

# Covariability of Equatorial Winds and Sea Surface Height

Ted Durland, J. Tom Farrar, Dudley Chelton and Michael Schlax  
June 2nd, 2014

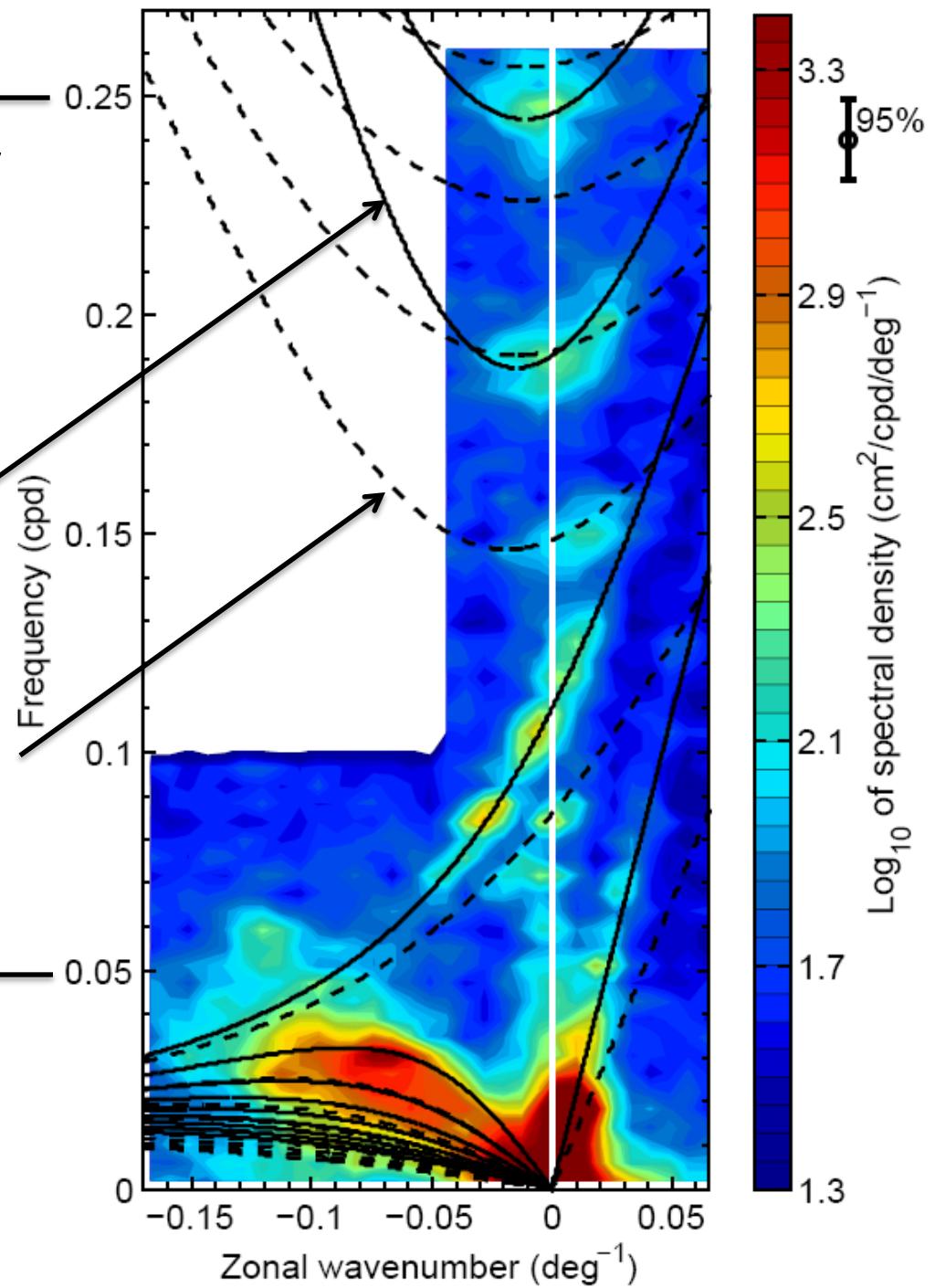


Supported by NASA OVWST program

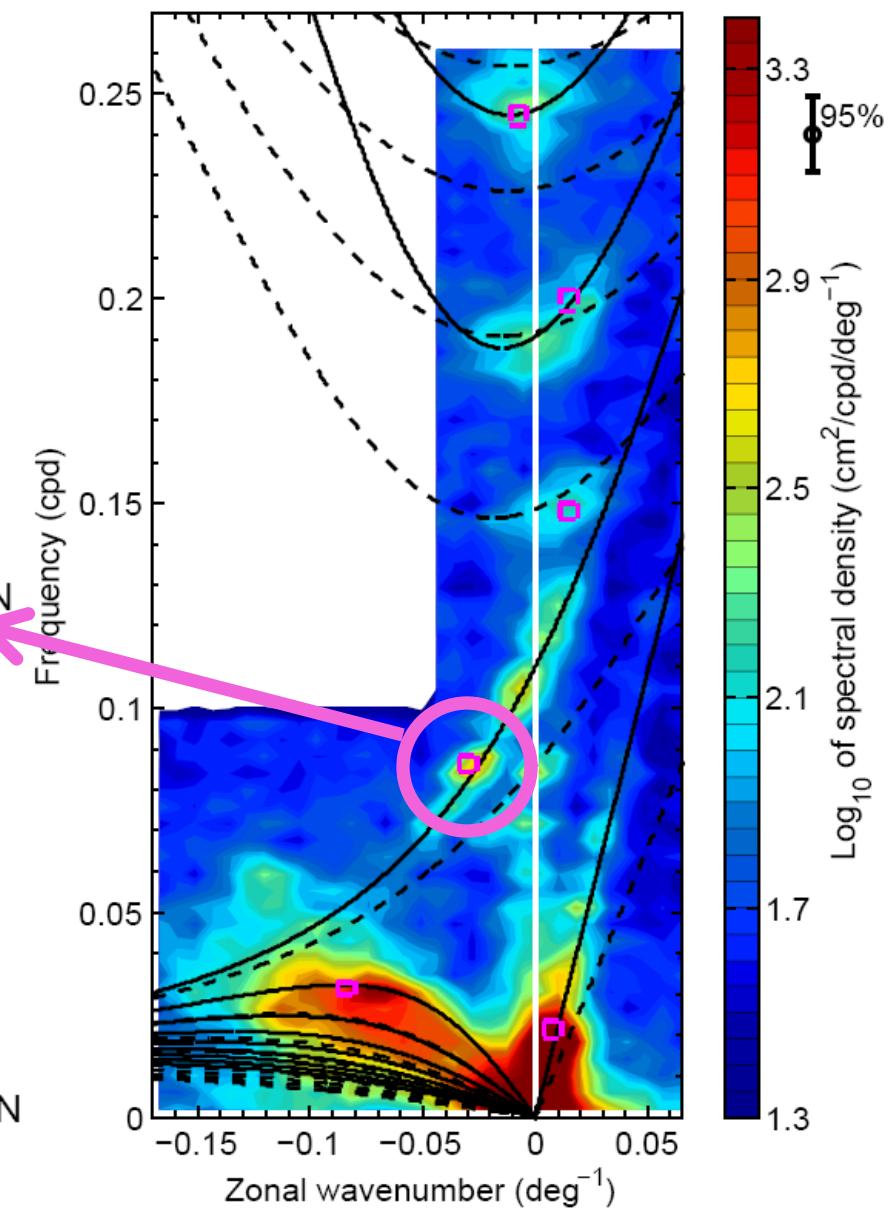
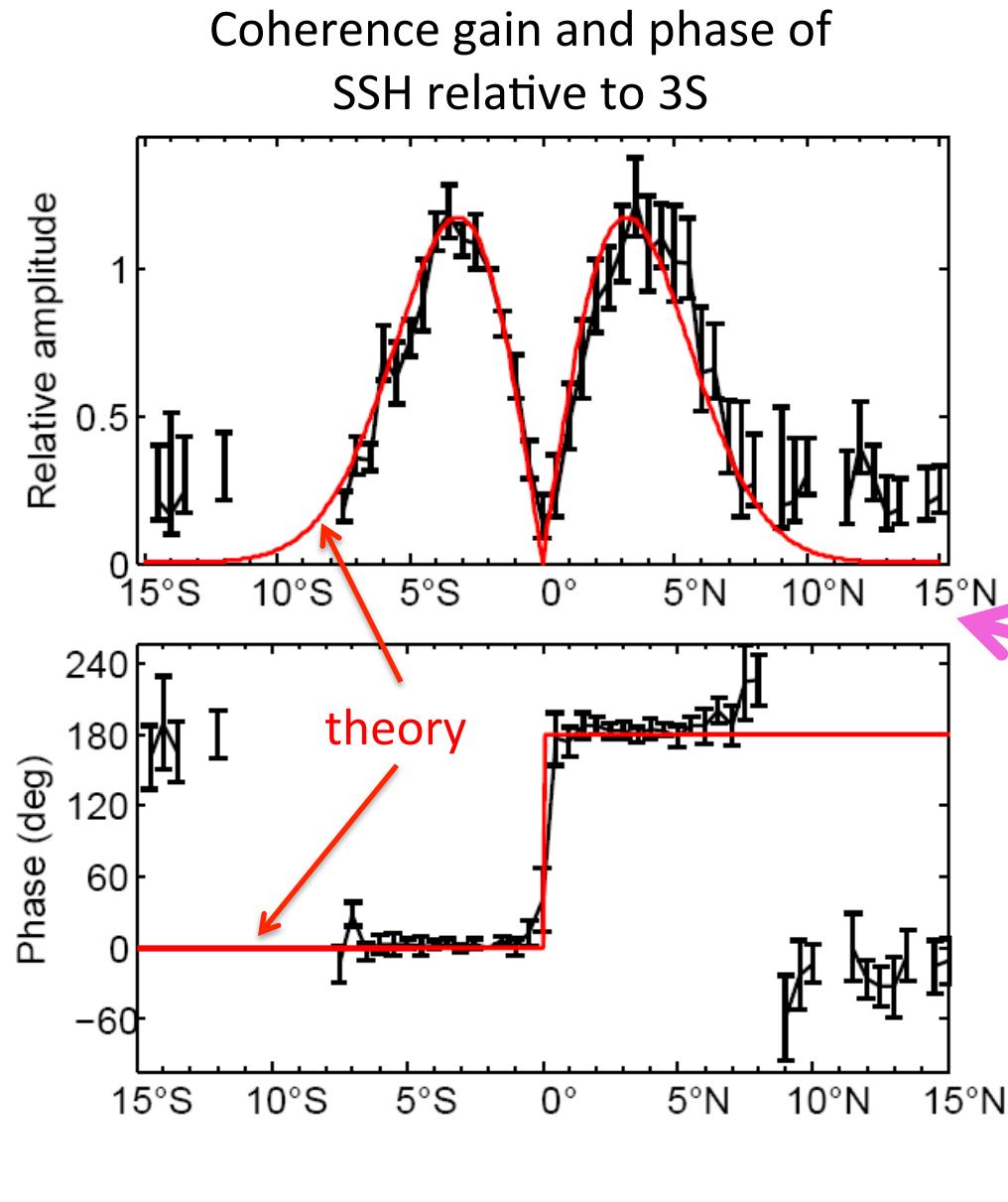
period  
4 day — 0.25  
Zonal-wavenumber/ frequency  
spectrum of SSH from  
along-track altimeter data  
in the equatorial Pacific Ocean  
(averaged 6°S-6°N, 2000-2005)

Solid dispersion curves:  
baroclinic mode 1  
  
Dashed dispersion curves:  
baroclinic mode 2

20 day — 0.05  
(Comparable spectra of TAO-TRITON  
dynamic height in Farrar and Durland,  
*JPO*, Nov, 2012)



# The 11-day, 33° Yanai wave (1<sup>st</sup> baroclinic mode)...

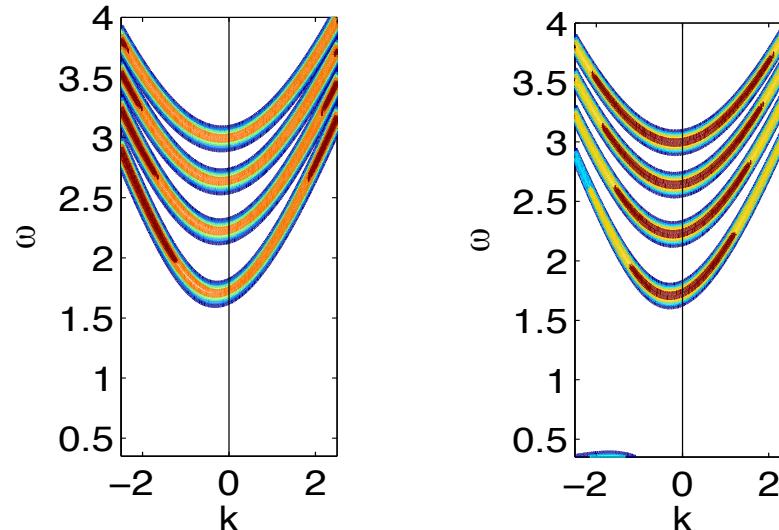


# Spectral solutions for forcing that is broadband in zonal wavenumber, frequency and meridional structure

log10  
meridionally  
integrated  
pressure (SSH)  
variance

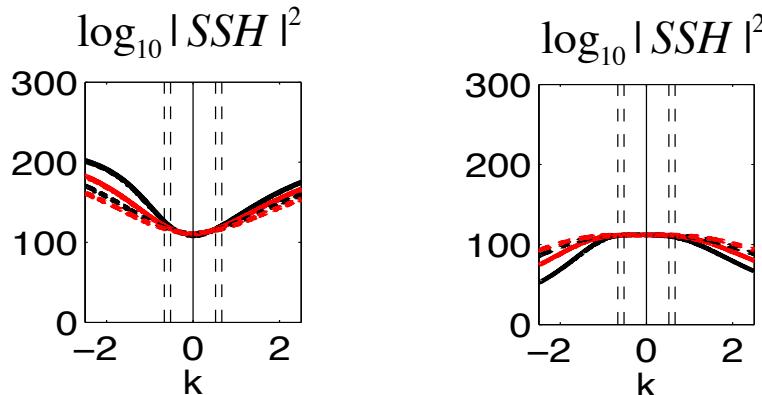
- || |  $m = 1$
- || |  $m = 2$
- · - |  $m = 3$
- · - |  $m = 4$

Wavenumber-frequency spectra (Non-dimensional  $\omega, k$ )  
Response to  $\tau^x$       Response to  $\tau^y$



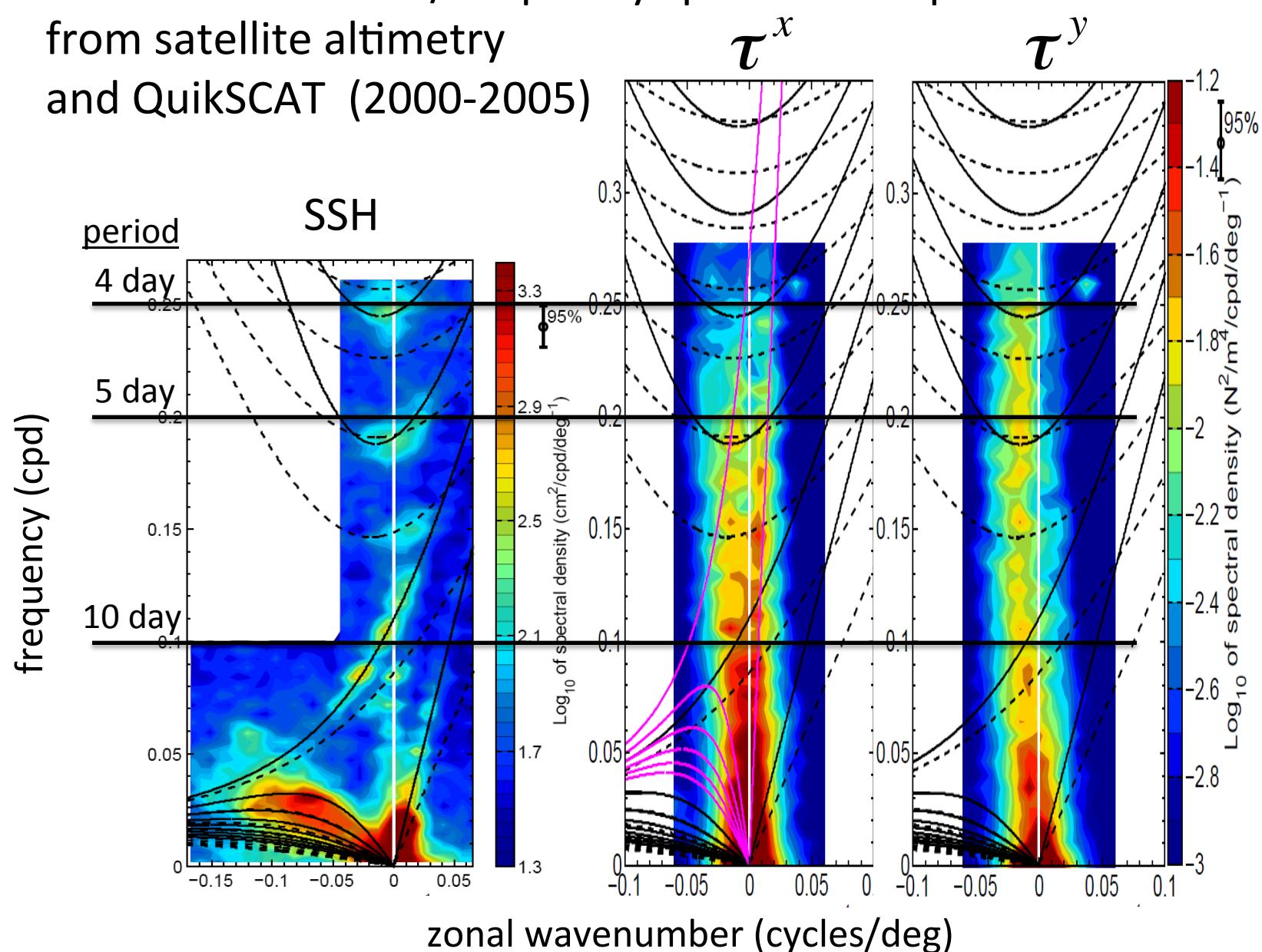
Infinite domain,  
Broadband forcing,  
No peaks where  
group velocity  
vanishes

Along dispersion curves (at max resonance)

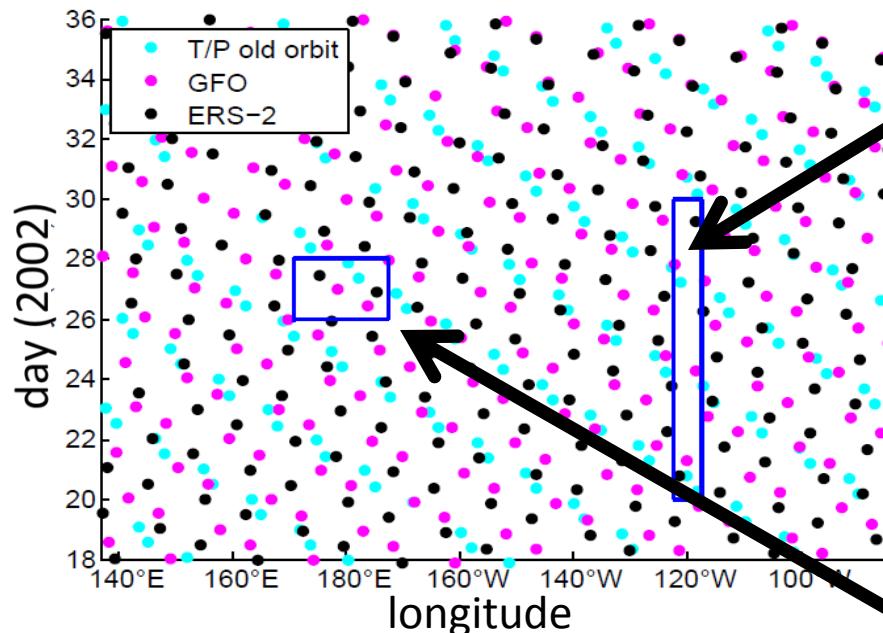


Durland and Farrar,  
*JPO*, 2012

zonal-wavenumber/frequency spectra for equatorial Pacific  
from satellite altimetry  
and QuikSCAT (2000-2005)

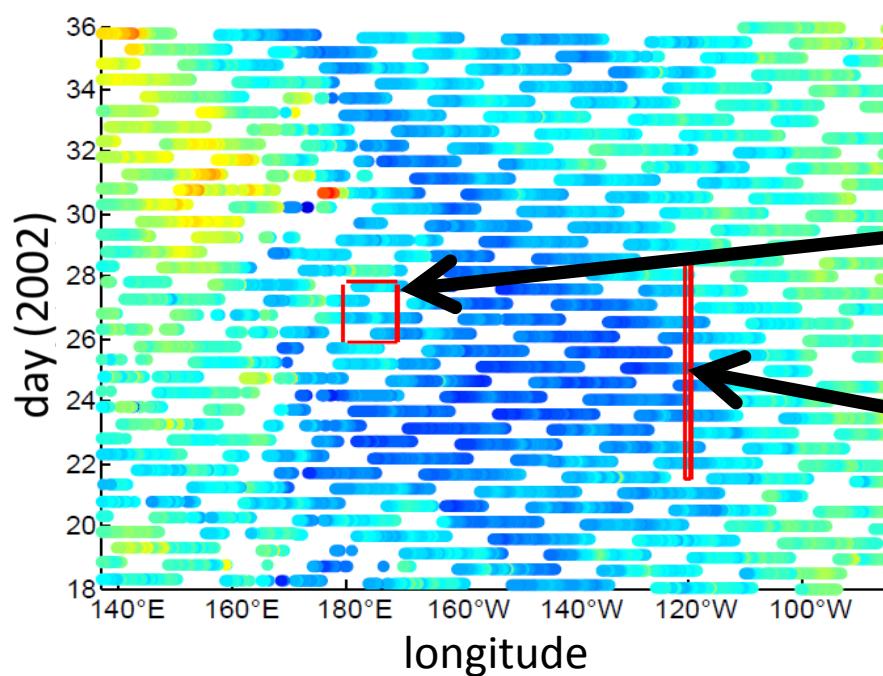


Longitude-time  
altimetry sampling  
at  $3.5^{\circ}\text{N}$



5° by 10 days  
similar to scales  
represented in  
AVISO gridded  
products

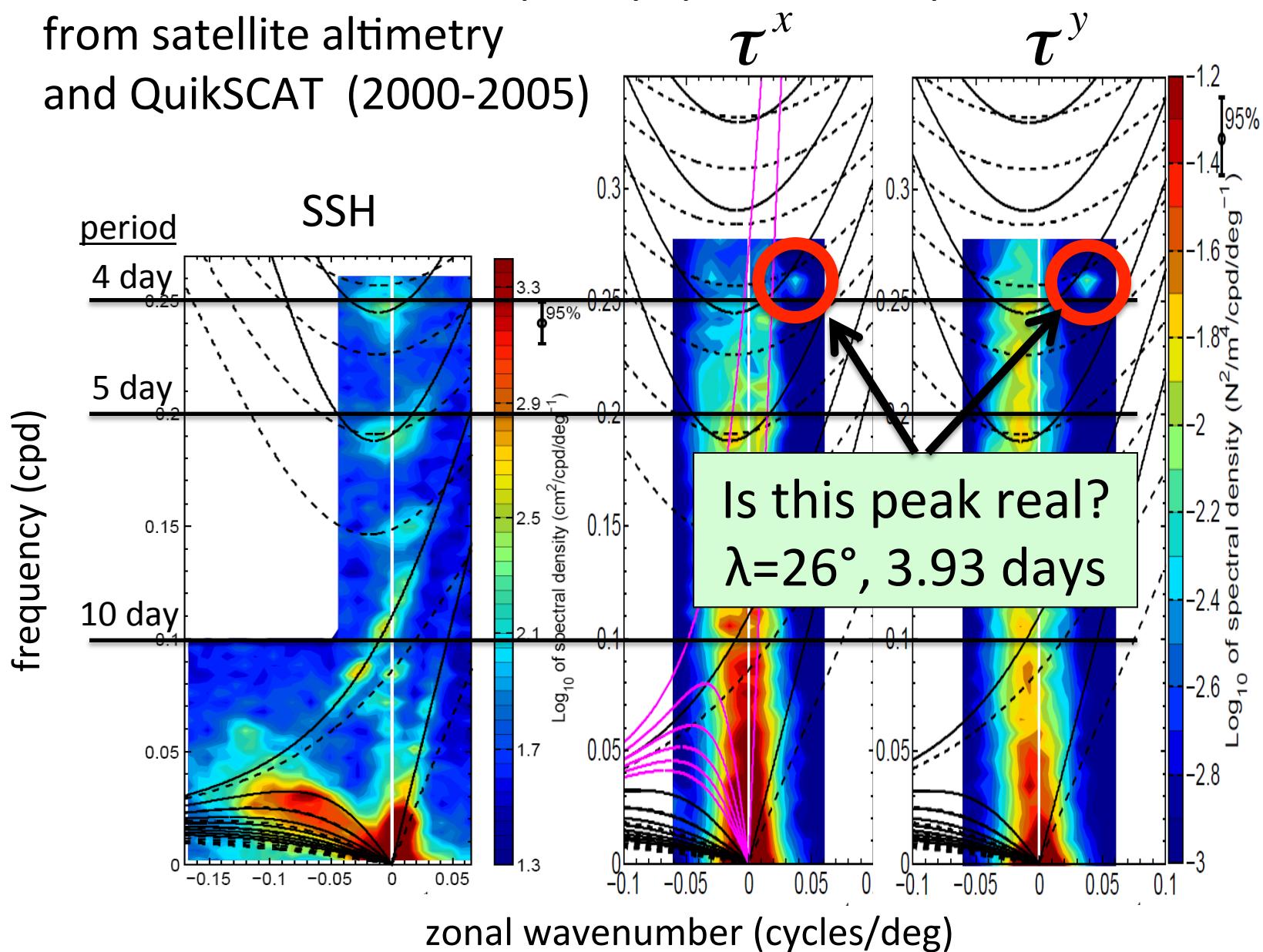
Longitude-time  
QuikSCAT sampling  
at  $3.5^{\circ}\text{N}$



Farrar sampling  
15° by 2 days  
~ 4 day “Nyquist  
period”  
8° x 2 days:

0.25° x 7 days:  
similar to weekly  
average of  
QuikSCAT data

zonal-wavenumber/frequency spectra for equatorial Pacific  
from satellite altimetry  
and QuikSCAT (2000-2005)



## Resolution Analysis

(Farrar, Schlax, Chelton and Durland, in prep.)

- 1) How can we decide if a spectral band is *resolved*? Specifically, how well can we determine the Fourier coefficients from the data?
  - 2) How can we tell if unresolved wavenumbers and frequencies of the true variability might contaminate a particular band of our estimate?
- 

$$\hat{\alpha} = \Gamma y = \Gamma(F^{-1}\alpha + n)$$

Vector of estimated Fourier coefficients      Vector of sampled data      Vector of “true” Fourier coefficients  
linear operations performed on raw data to estimate Fourier coefficients      noise      inverse Fourier transform

Assuming the relative phases of the true Fourier coefficients are random,

$$\langle |\hat{\alpha}|^2 \rangle = \mathbf{R}|\alpha|^2, \text{ where } \mathbf{R} = |\Gamma F^{-1}|^2 \text{ and } | |^2 \text{ is element-wise}$$

This “resolution matrix”  $\mathbf{R}$  is the basic result

$$\langle |\hat{\mathbf{a}}|^2 \rangle = \mathbf{R}|\mathbf{a}|^2$$

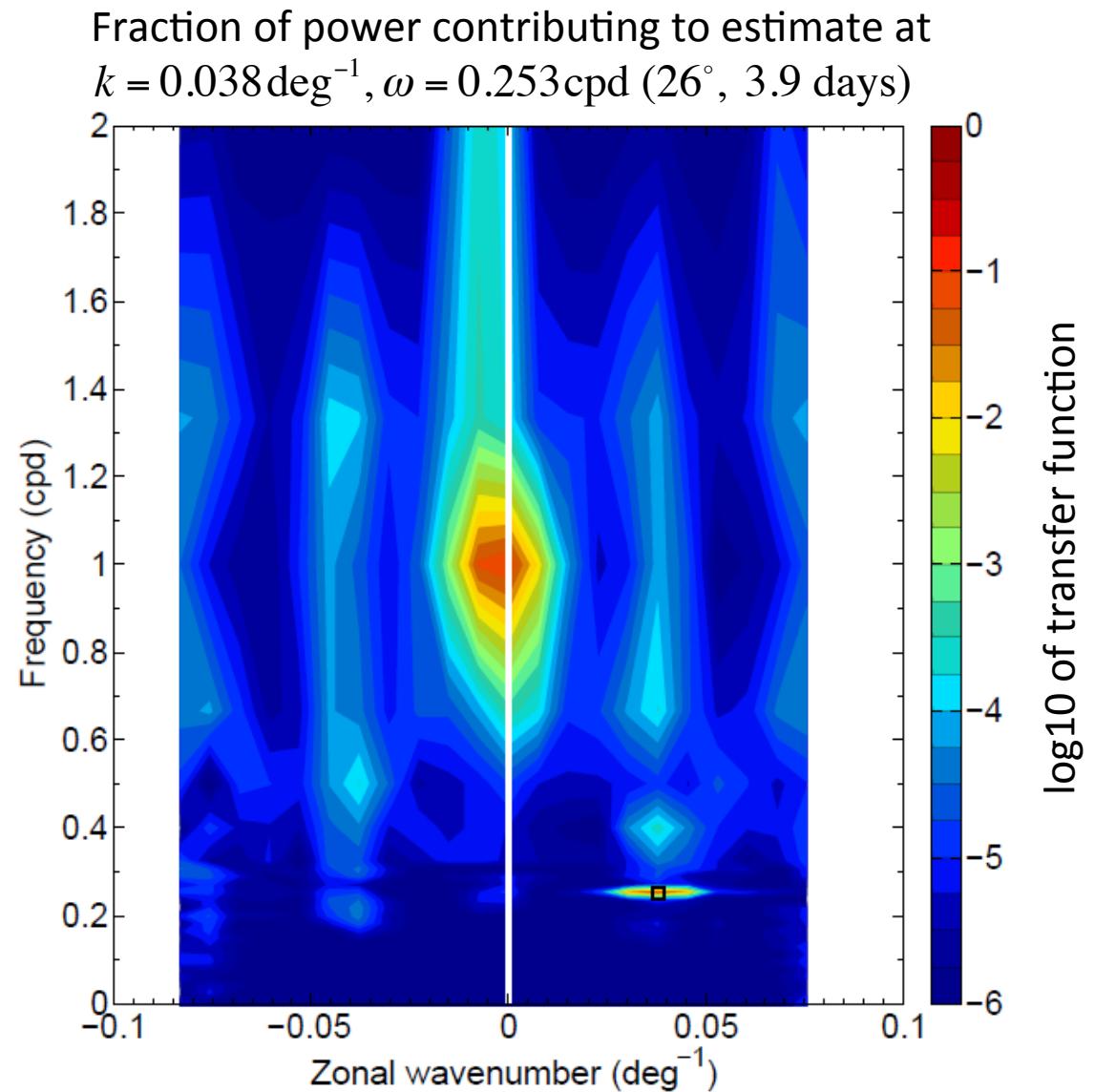
Vector of estimated spectral values      Vector of true spectral values

R gives us two powerful tools:

- A row of  $\mathbf{R}$  tells us how each frequency of the true spectrum influences our estimate at a particular frequency (a “transfer function”)
- A column of  $\mathbf{R}$  tells us the spectrum we would estimate if the true spectrum contained only a single frequency (a “response function”)

Row of resolution matrix: tells us the parts of the spectrum contributing to the “alias” peak in our estimate

Our estimate responds strongly to diurnal signals of global scale that propagate westward, passing ~10% of the variance (i.e., ~30% of the amplitude)



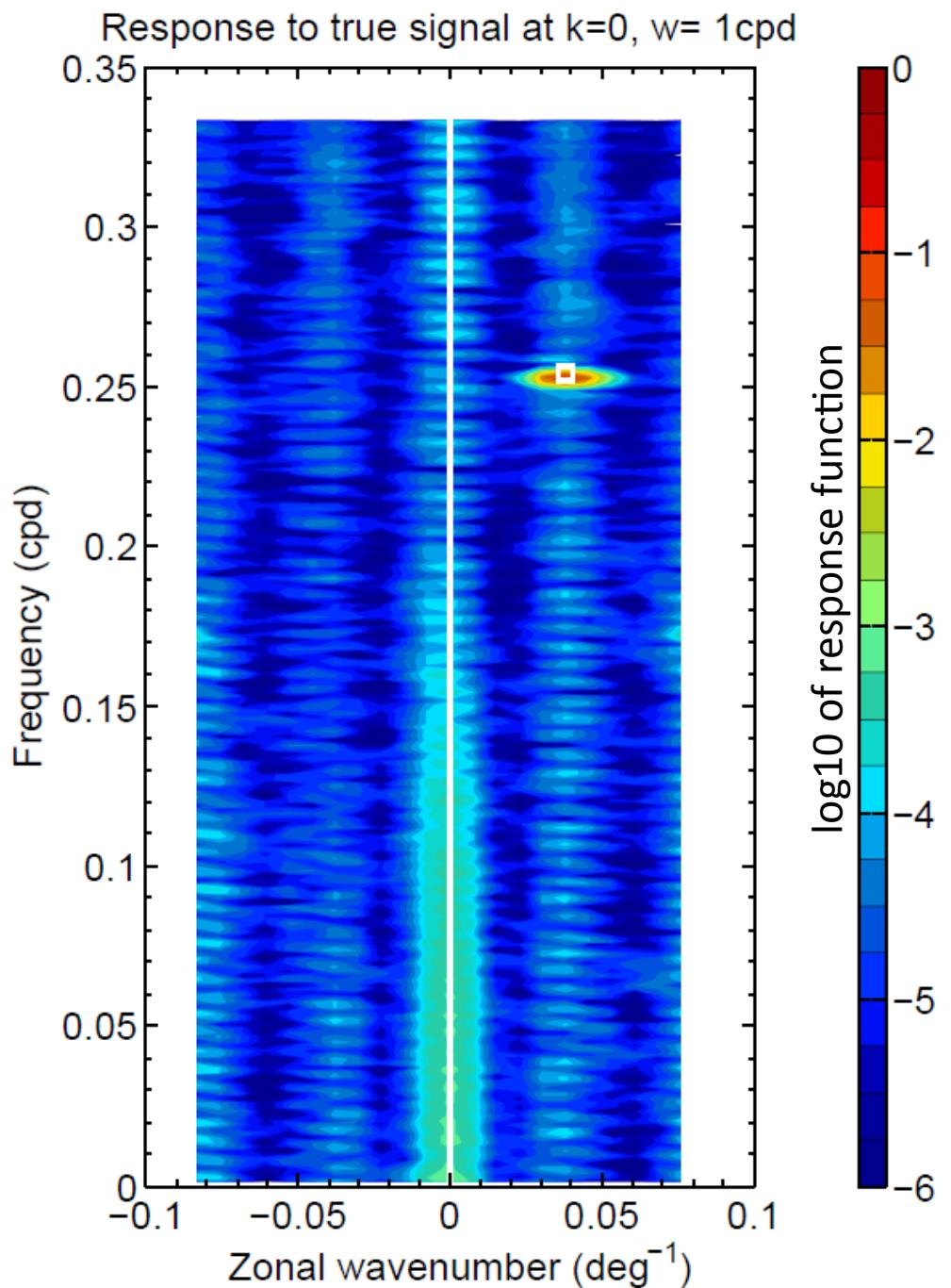
Column of resolution matrix:  
tells us how each band in our  
estimate responds to a  
global-scale diurnal signal

What other bands are  
contaminated by the diurnal cycle?

Besides the major alias peak we  
could see in the spectrum, the  
worst contamination is at basin  
scales and periods exceeding 7  
days, but the alias is  $\sim$ 100 times  
weaker.

The total variance of the aliasing to  
low frequencies is comparable,  
though, because there are lots  
affected bands ( $\sim$ 110 bands).

**About 50% of the variance of the  
global-scale diurnal cycle appears  
spuriously in my gridded field!**

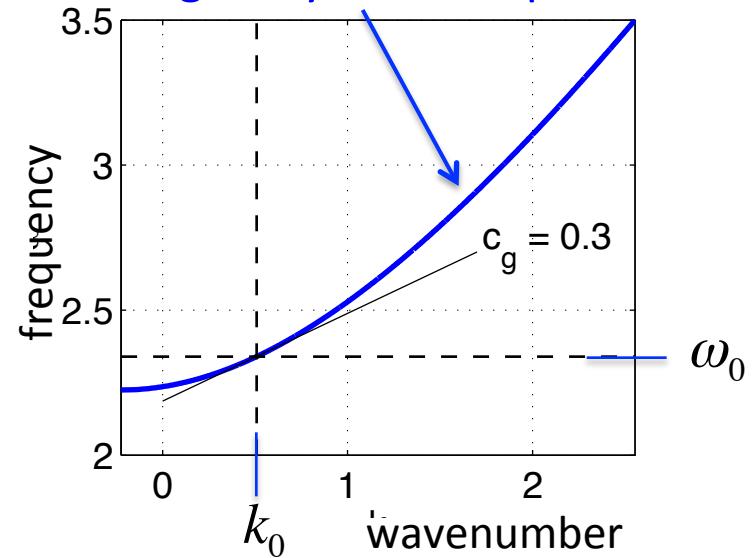


## To Do

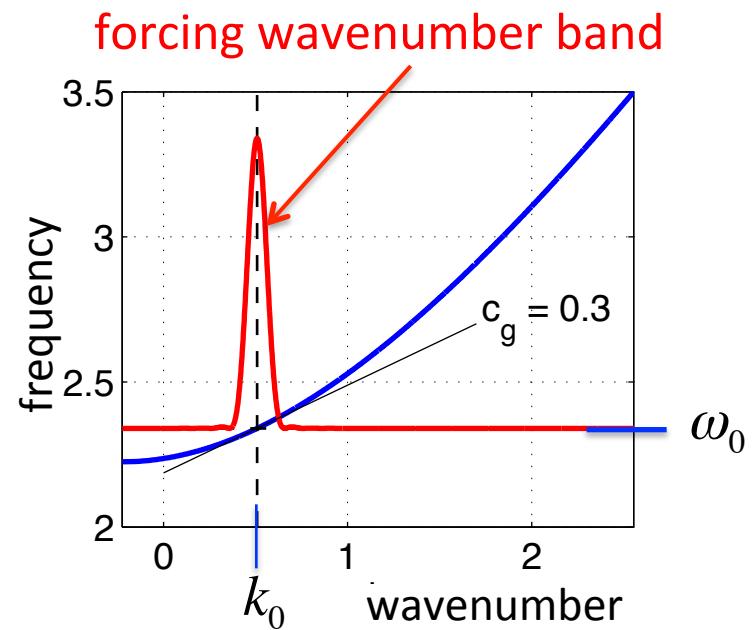
- 1) Contribute to understanding of zonal, meridional and temporal structure of diurnal and semidiurnal surface winds over oceans, with methods used to characterize oceanic modes.
- 2) Use resolution matrix to conduct comprehensive examination of possible artifacts in existing gridded products.

# Group velocity effect when forcing is isolated in space

inertia-gravity wave dispersion curve

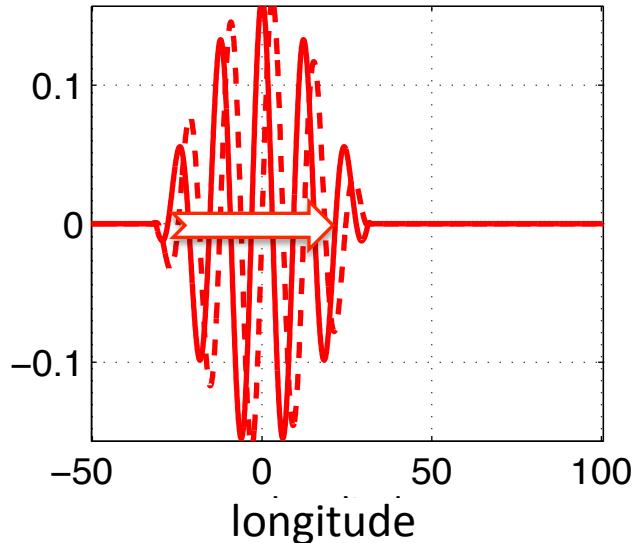
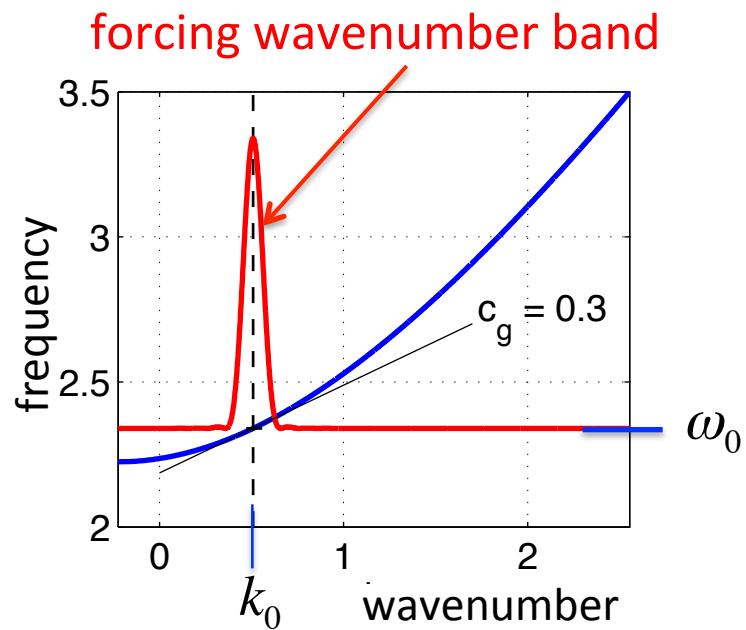


## Group velocity effect when forcing is isolated in space

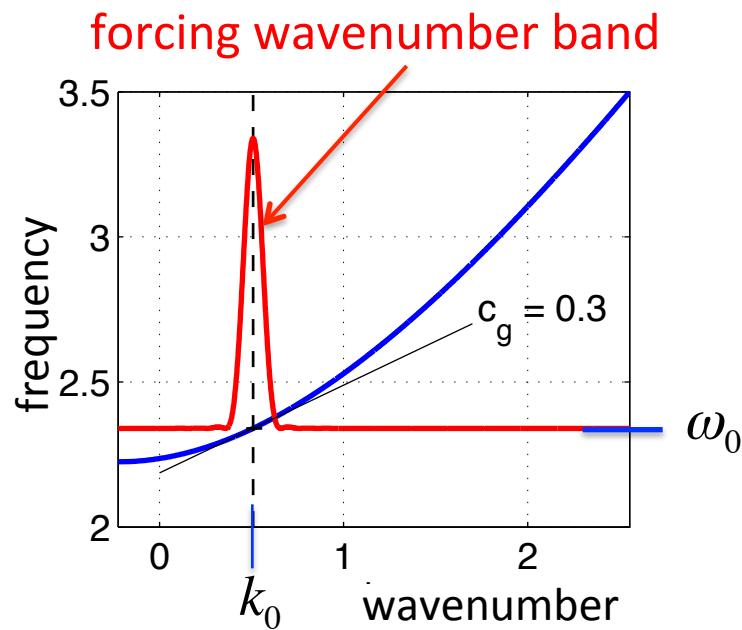


## Group velocity effect when forcing is isolated in space

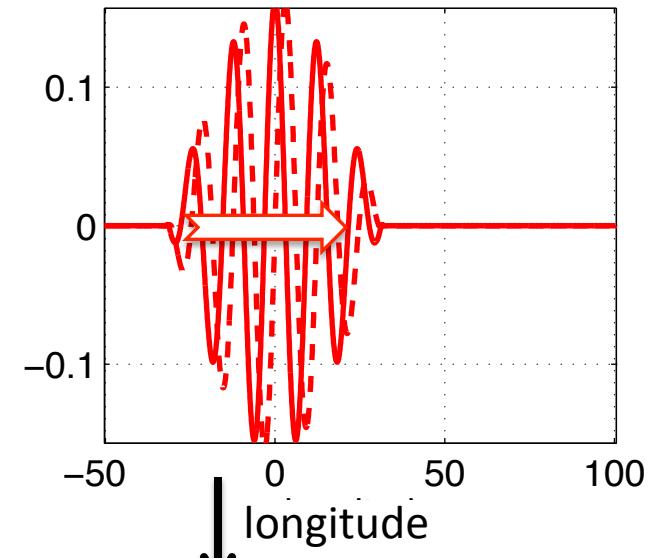
eastward propagating forcing,  
confined in longitude



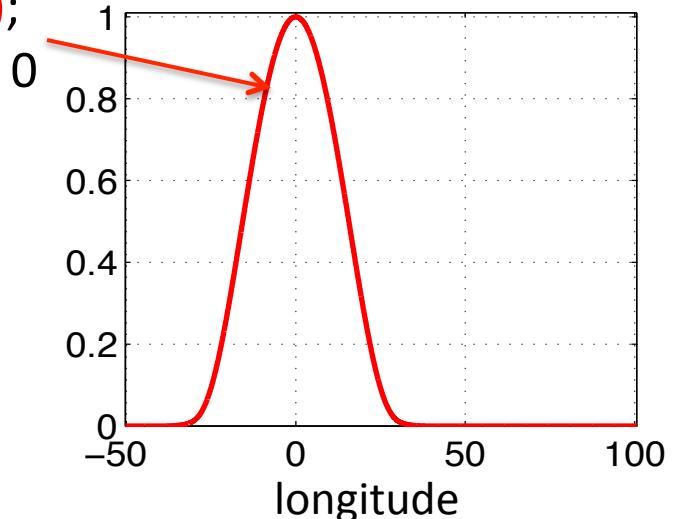
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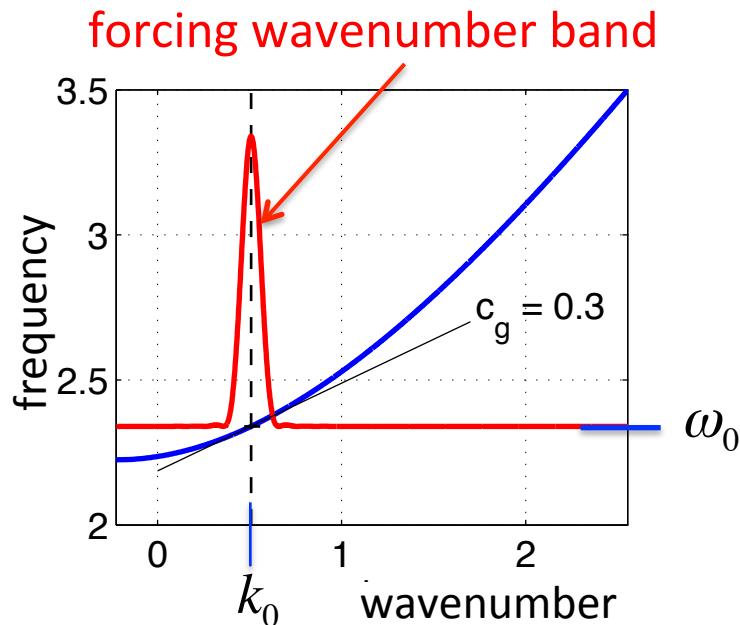
eastward propagating forcing,  
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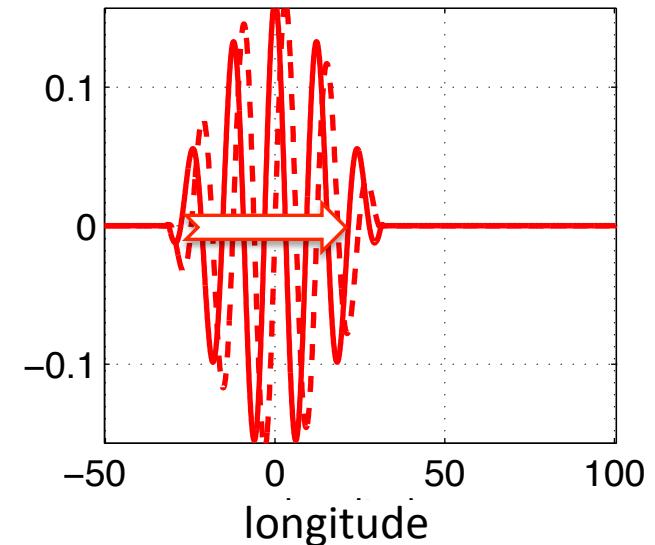
normalized forcing of energy equation,  $F(x)$ ;  
also oceanic energy density,  $E(x)$ , when  $c_g = 0$



## Group velocity effect when forcing is isolated in space



eastward propagating forcing,  
confined in longitude

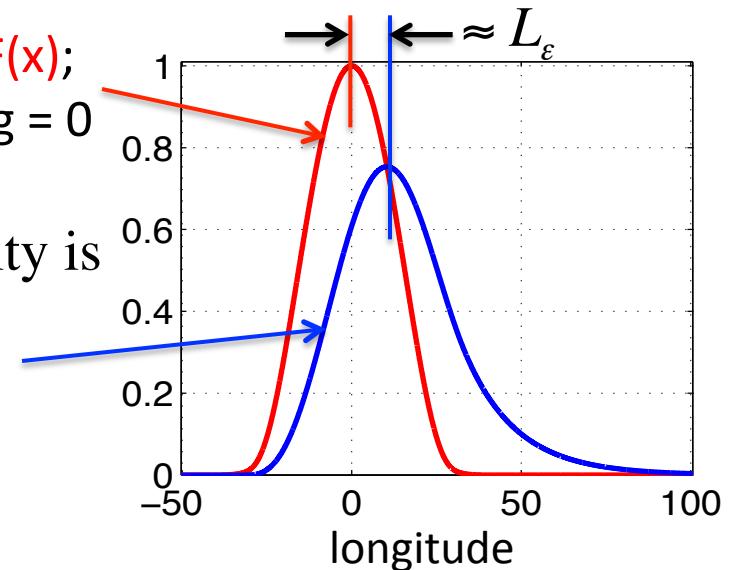


normalized forcing of energy equation,  $F(x)$ ;  
also oceanic energy density,  $E(x)$ , when  $c_g = 0$

when  $c_g \neq 0$ , oceanic response energy density is

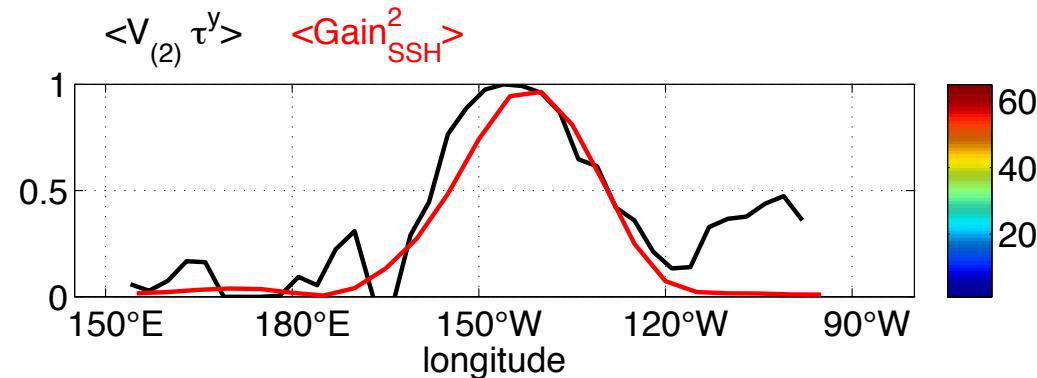
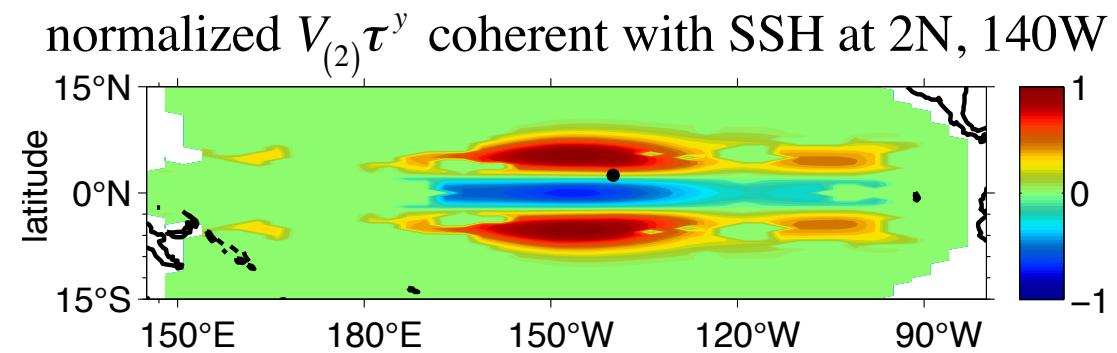
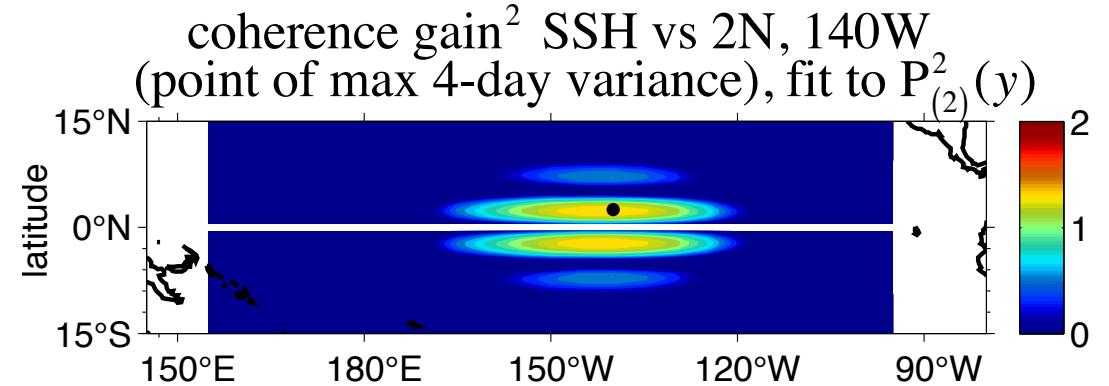
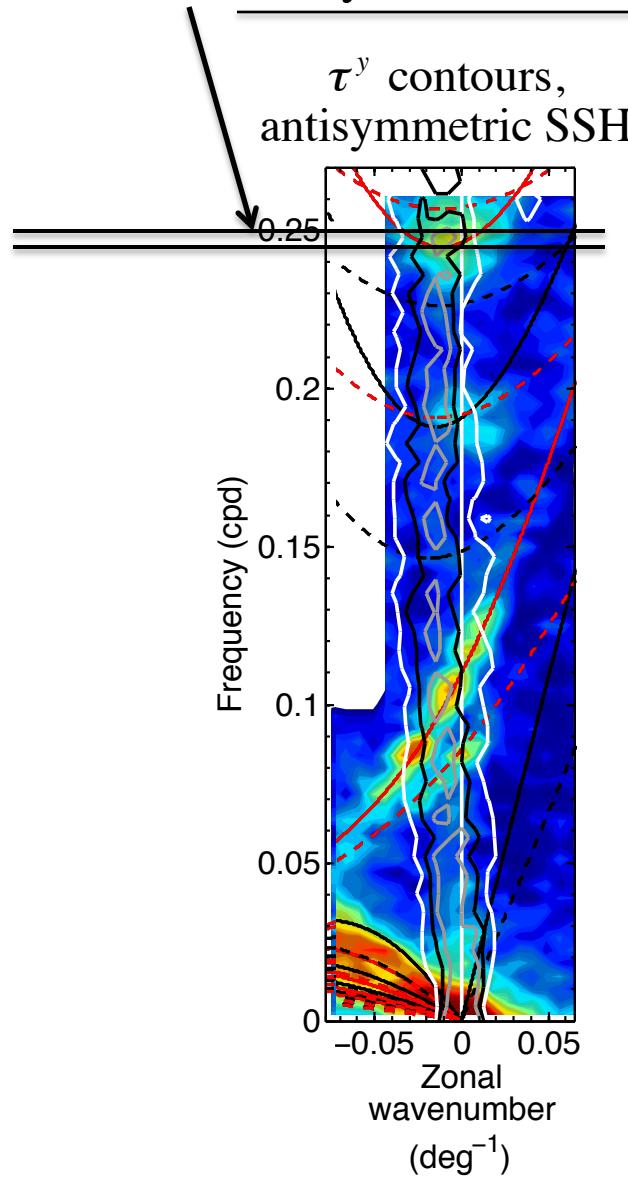
$$E(x) \approx \propto L_\varepsilon^{-1} \int_0^\infty F(x - \xi) \exp(-\xi / L_\varepsilon) d\xi,$$

$$T_\varepsilon = L_\varepsilon / c_g$$



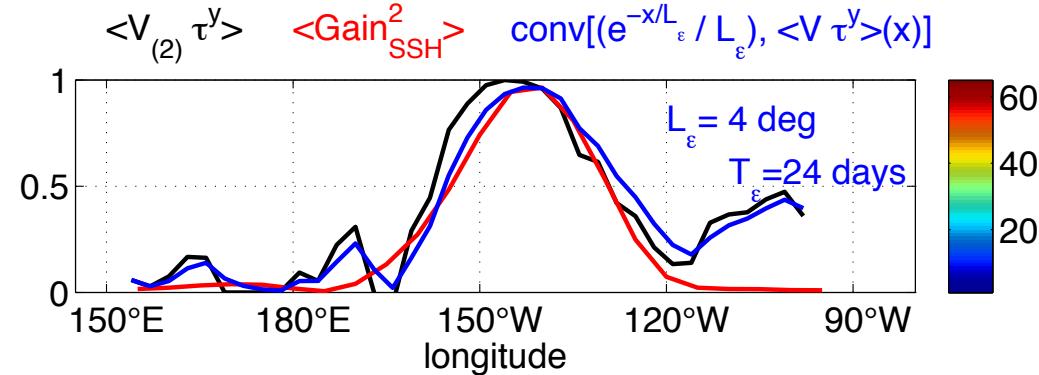
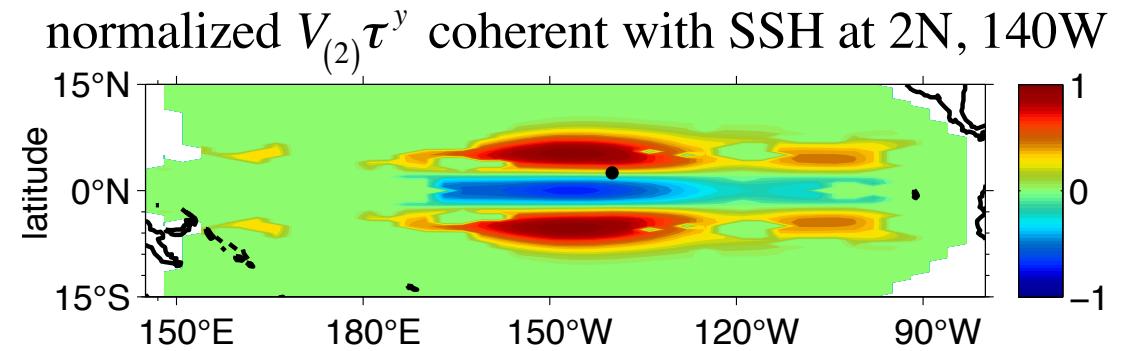
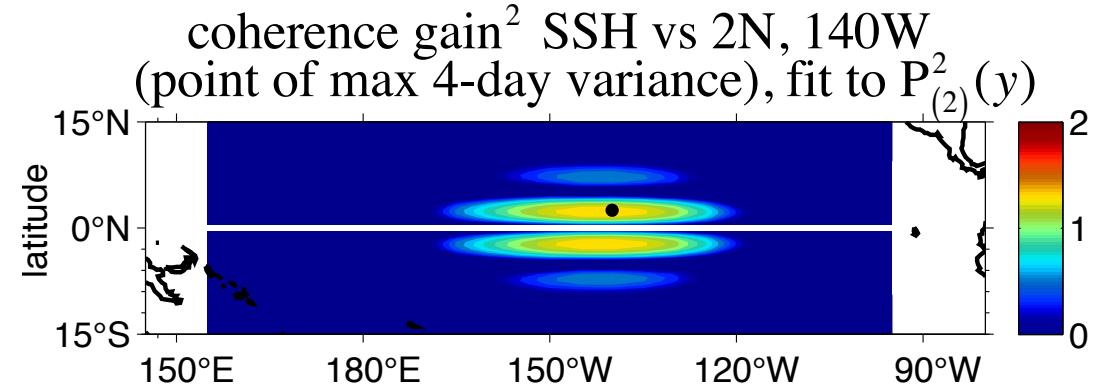
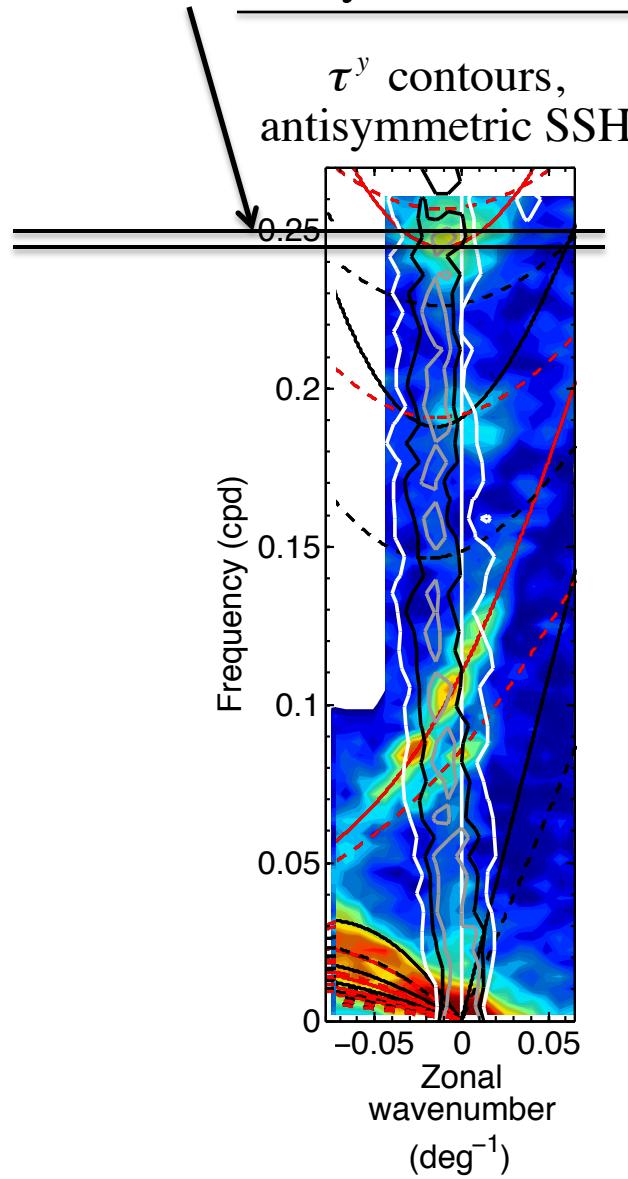
# SSH – wind stress comparisons using altimetry & QuikSCAT

4.0–4.1 day band, bc model1, merid mode 2 wave:  $c_{gx} = 0.2 \text{ deg/day}$  (0.2 m/s)



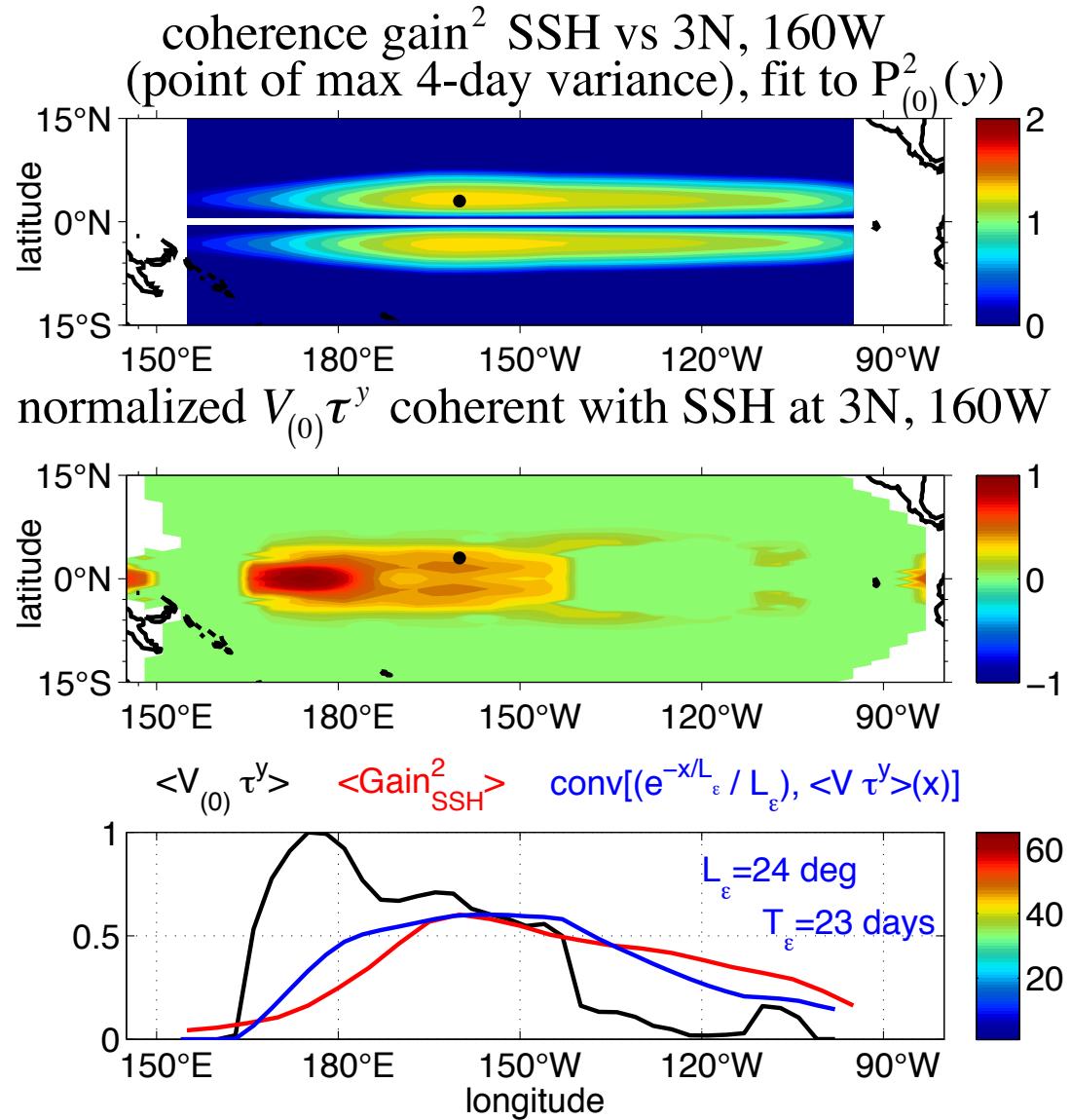
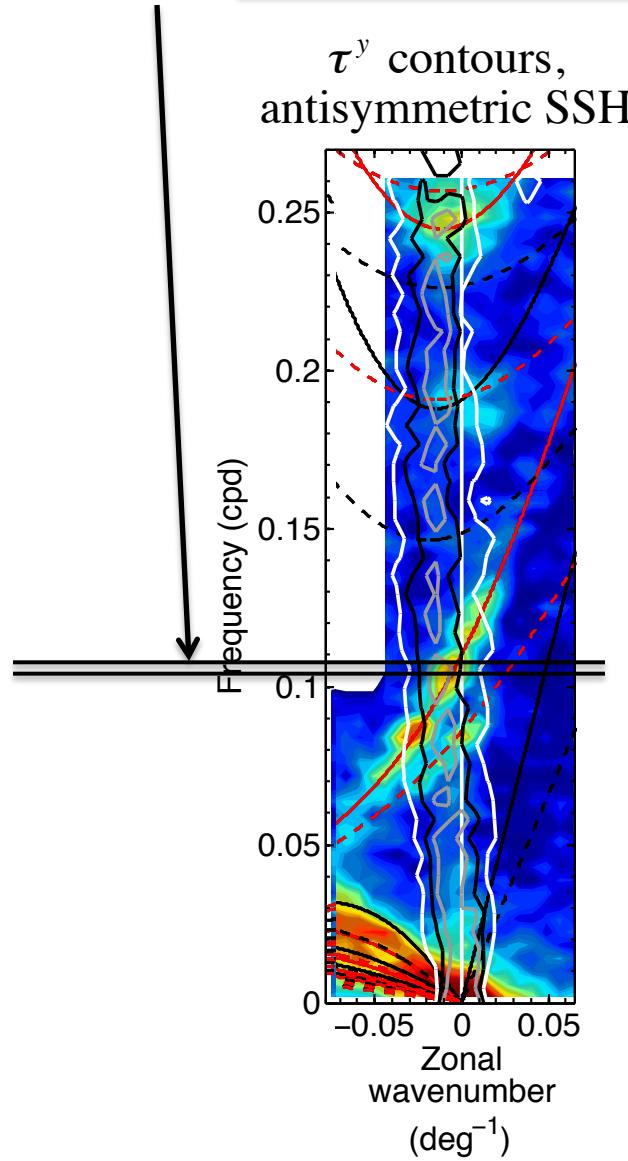
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# SSH – wind stress comparisons using altimetry & QuikSCAT

9.3–9.7 day band, bc mode1, merid mode 0 wave:  $c_{gx} = 1.0 \text{ deg/day}$  (1.3 m/s)



## To Do Summary

- 1) Contribute to understanding of zonal, meridional and temporal structure of diurnal and semidiurnal surface winds over oceans.
- 2) Use resolution matrix approach to conduct comprehensive examination of possible aliases in existing gridded products.
- 3) Improve spatial wind stress – SSH response model:
  - a) vertical modeling – set up of baroclinic modes,
  - b) inclusion of “minor forcing” – subtleties of phase variability,
  - c) extension from single frequency to frequency band.
- 4) Extend climatological analysis of equatorial SSH & wind stress variance to describe spatial & temporal (seasonal, intraseasonal, interannual) patterns.

## Motivation for studying these high-frequency waves

“Many of the biological processes included in the ecosystem model respond to environmental fluctuations with time scales between 1 and 10 days, which are not typically resolved by basin-to global-scale circulation models.”

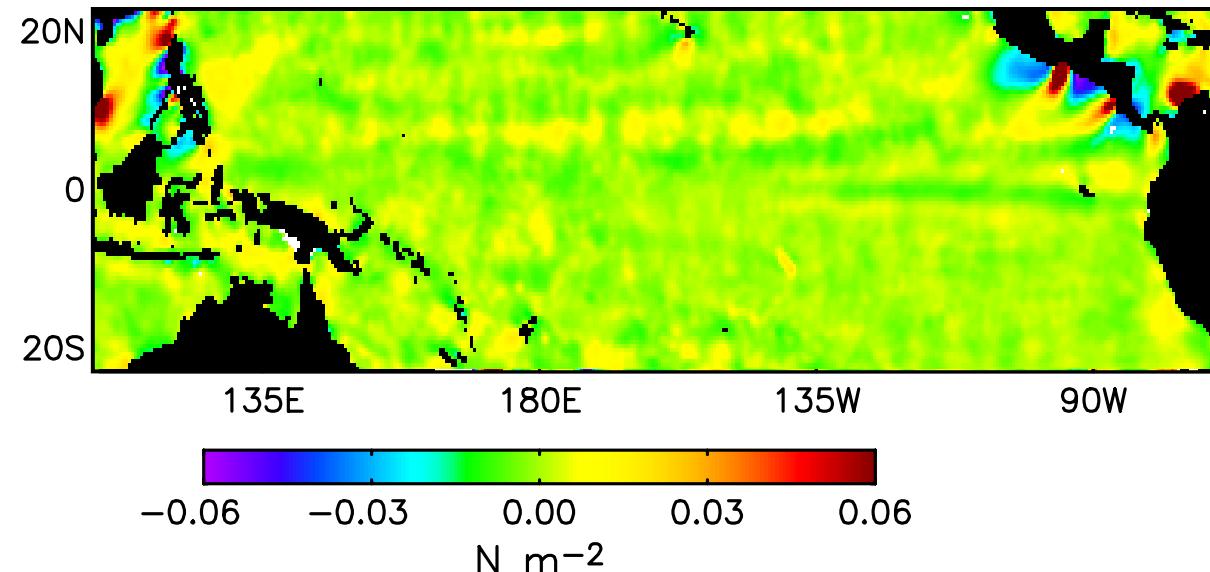
“Physical control of biological processes in the central equatorial Pacific Ocean”

Freidrichs and Hofmann, DSR, 2001.



High passed wind stress in the equatorial Pacific (11/2002 - 2/2003)  
from Maloney and Chelton, 2006

QuikSCAT



ECMWF  
and  
Reynolds

