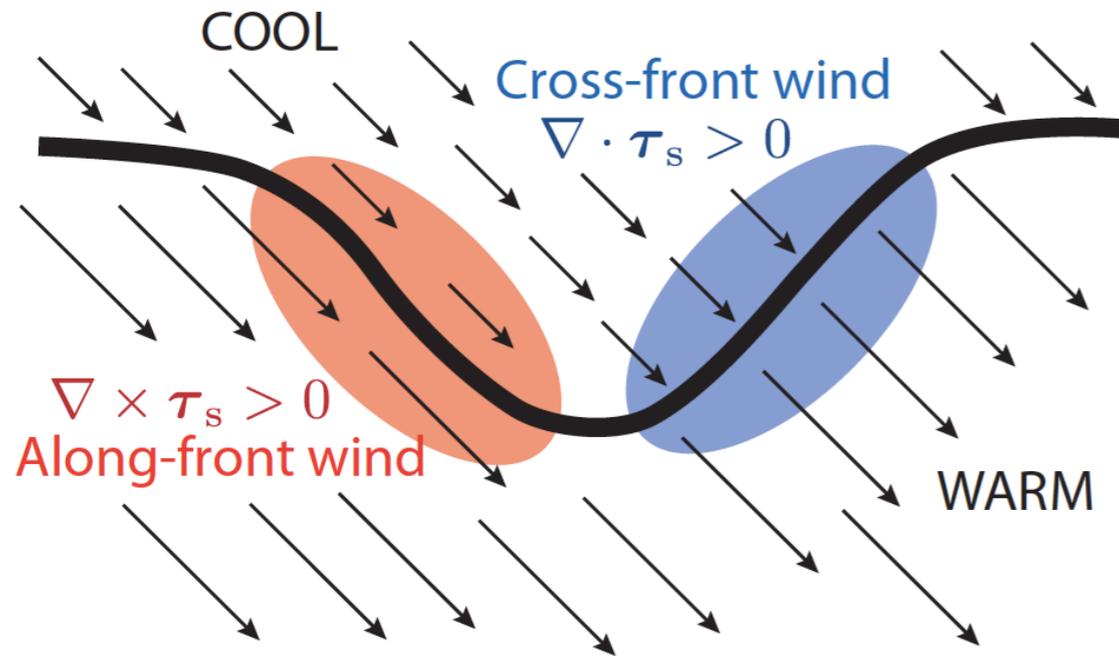


Characterization of the frontal air-sea interaction by transfer functions

Niklas Schneider and Kazuyoshi Kikuchi

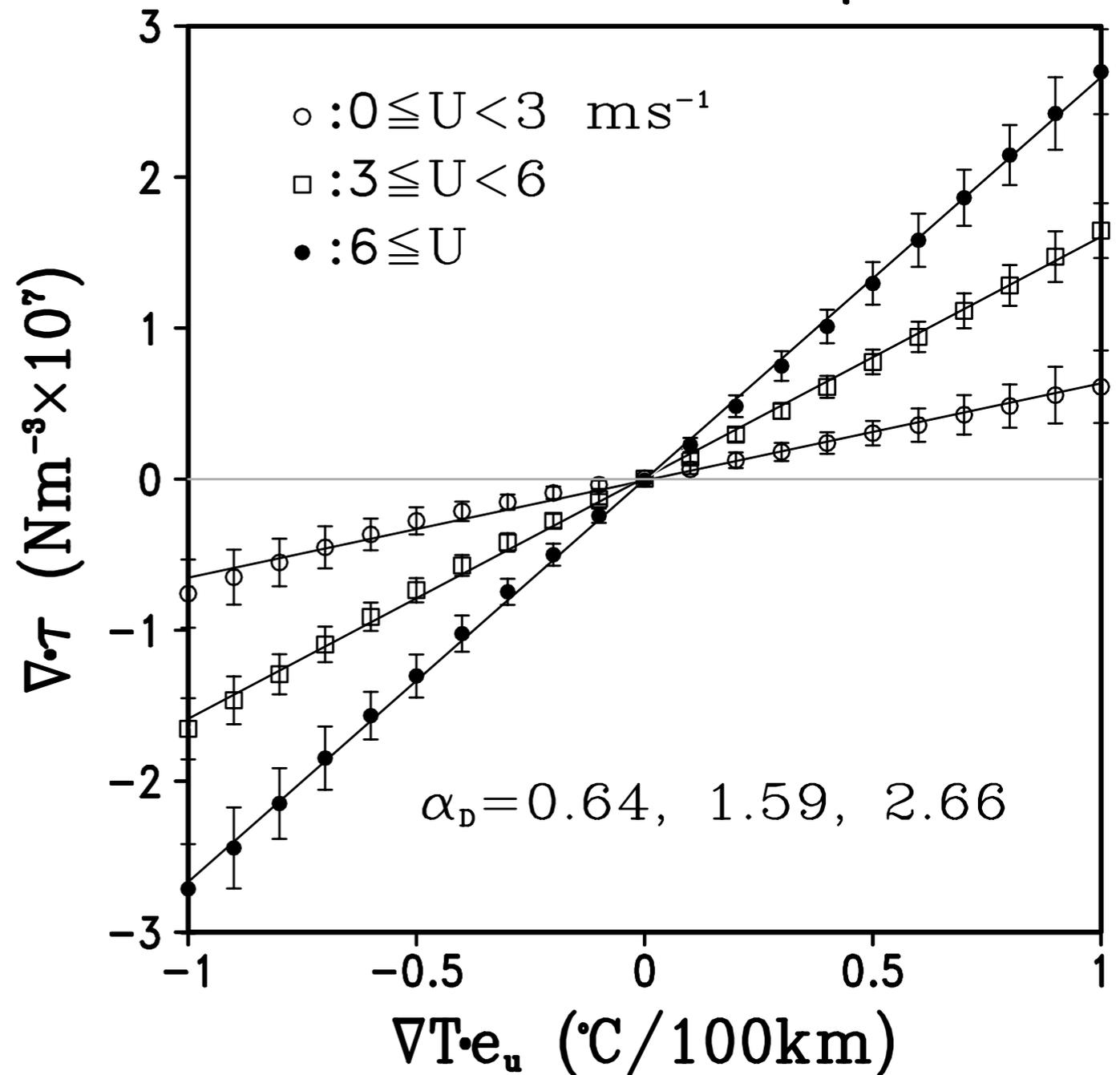
International Pacific Research Center, University of Hawaii at Manoa



Chelton et al., 2004

SCOW project:
 QuikSCAT wind stress
 AVHRR SST
 monthly averages
 Risien and Chelton 2008

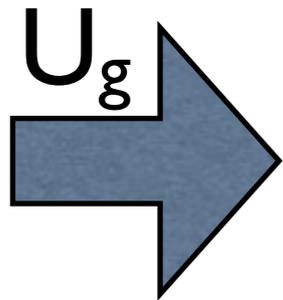
Southern Hemisphere



Model for air-sea interaction at SST fronts

- Reduced gravity model capped by sharp inversion
- Forced by barotropic tropospheric pressure gradient
- Background state: SST constant

$h^{(0)}$ ————— inversion, $\Delta\Theta$, no flux



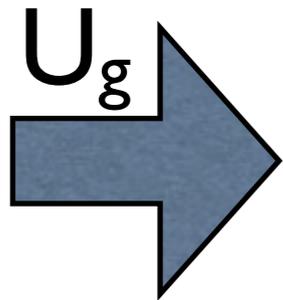
$u^{(0)}, v^{(0)}$ Ekman spiral
 $\Theta^{(0)}$ constant

—————
 $T^{(0)}$ constant

Model for air-sea interaction at SST fronts

- Reduced gravity model capped by sharp inversion
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$h^{(0)}$ ————— inversion, $\Delta\Theta$, no flux



$u^{(0)}, v^{(0)}$ Ekman spiral
 $\Theta^{(0)}$ constant

—————
 $T^{(0)}$ constant

- consider weak front $T^{(0)} + \varepsilon T^{(1)}$, linear response

Air-sea interaction at weak SST front

1st order (linear) response

$$\bar{\vec{u}}^{(0)} \cdot \nabla \Theta^{(1)} = \gamma (T^{(1)} - \Theta^{(1)}) + A_h \nabla^2 \Theta^{(1)}$$

nondimensionalized by Rossby radius of deformation, boundary layer height, inversion strength etc.

Air-sea interaction at weak SST front

1st order (linear) response

$$\bar{\vec{u}}^{(0)} \cdot \nabla \Theta^{(1)} = \gamma (T^{(1)} - \Theta^{(1)}) + A_h \nabla^2 \Theta^{(1)}$$

$$\underbrace{\vec{u}^{(0)} \cdot \nabla \vec{u}^{(1)} + w^{*(1)} \partial_s \vec{u}^{(0)}}_{\text{advection}} + \underbrace{\hat{e}_3 \times \vec{u}^{(1)}}_{\text{Coriolis}} + \underbrace{\nabla h^{(1)}}_{\text{back pressure}} - \partial_s \underbrace{E^{(0)} \partial_s \vec{u}^{(1)}}_{\text{background mixing}} = \vec{F}$$

$$\vec{u}^{(0)} \cdot \nabla h^{(1)} + \nabla \cdot \vec{u}^{(1)} + \partial_s w^{*(1)} = 0$$

nondimensionalized by Rossby radius of deformation, boundary layer height, inversion strength etc.

Air-sea interaction at weak SST front

1st order (linear) response

$$\bar{\vec{u}}^{(0)} \cdot \nabla \Theta^{(1)} = \gamma (T^{(1)} - \Theta^{(1)}) + A_h \nabla^2 \Theta^{(1)}$$

$$\vec{F} = \nabla \underbrace{\int_s^1 \Theta^{(1)} ds'}_{\text{pressure gradient mechanism}} + \partial_s \left(\underbrace{\delta^{(1)} \frac{\partial E}{\partial \delta} \Big|_{\delta^{(0)}} \partial_s \vec{u}^{(0)}}_{\text{vertical mixing mechanism}} \right)$$

$\delta^{(1)} = T^{(1)} - \Theta^{(1)}$
 air-sea temperature differences modulates vertical eddy viscosity

$$\underbrace{\vec{u}^{(0)} \cdot \nabla \vec{u}^{(1)} + w^{*(1)} \partial_s \vec{u}^{(0)}}_{\text{advection}} + \underbrace{\hat{e}_3 \times \vec{u}^{(1)}}_{\text{Coriolis}} + \underbrace{\nabla h^{(1)}}_{\text{back pressure}} - \partial_s \underbrace{E^{(0)} \partial_s \vec{u}^{(1)}}_{\text{background mixing}} = \vec{F}$$

$$\vec{u}^{(0)} \cdot \nabla h^{(1)} + \nabla \cdot \vec{u}^{(1)} + \partial_s w^{*(1)} = 0$$

nondimensionalized by Rossby radius of deformation, boundary layer height, inversion strength etc.

Transfer function

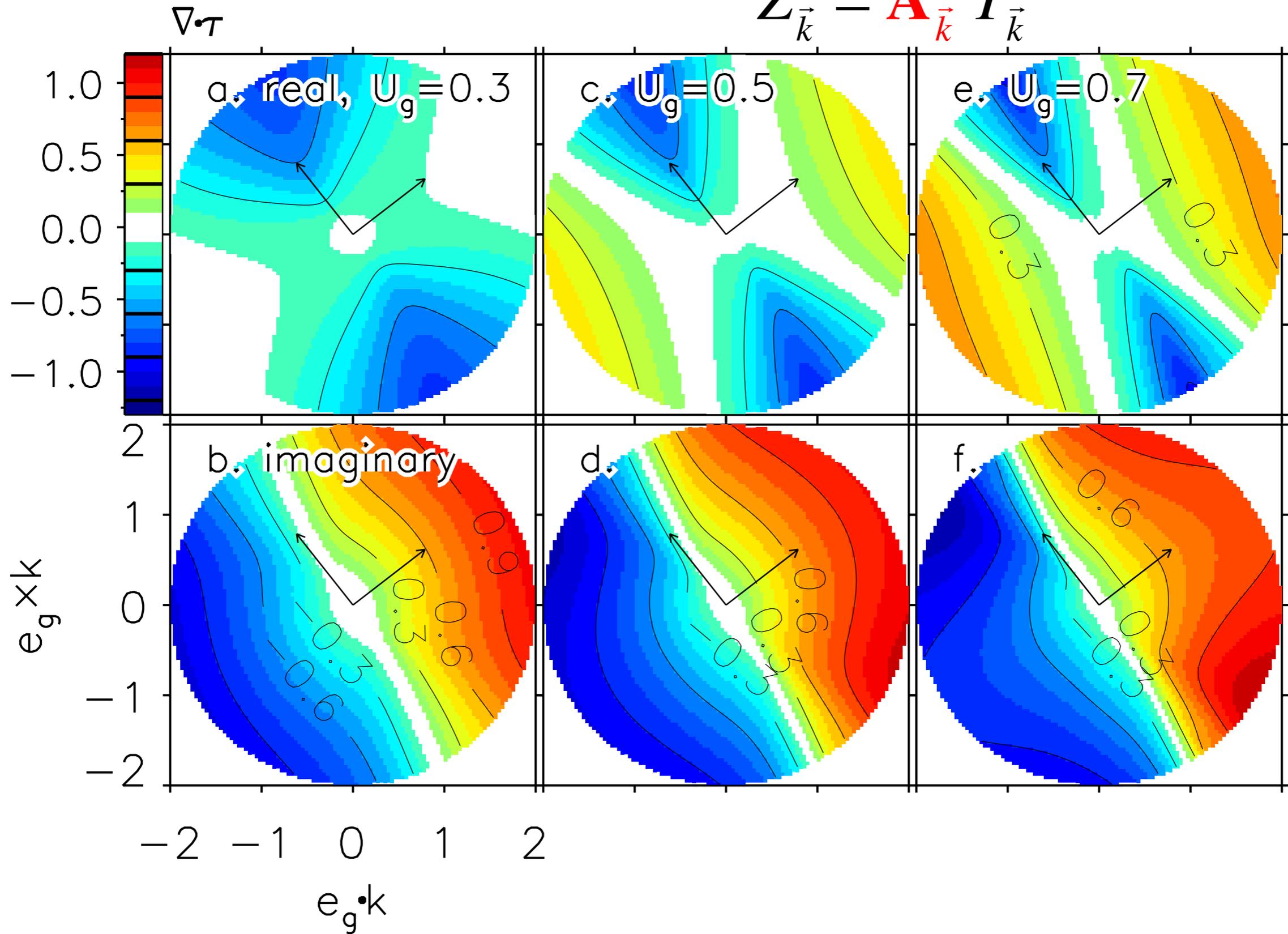
$$\hat{Z}_{\vec{k}} = \hat{A}_{\vec{k}} \hat{T}_{\vec{k}}$$

real part: in phase relationship between Z and T
imaginary part: 90° phase shifted

Transfer function for wind-stress divergence

Linear model

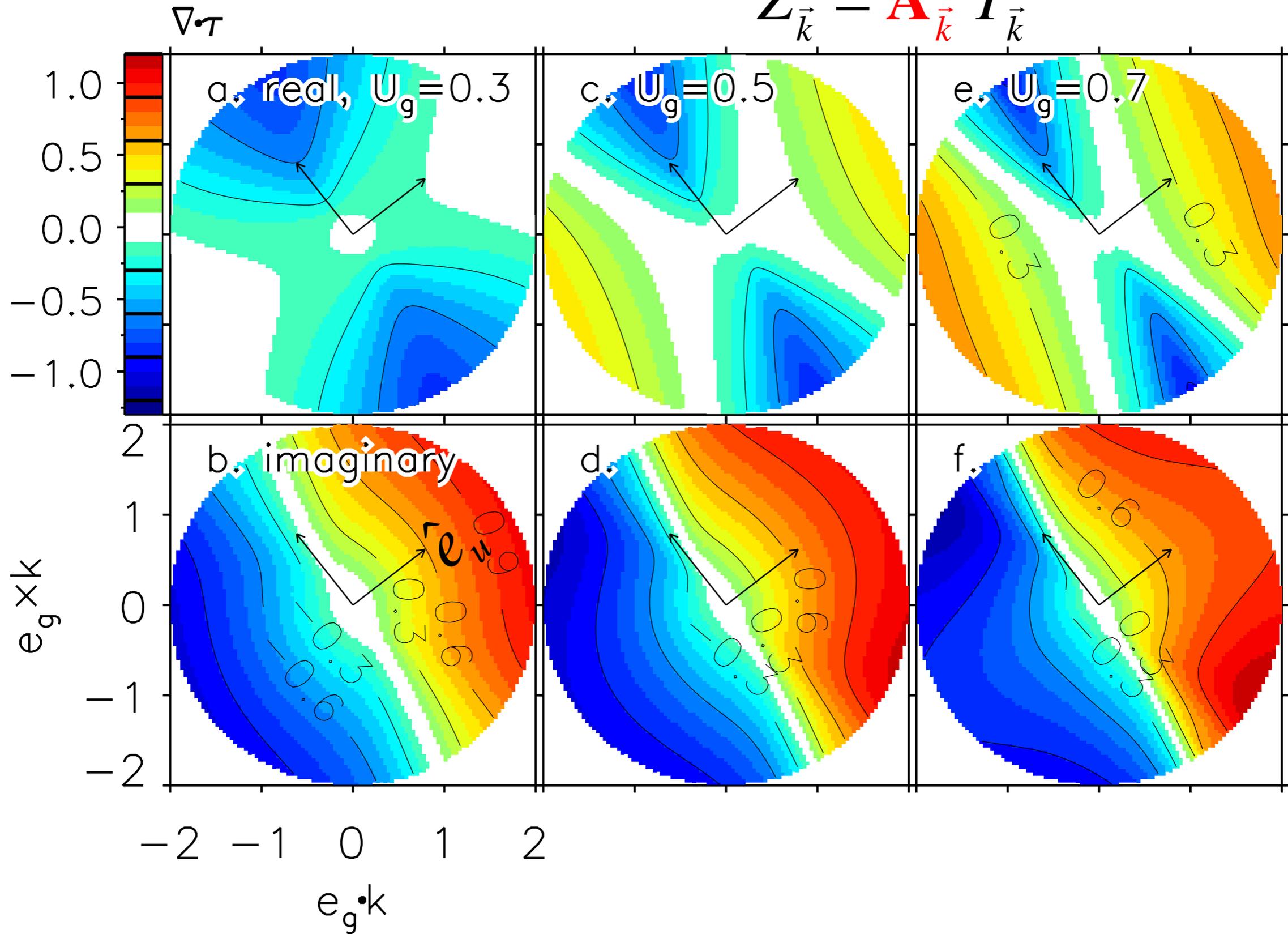
$$\hat{Z}_{\vec{k}} = \hat{A}_{\vec{k}} \hat{T}_{\vec{k}}$$



Transfer function for wind-stress divergence

Linear model

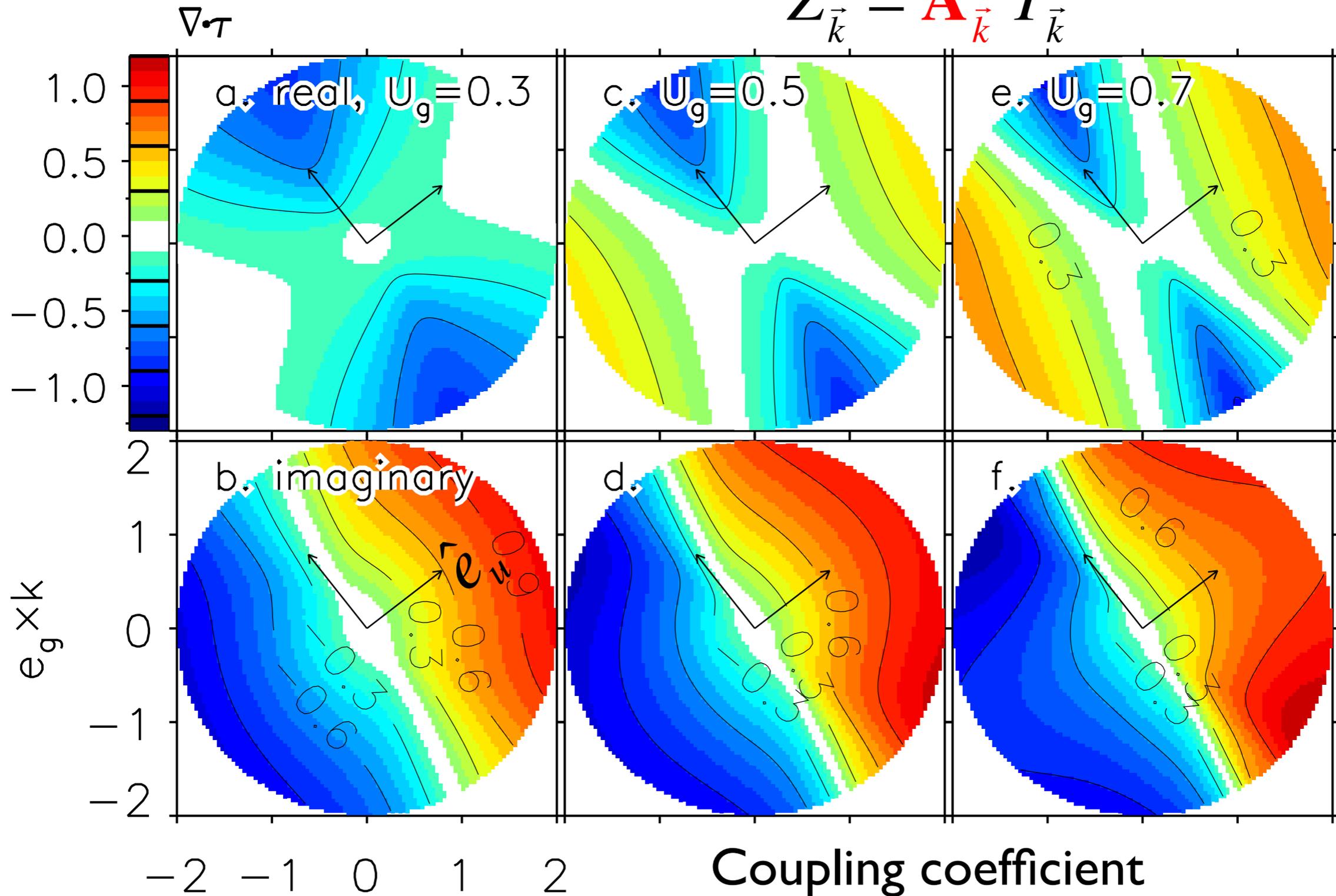
$$\hat{Z}_{\vec{k}} = \hat{A}_{\vec{k}} \hat{T}_{\vec{k}}$$



Transfer function for wind-stress divergence

Linear model

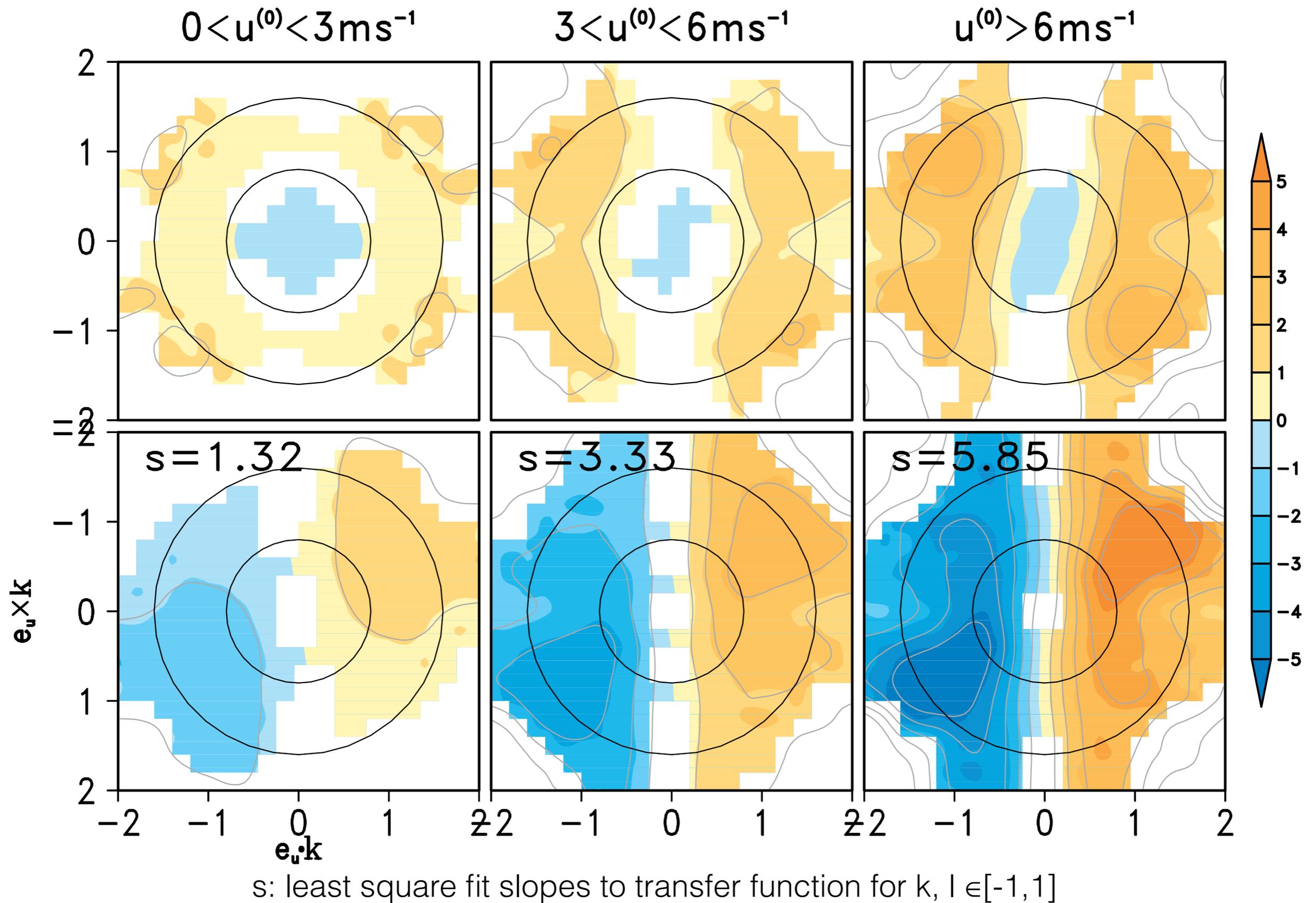
$$\hat{Z}_{\vec{k}} = \hat{A}_{\vec{k}} \hat{T}_{\vec{k}}$$



$$\hat{Z}_{\vec{k}} = \alpha i \hat{e}_u \cdot \vec{k} \hat{T}_{\vec{k}} + \hat{\epsilon}$$

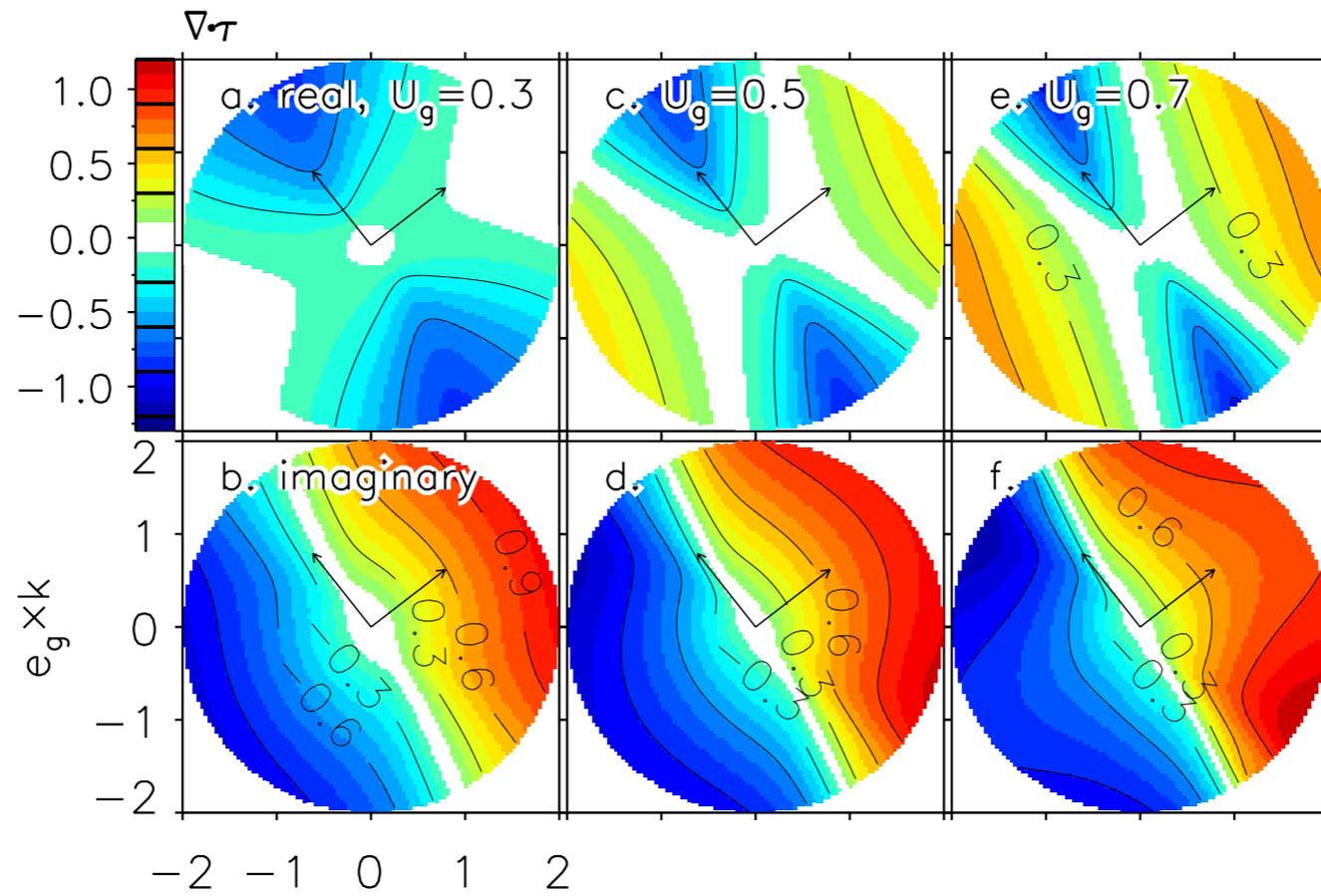
Transfer function for wind-stress divergence

SCOW QuikScat/AVHRR (Risien and Chelton 2008)
1999-2009 monthly averages, Southern Hemisphere

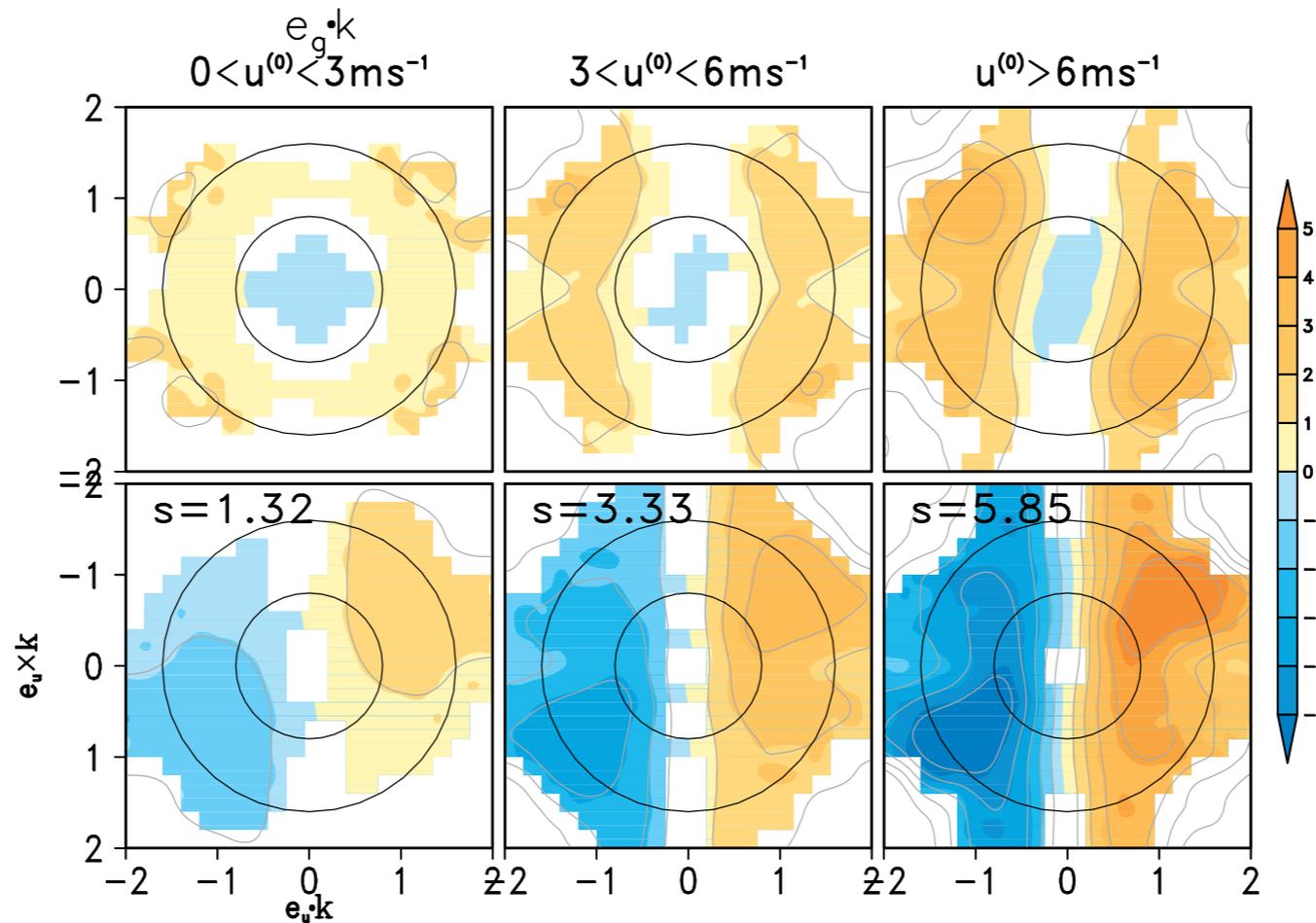


Transfer function comparison

Linear model

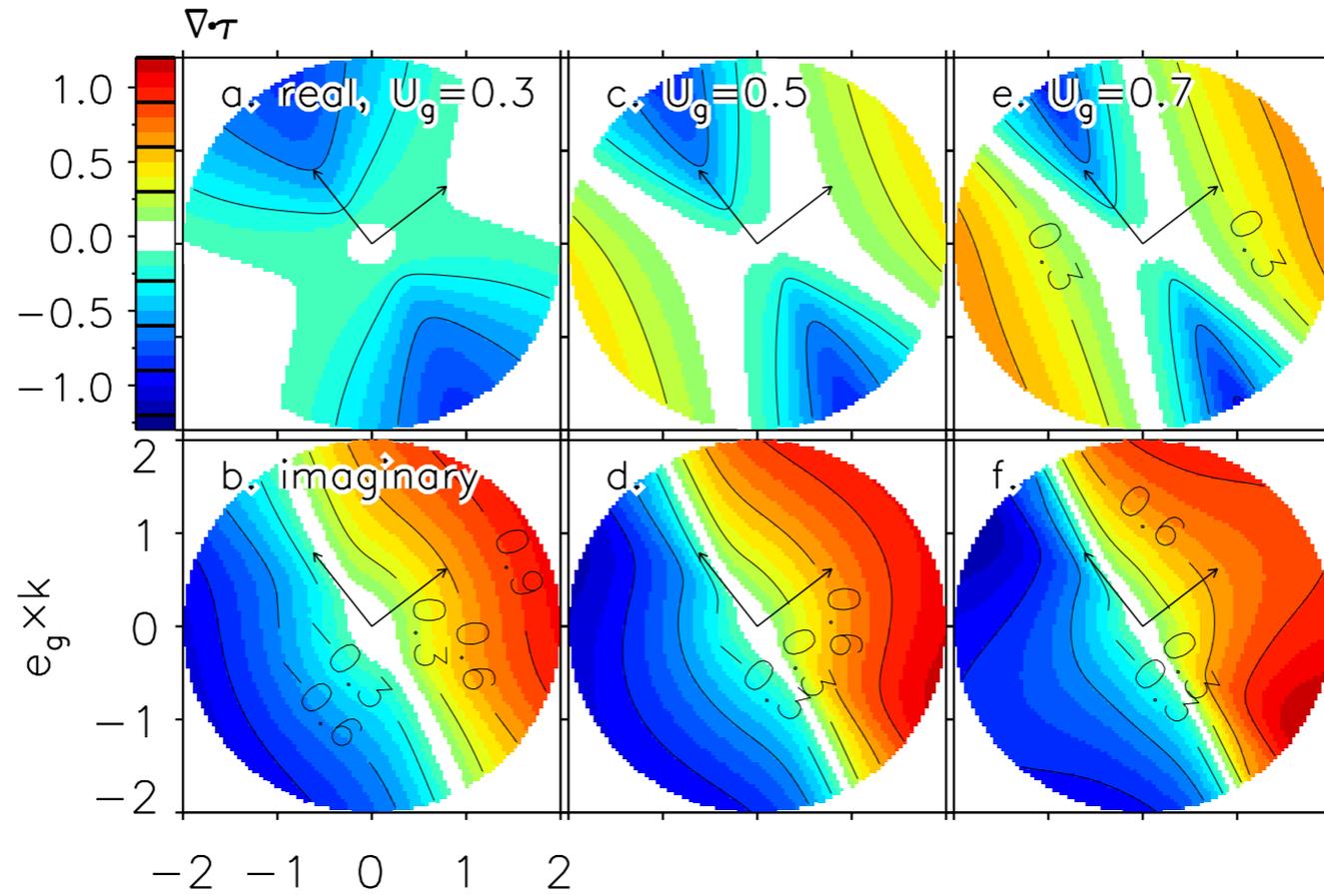


Observations

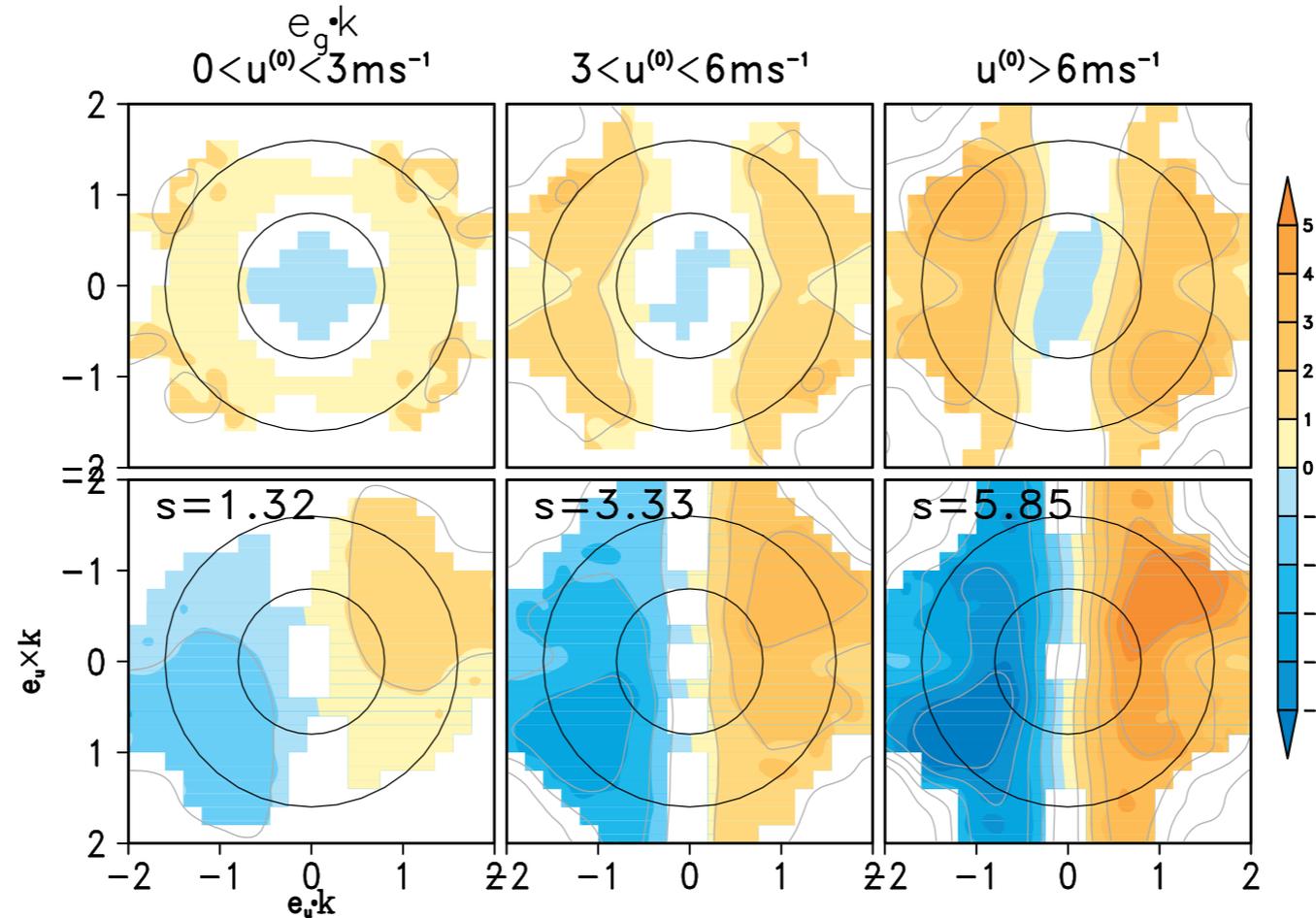


Transfer function comparison

Linear model



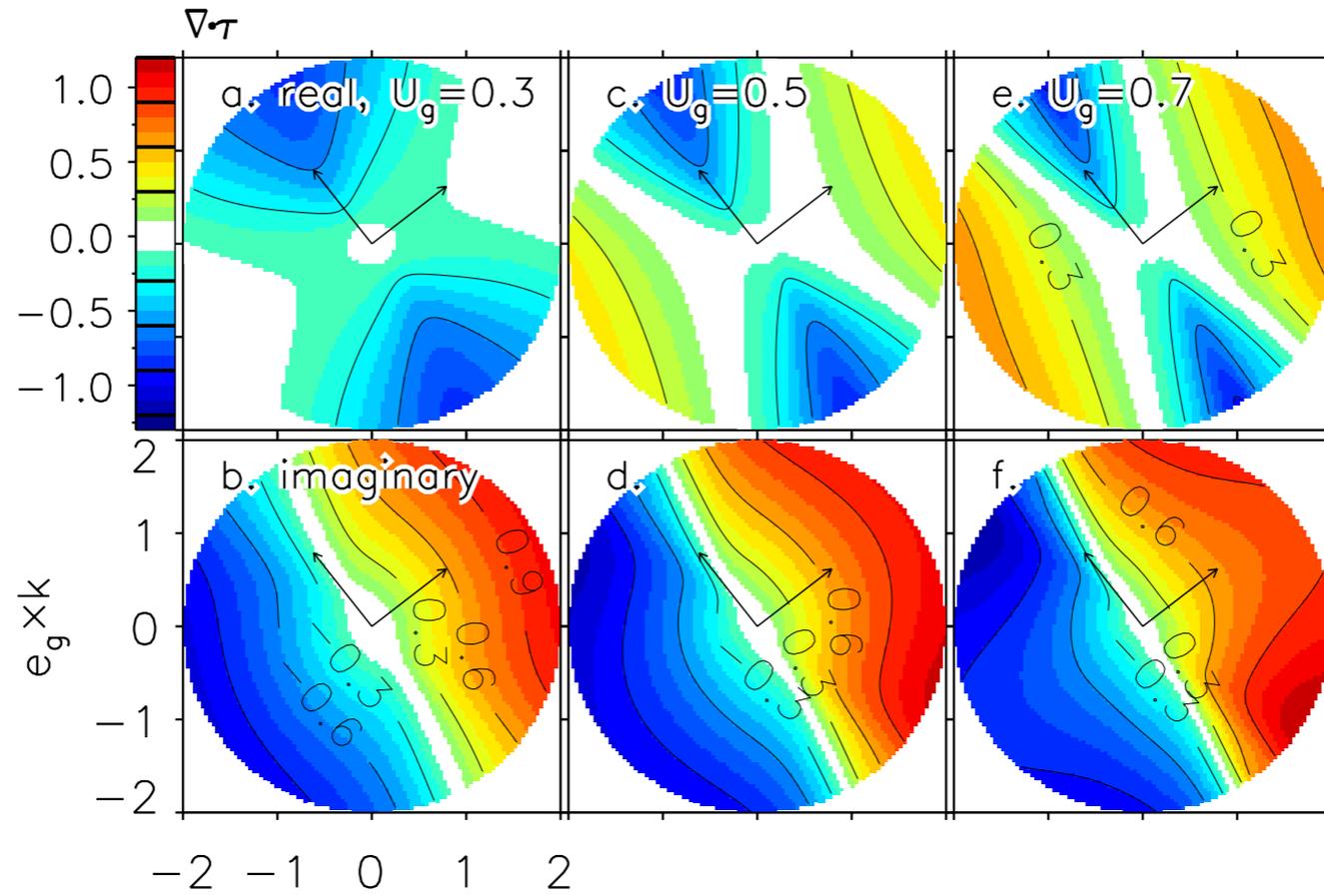
Observations



monthly
average

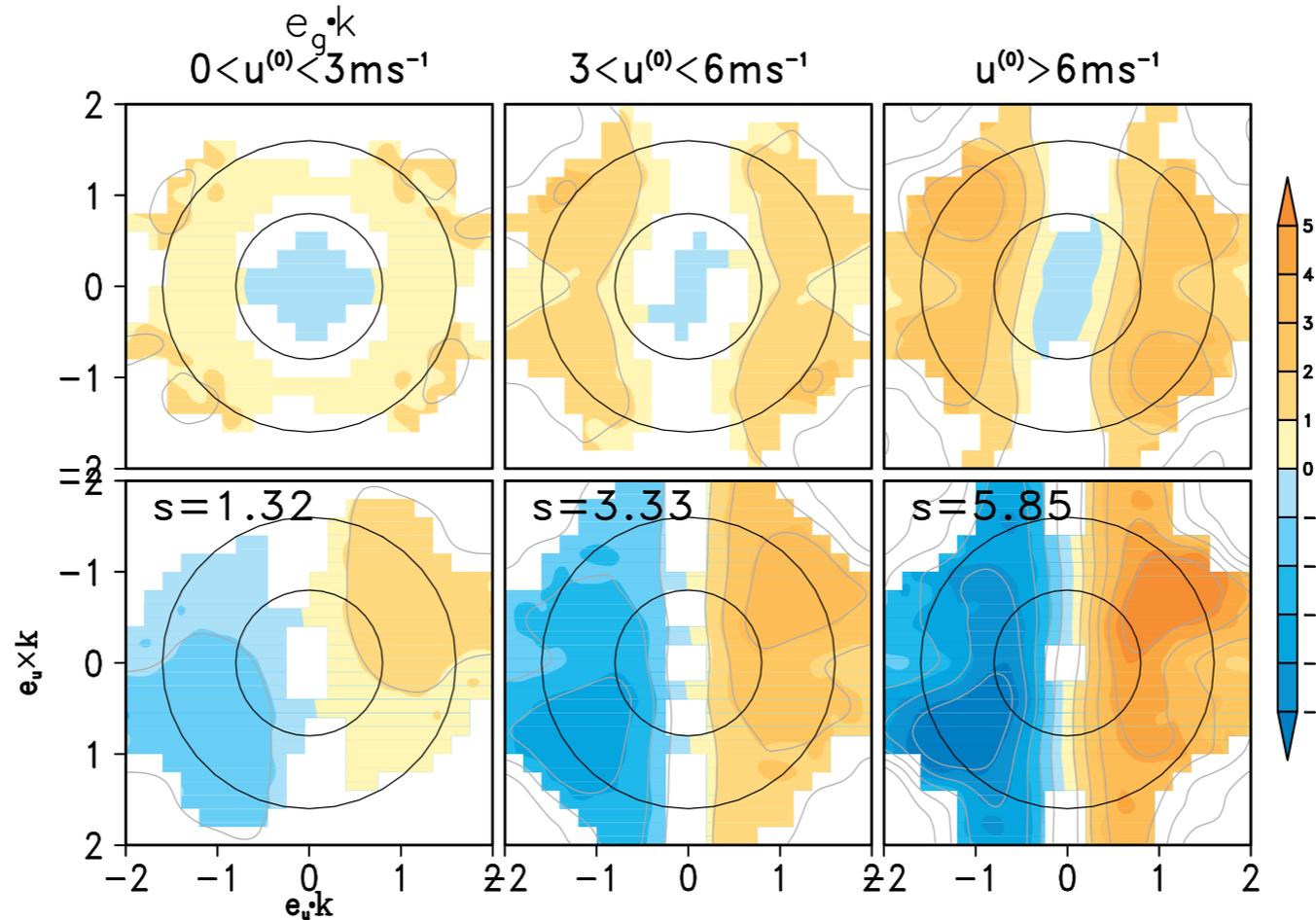
Transfer function comparison

Linear model



few inertial periods

Observations

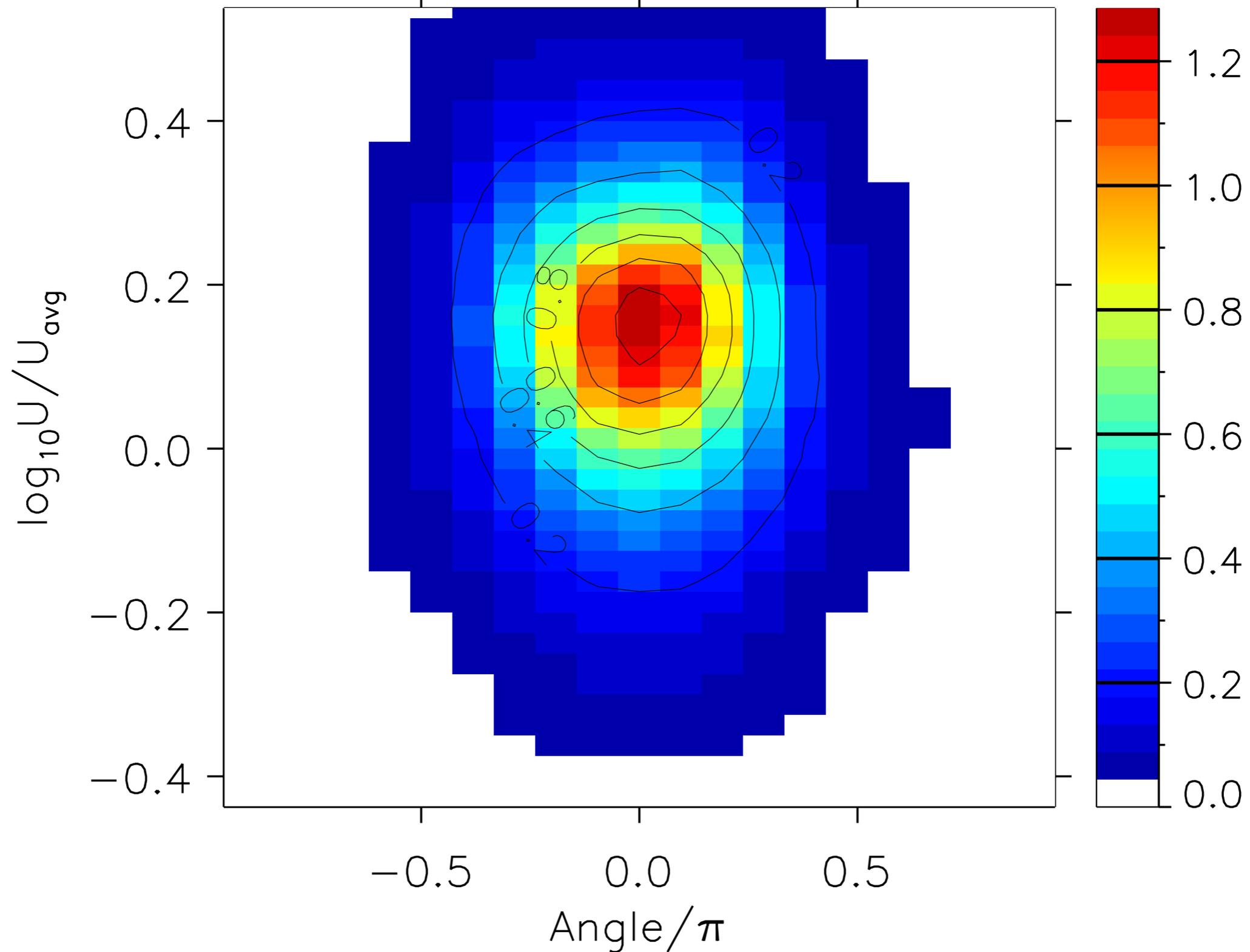


monthly average

Sub-monthly variations of background wind

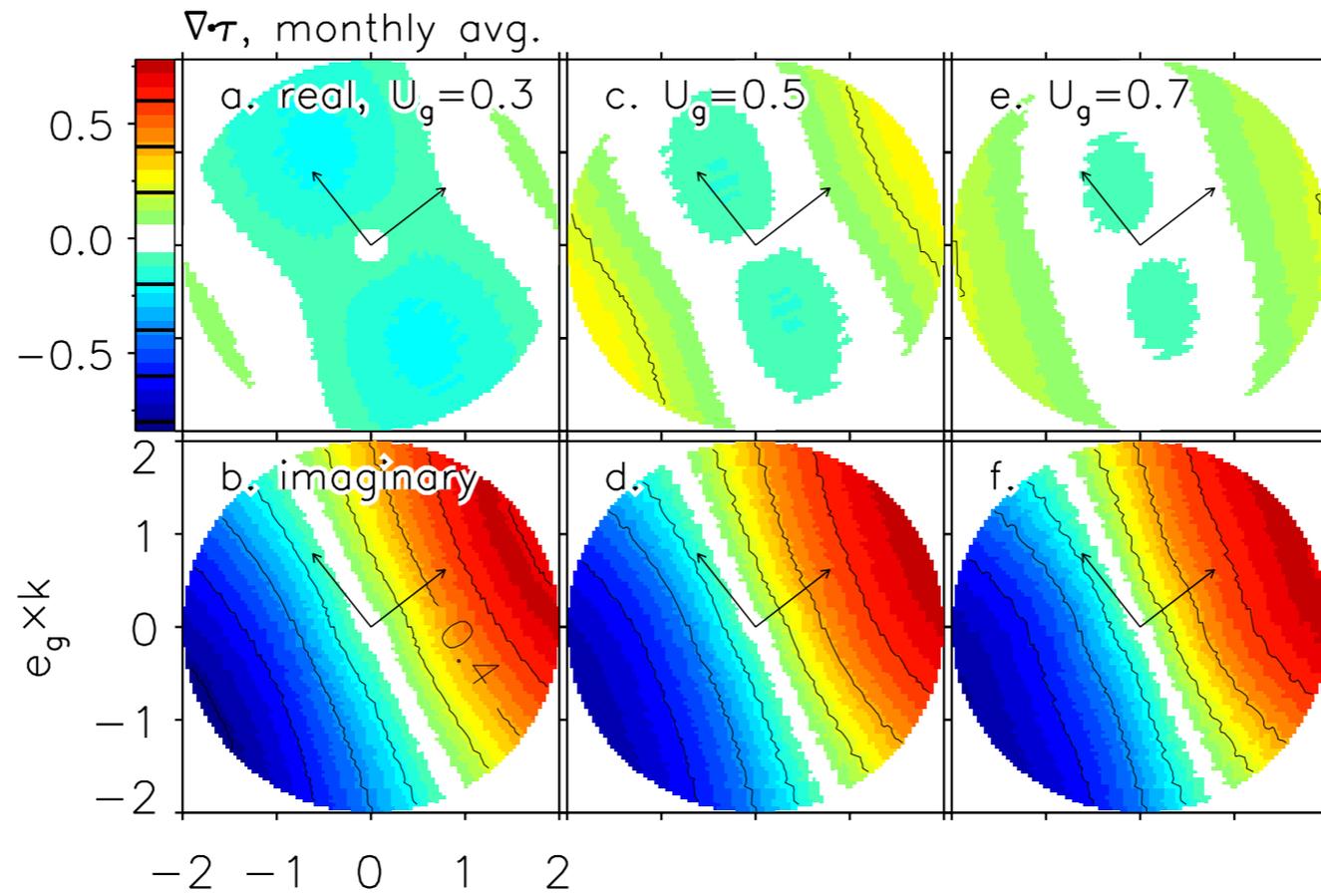
CAM5 multi year integration

$u^{(0)}$, submonthly variability, Agulhas, CAM5.HF.32



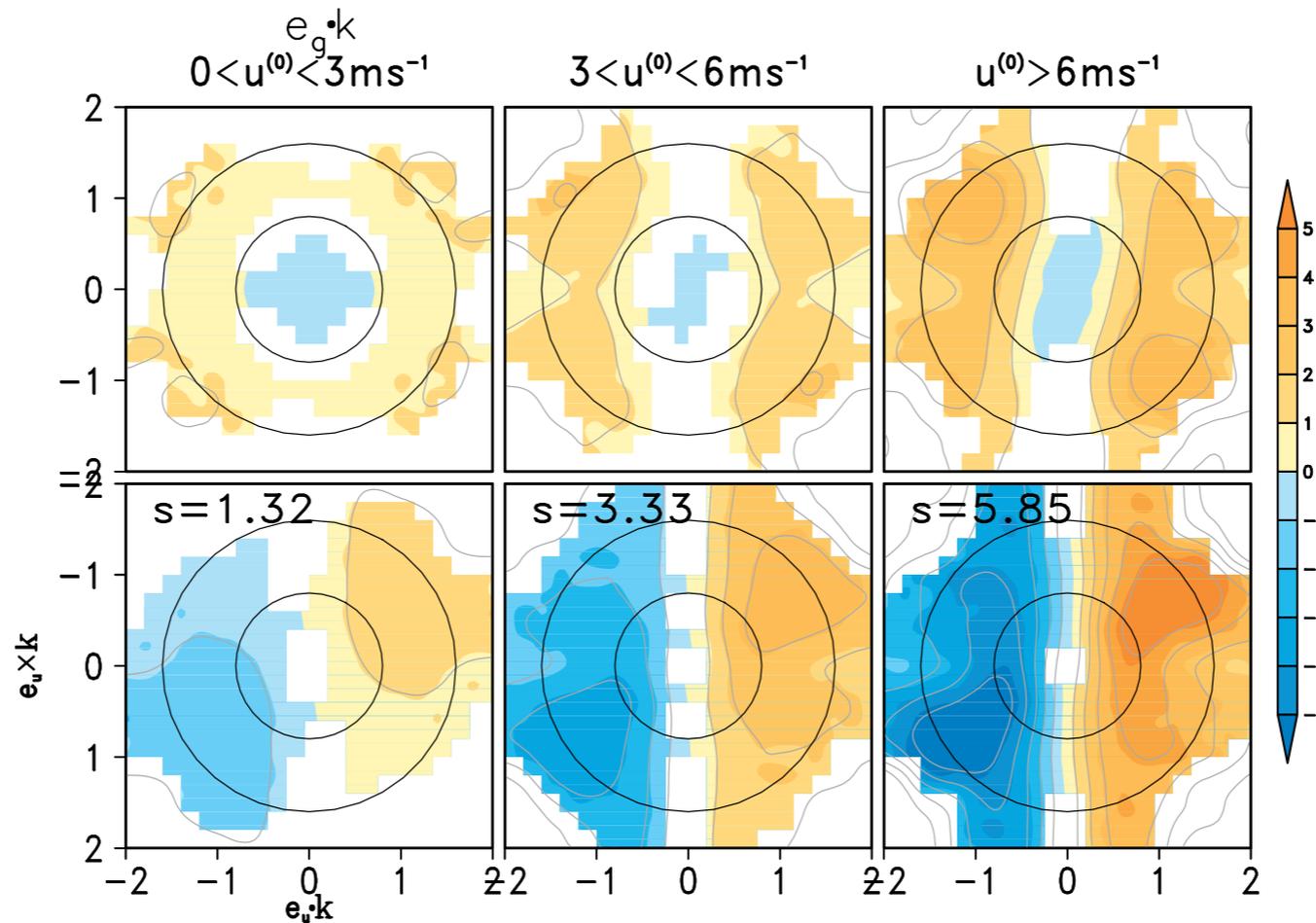
Transfer function comparison

Linear model



smoothed to
monthly average

Observations



monthly
average

Conclusions

- Transfer functions provide scale-dependent and lagged characterization of atmospheric boundary layer response to mesoscale sea surface temperatures
- Linear model suggests that the transfer function depends on the background wind speed and direction
- Monthly averages smooth out the structure of the transfer function
- Comparison of linear model with estimates of the transfer function is far from perfect but encouraging
- Testing with high resolution AMSR wind and SST products and adjustment of linear model parameters/physics under way