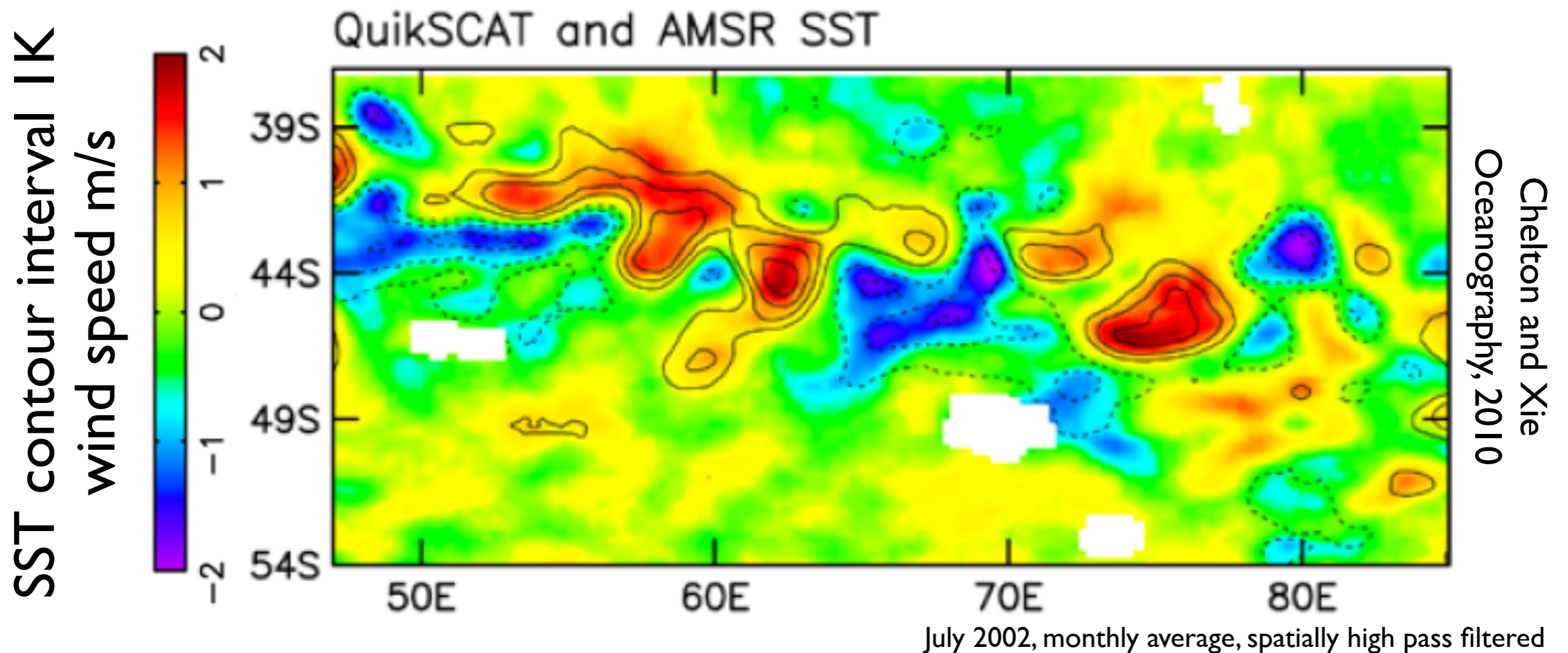


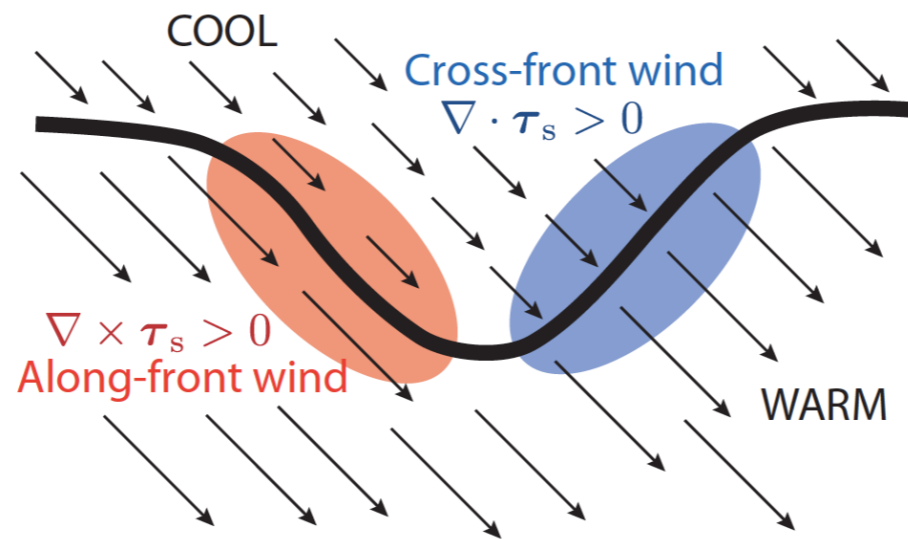
# The atmospheric to weak sea surface temperature fronts

Niklas Schneider and Bo Qiu

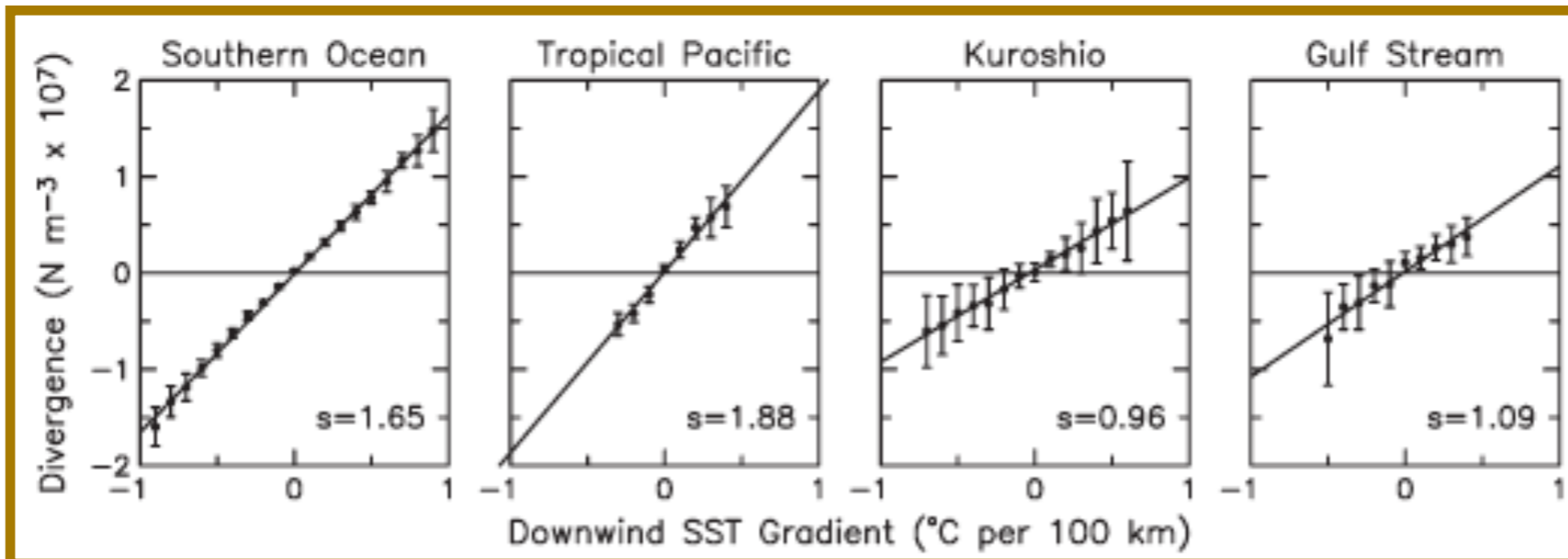
International Pacific Research Center and Department of Oceanography  
University of Hawaii at Manoa, Honolulu, HI, USA



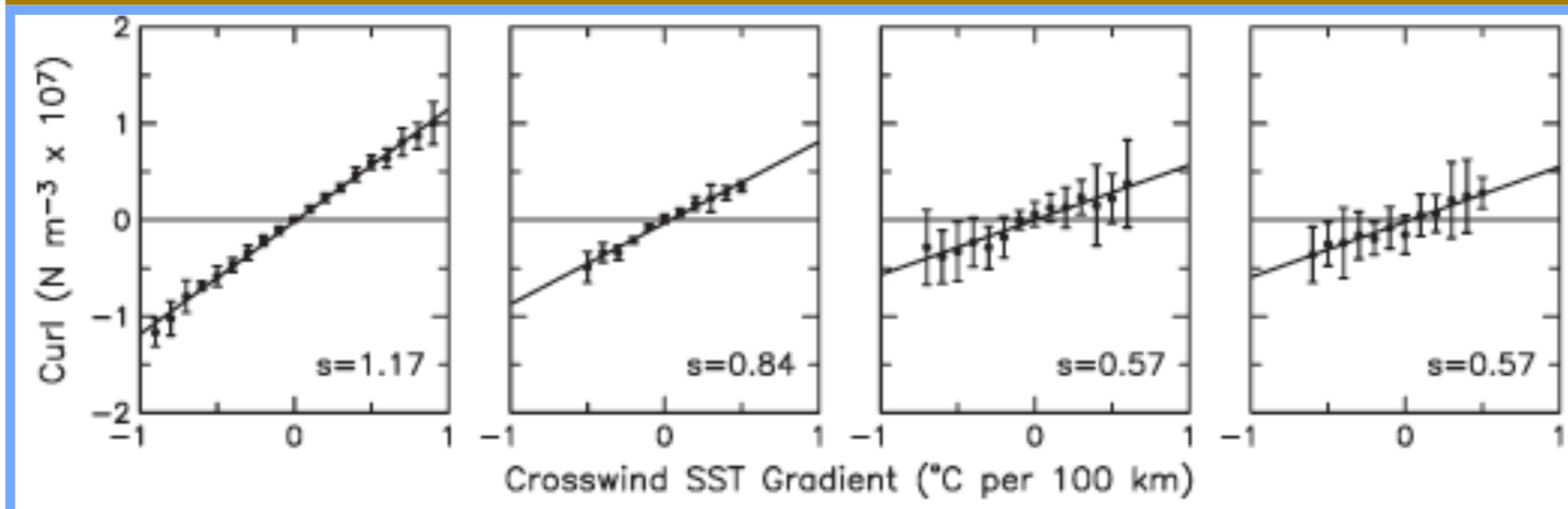
Imprint of ocean mesoscale on extratropical atmosphere  
warm SST associated with high wind speed



## Coupling coefficients



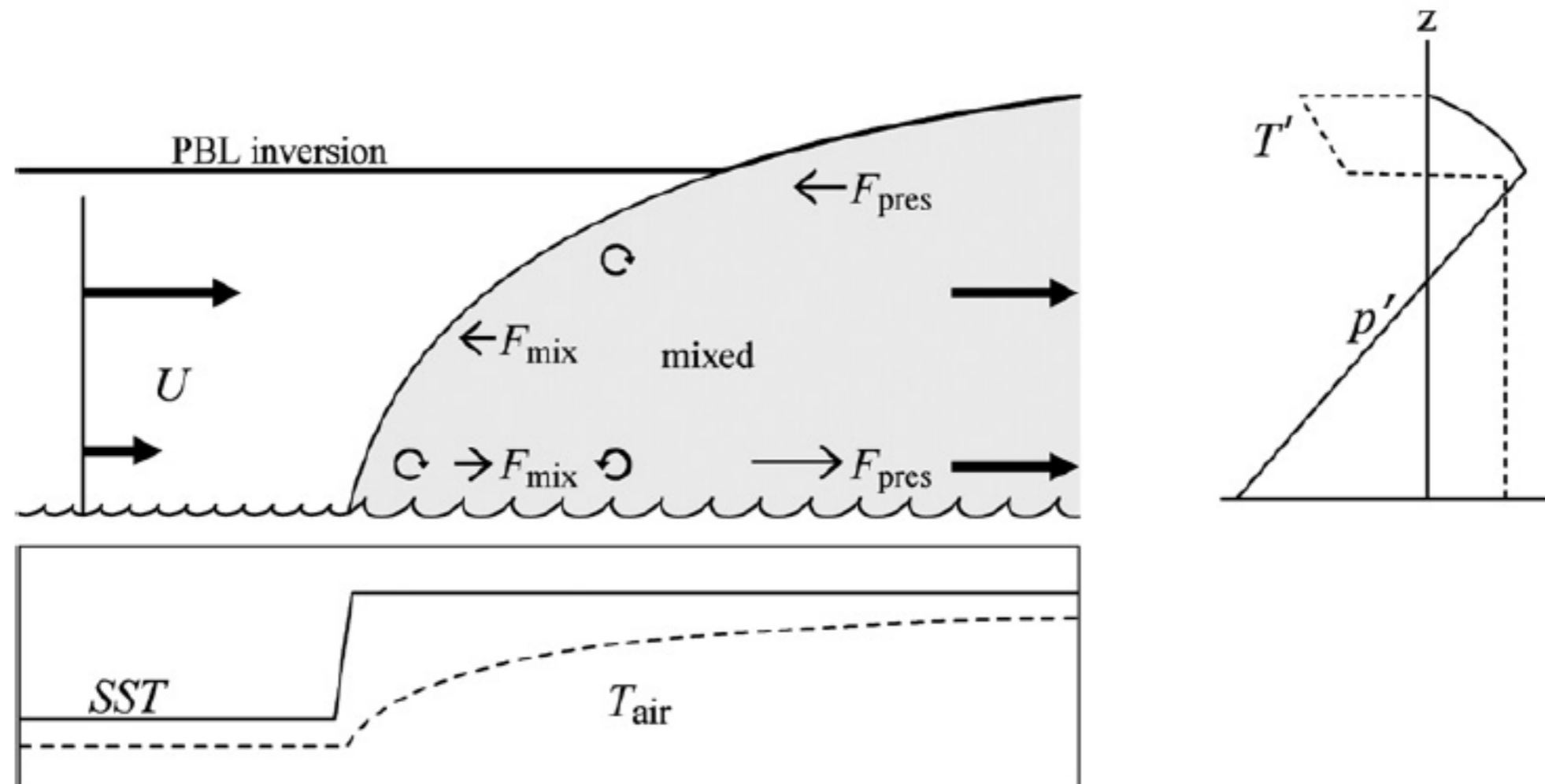
wind stress **divergence** function of **downwind** SST gradient



wind stress **curl** function of **crosswind** SST gradient

**slope s consistently larger for divergence than for curl**

# Impact of SST fronts on lower troposphere



Small et al. 2008

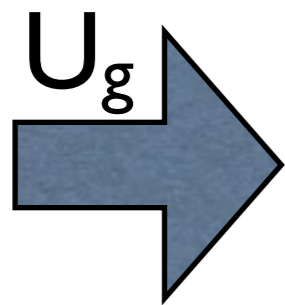
## Downstream of a warm SST front:

1. a **“Vertical Mixing”** increase in mixing entrains higher momentum from aloft (Wallace et al. 1989, Hayes et al. 1989, Samelson et al. 2006)
2. **“Pressure Effect”** imprint of SST gradient on boundary layer virtual temperature and baroclinic pressure gradients (Lindzen and Nigam 1987)
3. **“Spin-down”** Ekman transport convergences adjusts inversion height gradients to diminish surface Ekman pumping (Feliks et al. 2004)

# Model for air-sea interaction at SST fronts

- Reduced gravity model capped by sharp inversion
- Forced by barotropic tropospheric pressure gradient
- Background Ekman spiral
- Linear response to weak SST front

$h^{(0)}$  — inversion,  $\Delta\Theta$ , no flux



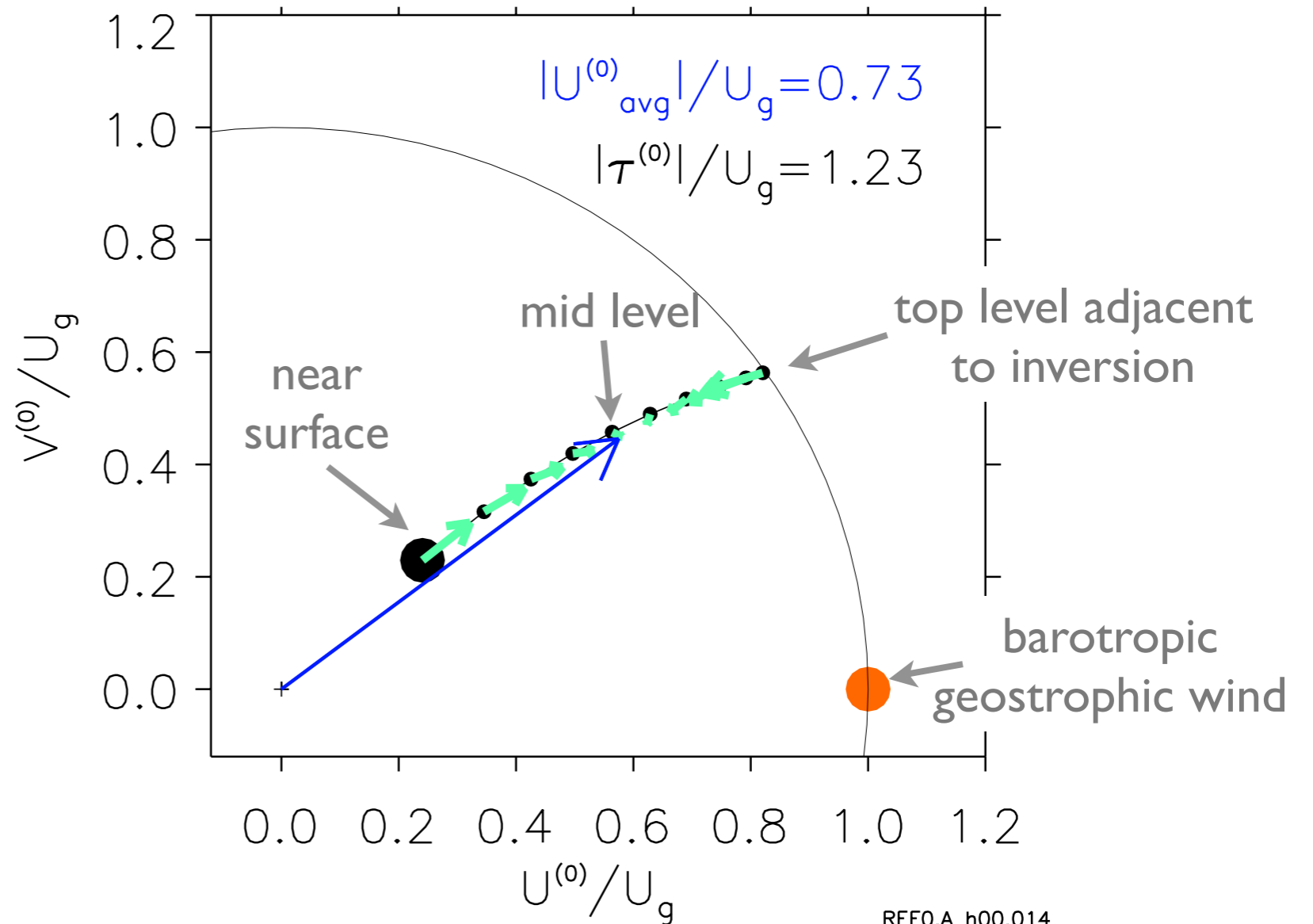
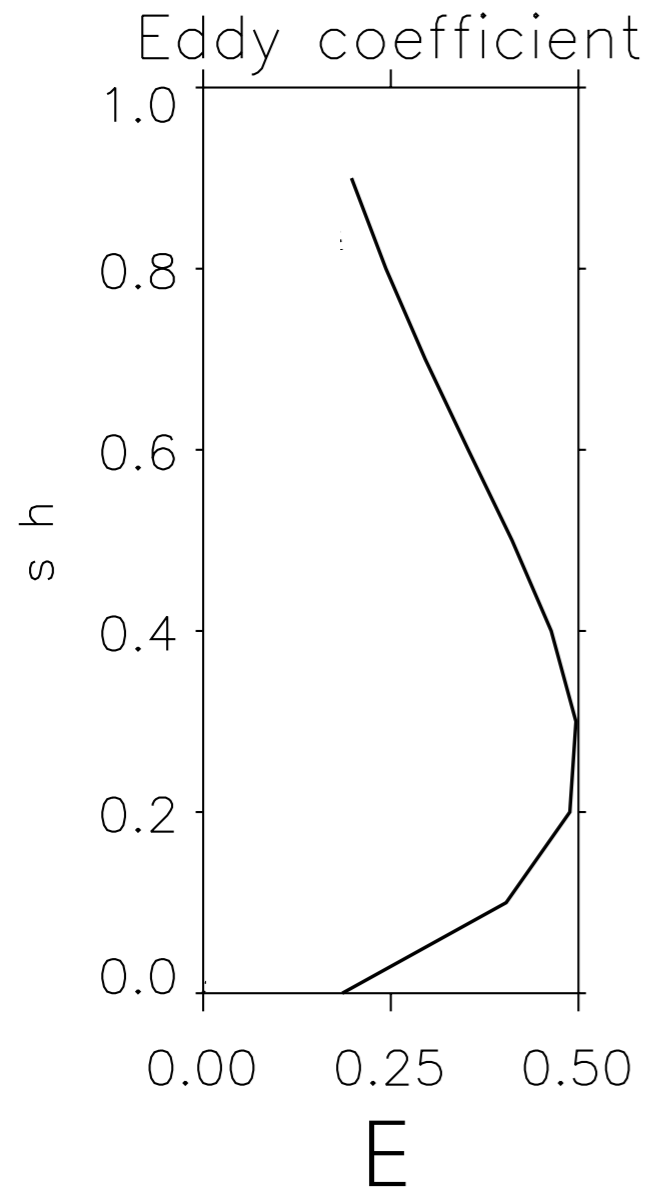
$u^{(0)}, v^{(0)}$  Ekman spiral  
 $\Theta^{(0)}$  constant

---

no ocean current, constant SST

# Background Ekman spiral

$$\hat{e}_3 \times (\vec{u}^{(0)} - \vec{U}_g) = \partial_s E^{(0)} \partial_s \vec{u}^{(0)}$$



# Air-sea interaction at weak SST front

## 1<sup>st</sup> order (linear) response

$$\bar{\vec{u}}^{(0)} \cdot \nabla \Theta^{(1)} = \gamma (T^{(1)} - \Theta^{(1)}) + A_h \nabla^2 \Theta^{(1)}$$

$$\underbrace{\vec{u}^{(0)} \cdot \nabla \vec{u}^{(1)} + w^{*(1)} \partial_s \vec{u}^{(0)}}_{\text{advection}} + \underbrace{\hat{e}_3 \times \vec{u}^{(1)}}_{\text{Coriolis}} + \underbrace{\nabla h^{(1)}}_{\text{back pressure}} - \partial_s \underbrace{E^{(0)} \partial_s \vec{u}^{(1)}}_{\text{background mixing}} = \vec{F}$$

$$\vec{u}^{(0)} \cdot \nabla h^{(1)} + \nabla \cdot \vec{u}^{(1)} + \partial_s w^{*(1)} = 0$$

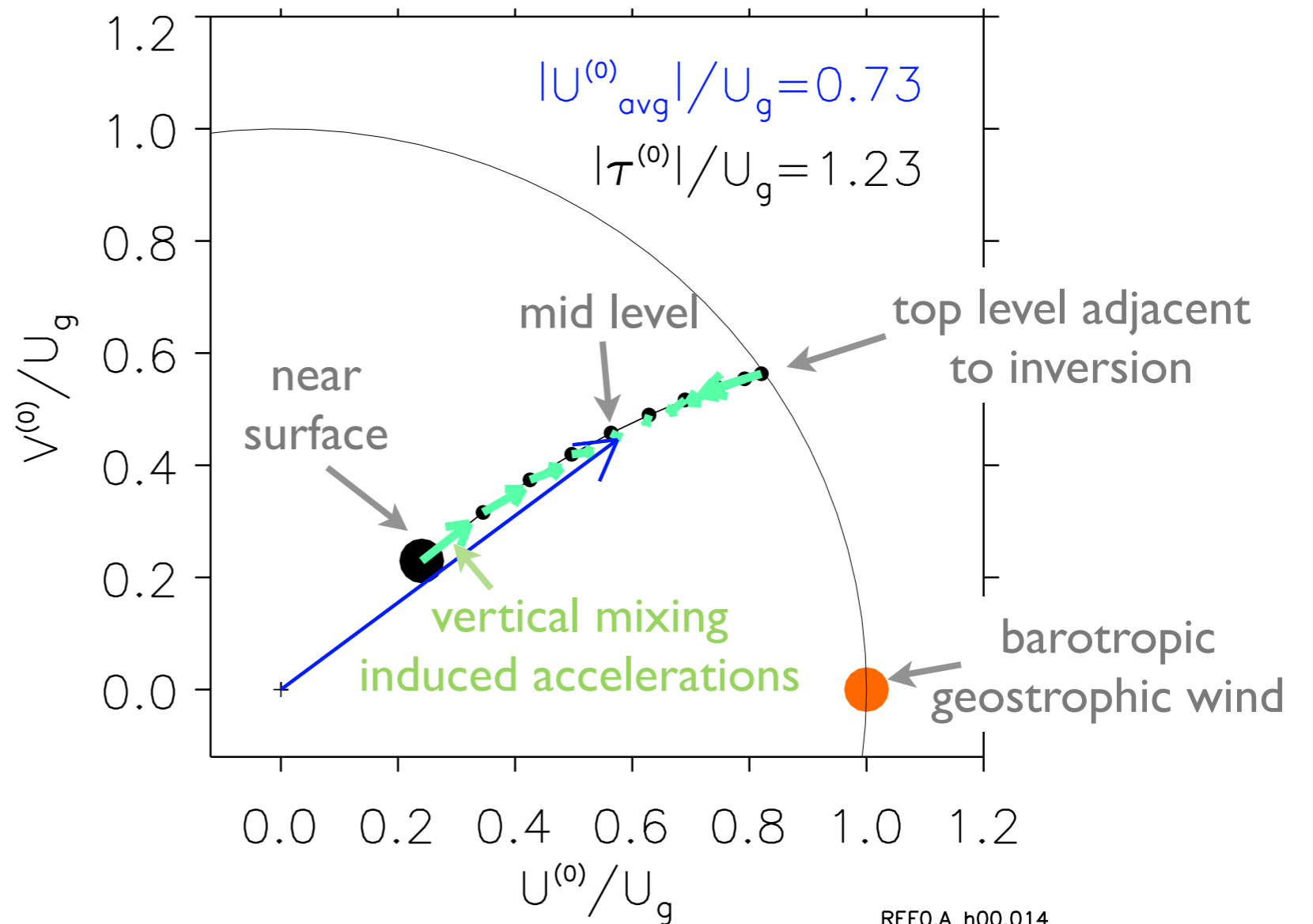
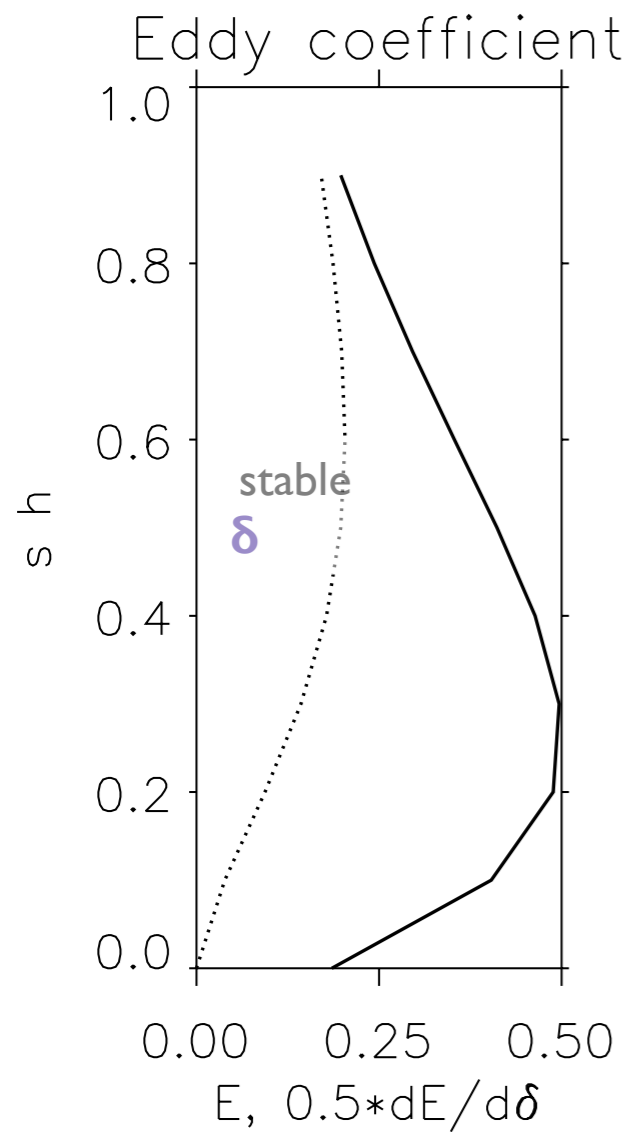
$$\vec{F} = \nabla \underbrace{\int_s^1 \Theta^{(1)} ds'}_{\text{pressure gradient mechanism}} + \partial_s \left( \underbrace{\delta^{(1)} \frac{\partial E}{\partial \delta} \Big|_{\delta^{(0)}} \partial_s \vec{u}^{(0)}}_{\text{vertical mixing mechanism}} \right)$$

$$\delta^{(1)} = T^{(1)} - \Theta^{(1)}$$

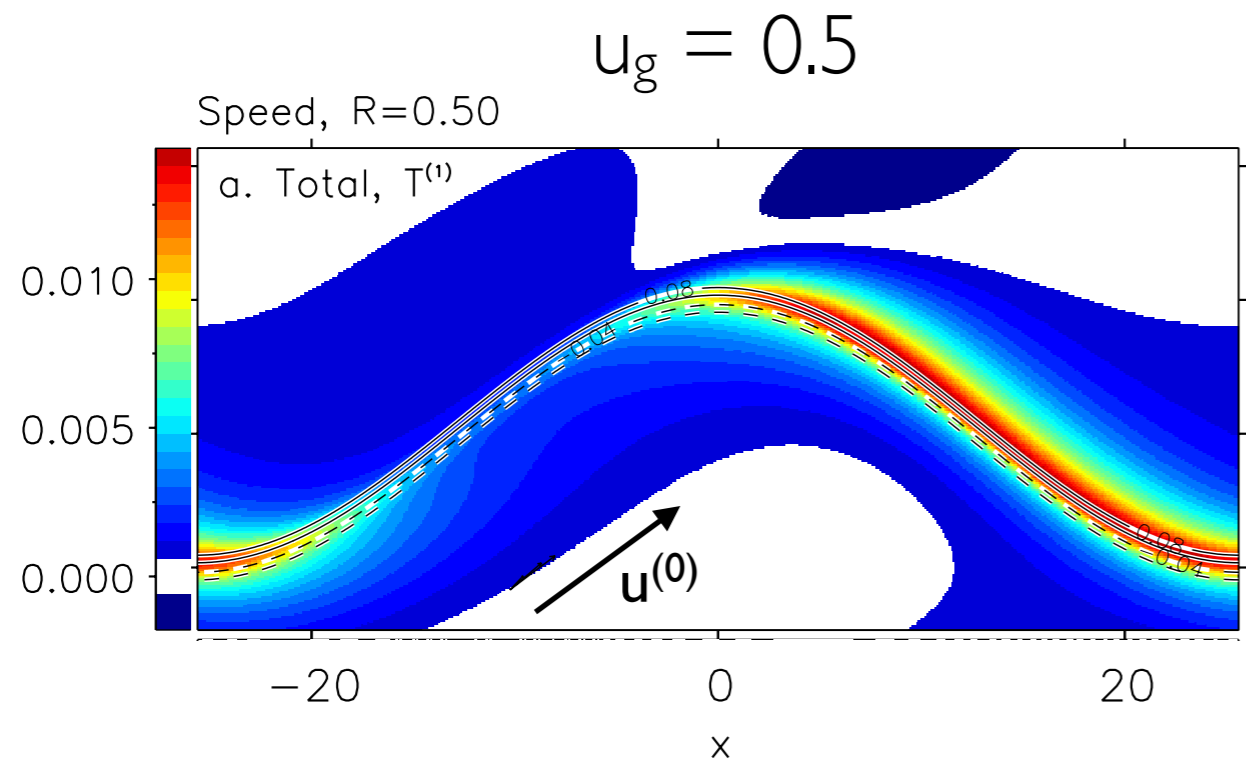
air-sea temperature  
differences modulates vertical  
eddy viscosity

nondimensionalized by Rossby Radius of deformation, boundary layer height, inversion strength etc.

# Vertical mixing effect



# Surface wind speed

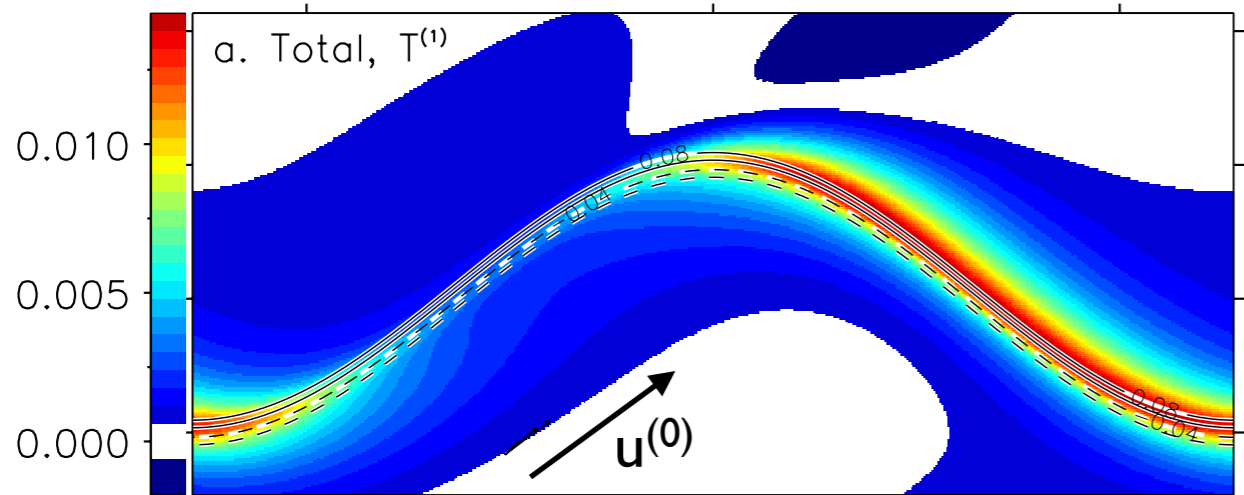




# Surface wind speed

$$u_g = 0.5$$

Speed,  $R=0.50$



b. Mixing,  $T^{(1)} - \Theta^{(1)}$

0.08  
0.04

This panel shows the mixing wind speed  $T^{(1)} - \Theta^{(1)}$  as a function of  $x$  and  $y$ . The color scale ranges from 0.04 (dark blue) to 0.08 (dark red). The plot shows a wave-like structure with a peak in mixing wind speed around  $x=0$  and  $y=0$ . A dashed line represents the surface wind speed  $u^{(0)}$ , and a solid line represents the mixing wind speed. An arrow labeled  $u^{(0)}$  points in the direction of the surface wind.

c. BCP,  $\Theta^{(1)}$

12  
6  
0  
-6

0.08  
0.04

This panel shows the BCP wind speed  $\Theta^{(1)}$  as a function of  $x$  and  $y$ . The color scale ranges from 0.04 (dark blue) to 0.08 (dark red). The plot shows a wave-like structure with a peak in BCP wind speed around  $x=0$  and  $y=0$ . A dashed line represents the surface wind speed  $u^{(0)}$ , and a solid line represents the BCP wind speed. An arrow labeled  $u^{(0)}$  points in the direction of the surface wind.

vertical  
mixing

baroclinic  
pressure

-20

x

20

y

y

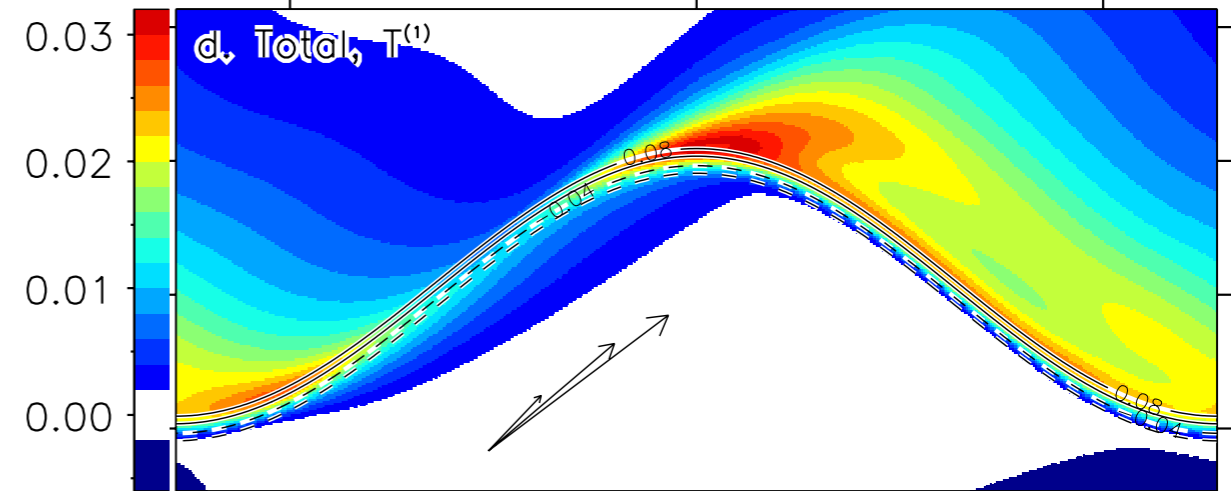
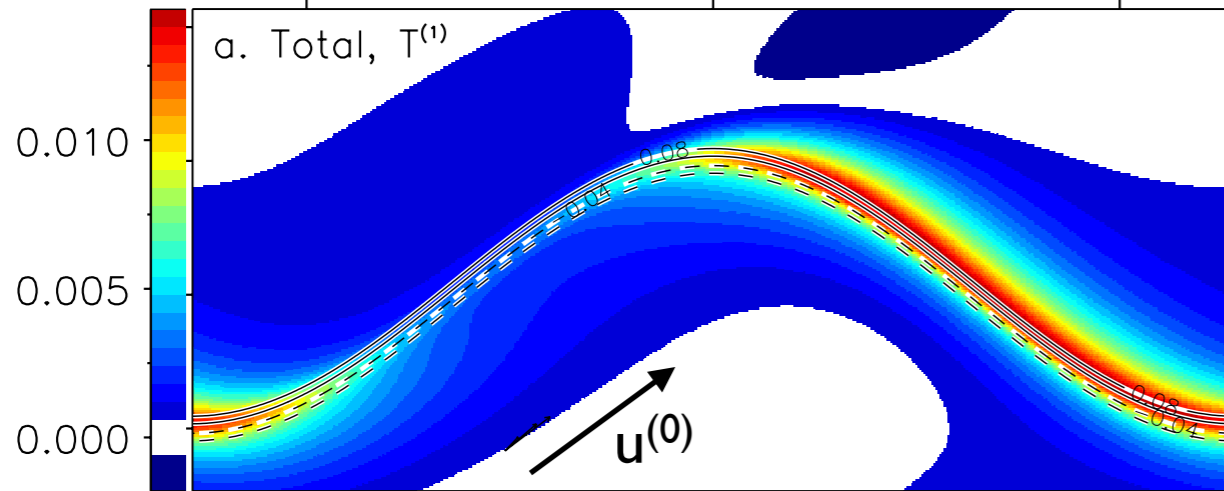
# Surface wind speed

$u_g = 0.5$

$u_g = 2$

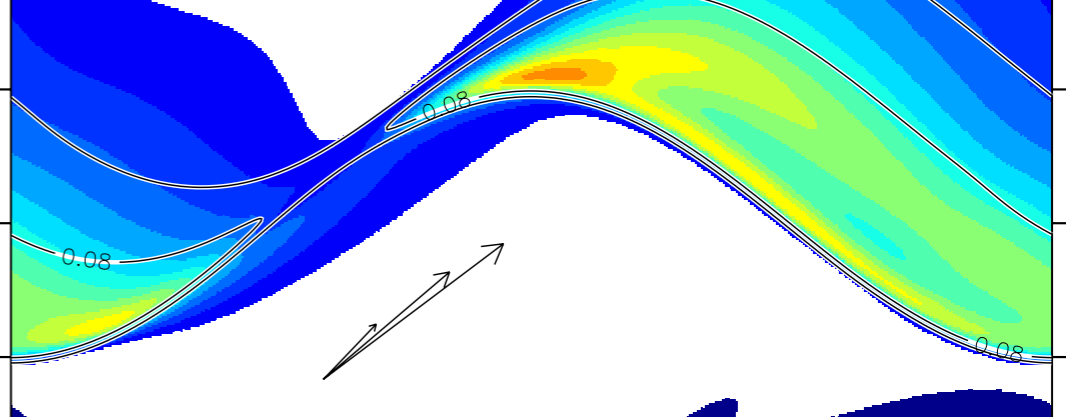
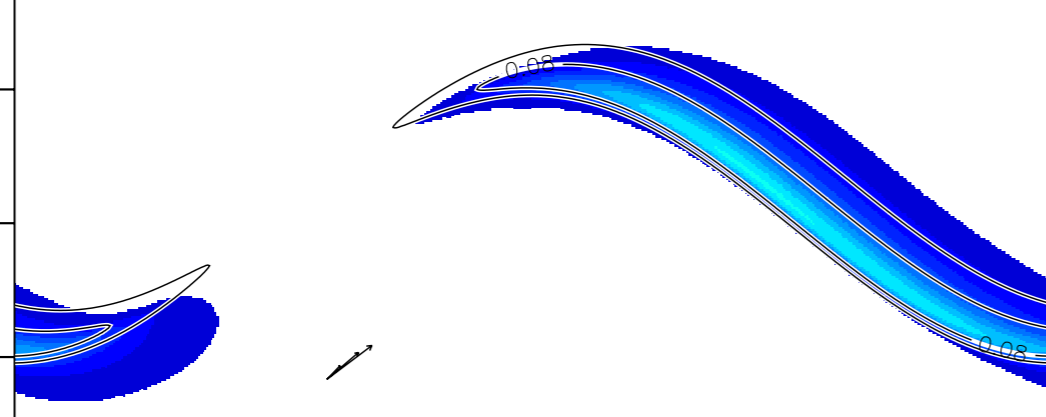
Speed,  $R=0.50$

Speed,  $R=2.00$



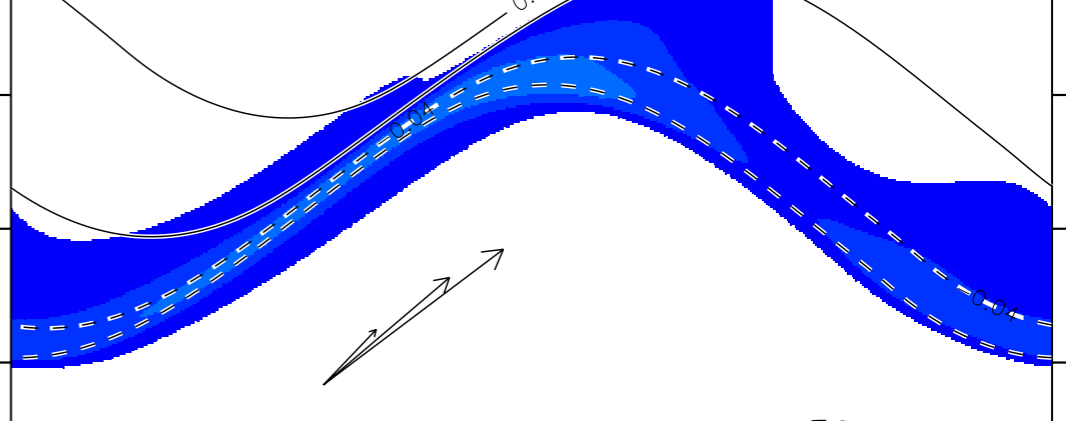
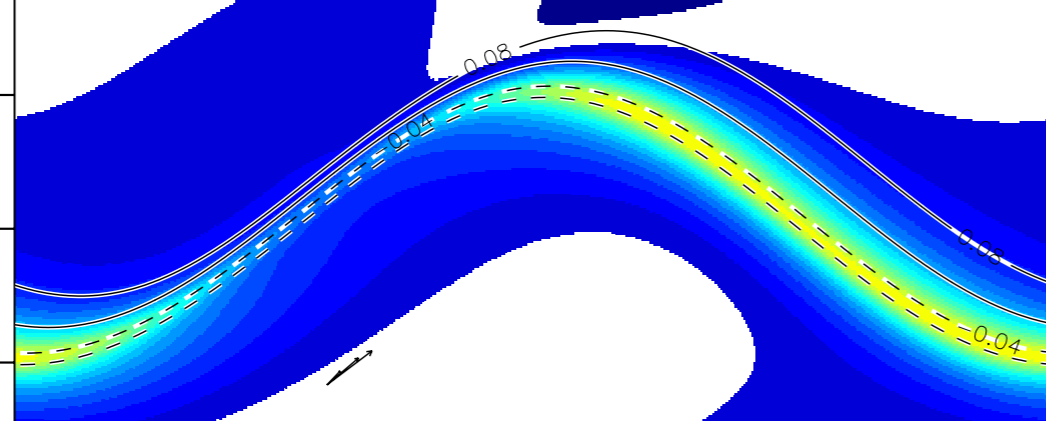
b. Mixing,  $T^{(1)} - \Theta^{(1)}$

e. Mixing,  $T^{(1)} - \Theta^{(1)}$



c. BCP,  $\Theta^{(1)}$

f. BCP,  $\Theta^{(1)}$



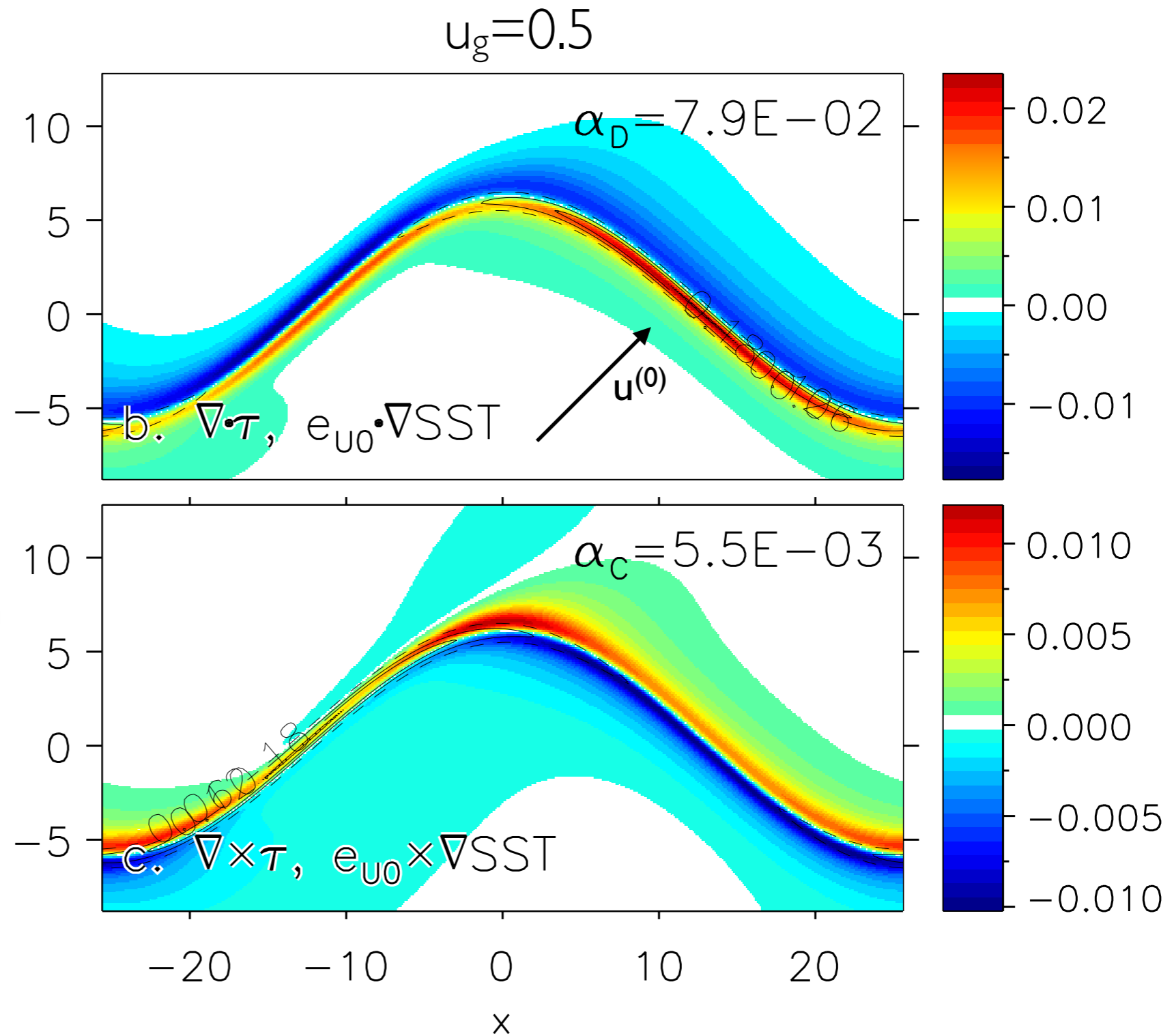
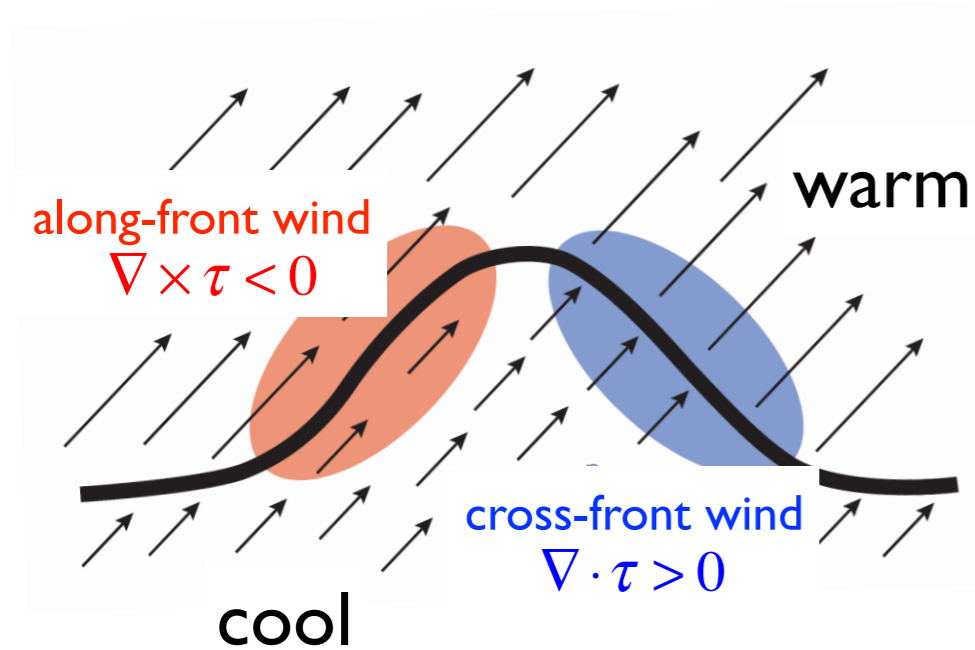
vertical mixing

baroclinic pressure

see Kilpatrick et al. 2014

REF1.ALH00.014

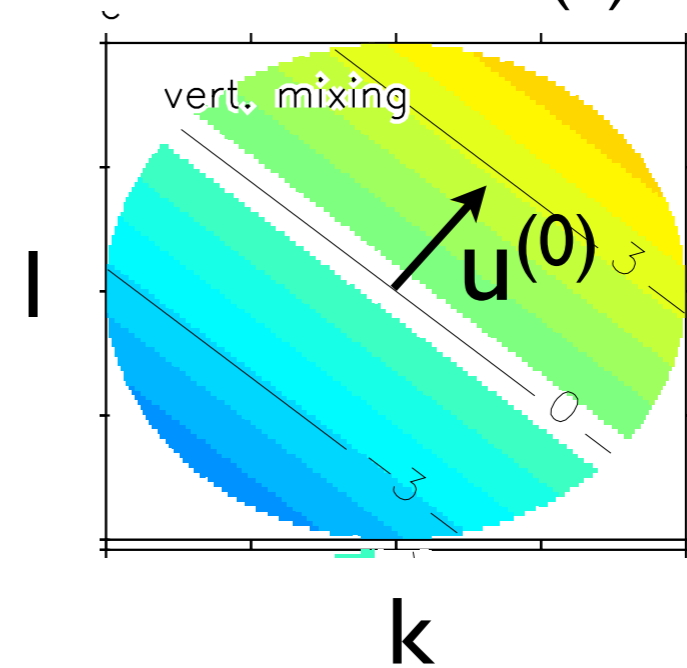
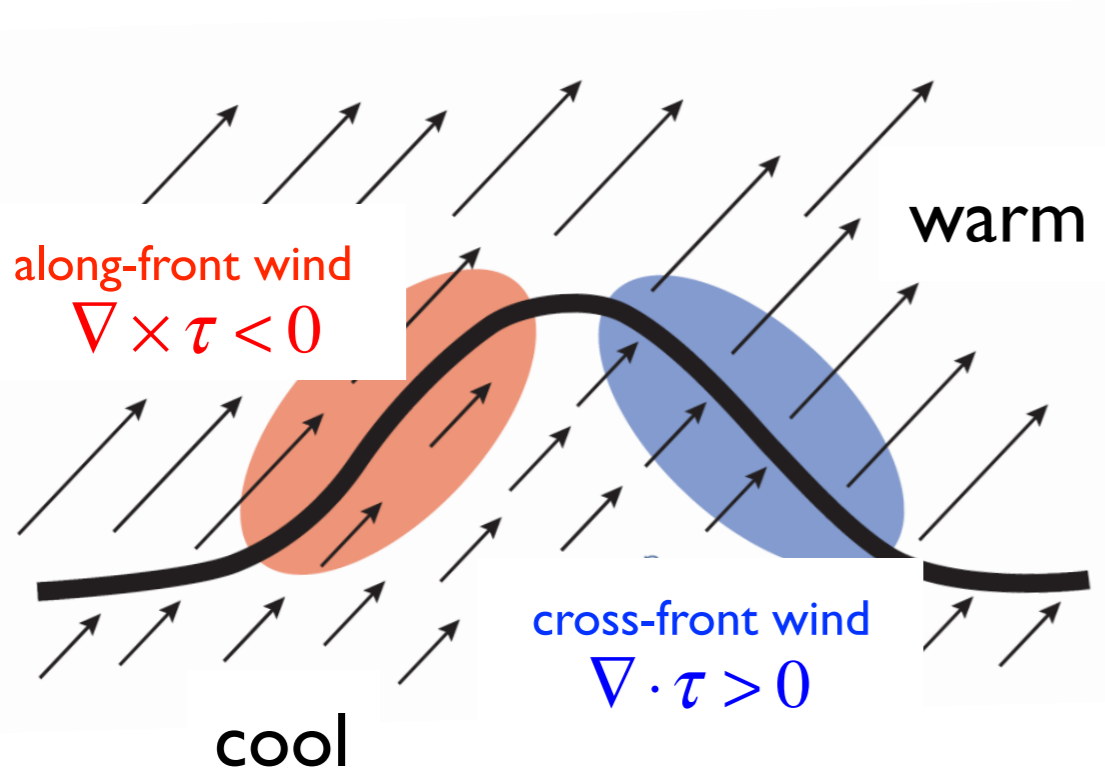
# Wind stress divergence and curl



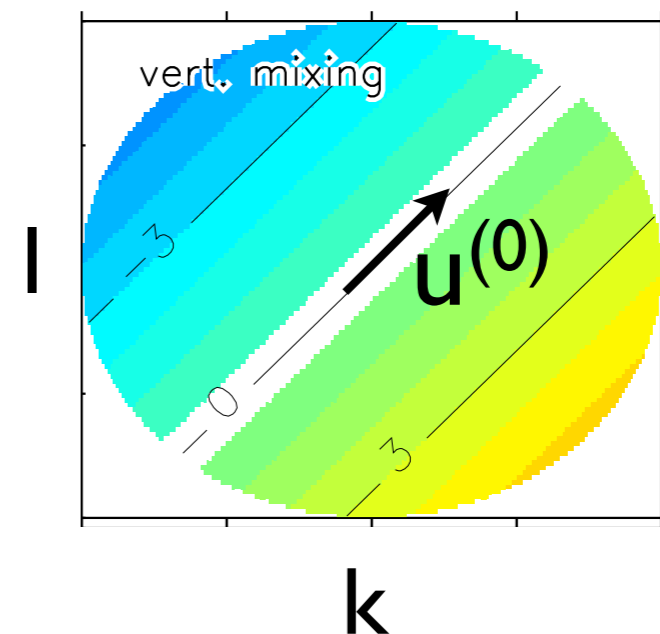
# Coupling Coefficients

in wavenumber space

$$\nabla \cdot \tau(k, l) = \alpha_D i k \cdot e_{u(0)} T^{(l)}(k, l)$$



$$\nabla \times \tau(k, l) = \alpha_C i k \times e_{u(0)} T^{(l)}(k, l)$$



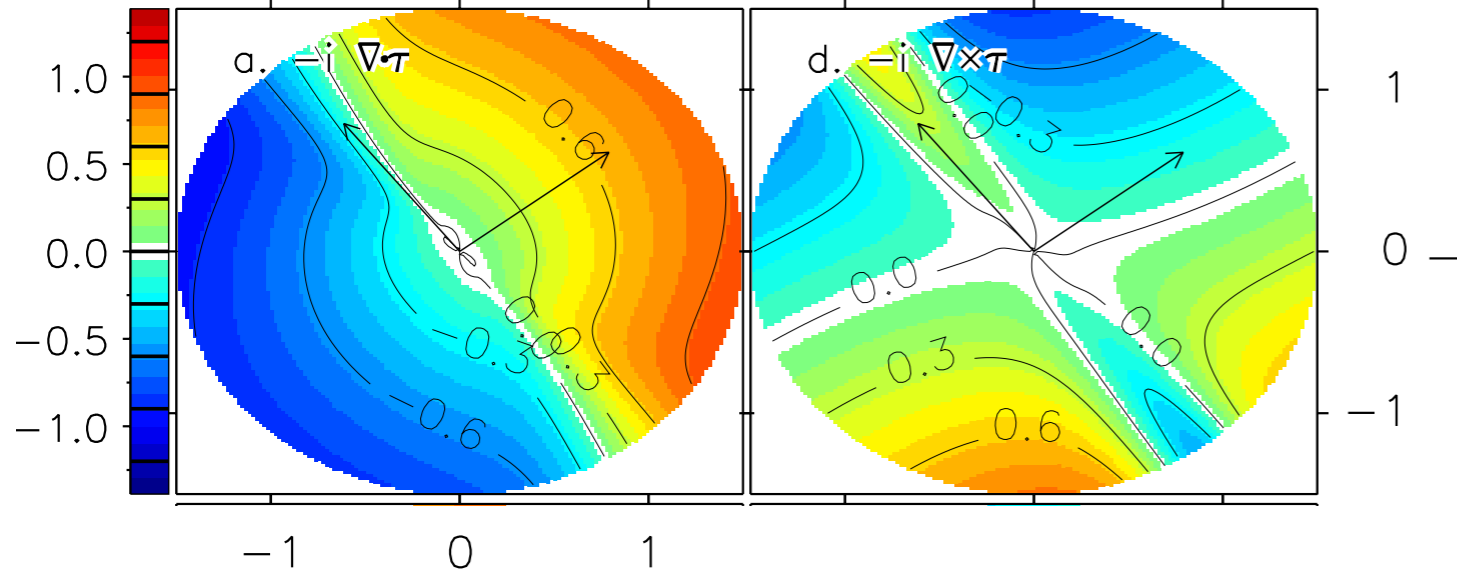
# Spin-down

## Wave-number space

divergence

curl

R=0.50



$$\nabla \cdot \tau(k, l) = \alpha_D i k \cdot \mathbf{e}_{u(0)} T^{(l)}(k, l)$$

$$\nabla \times \tau(k, l) = \alpha_C i k \times \mathbf{e}_{u(0)} T^{(l)}(k, l)$$

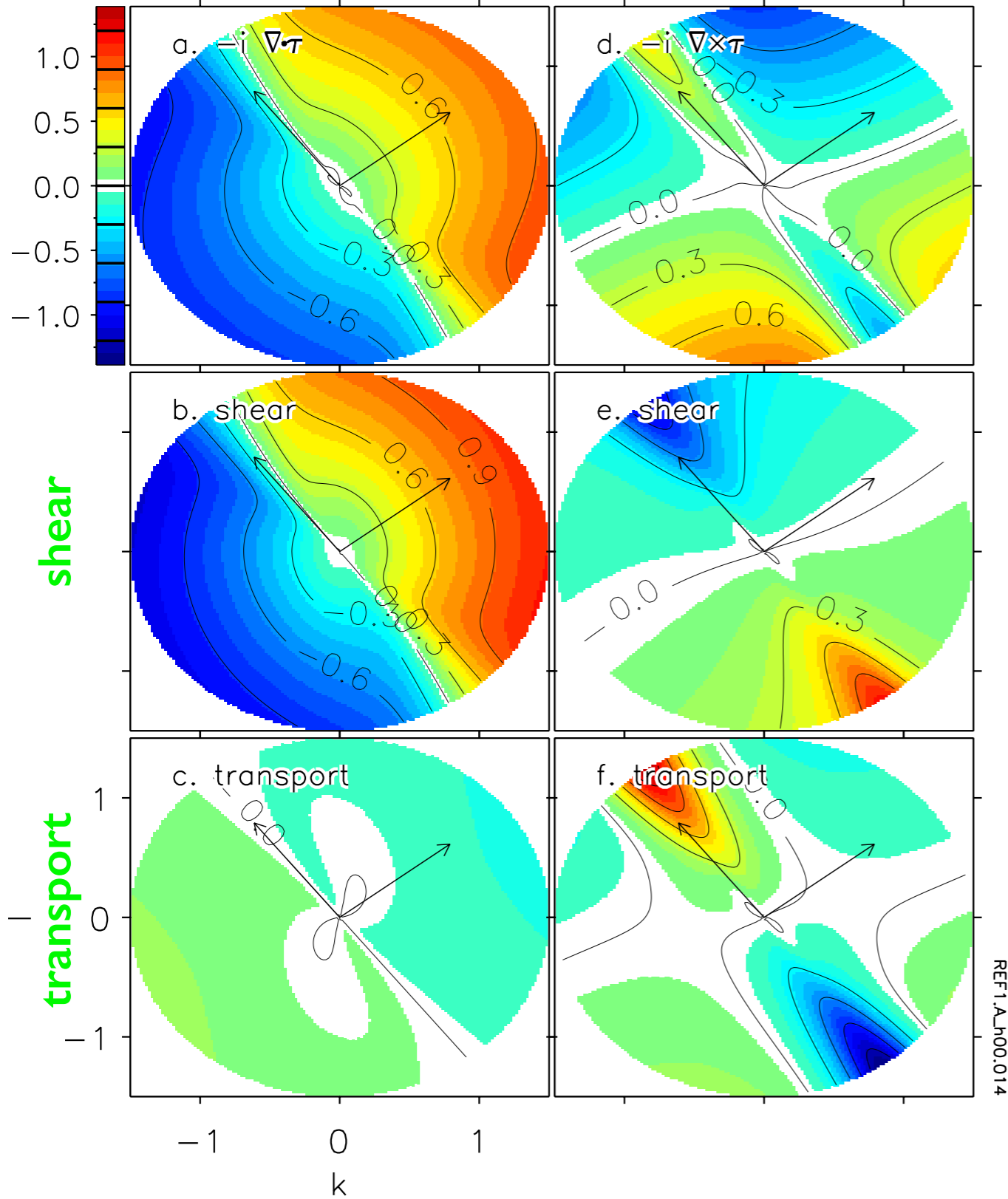
# Spin-down

## Wave-number space

divergence

curl

R=0.50



# Spin-down

Wave-number space

Physical space

divergence

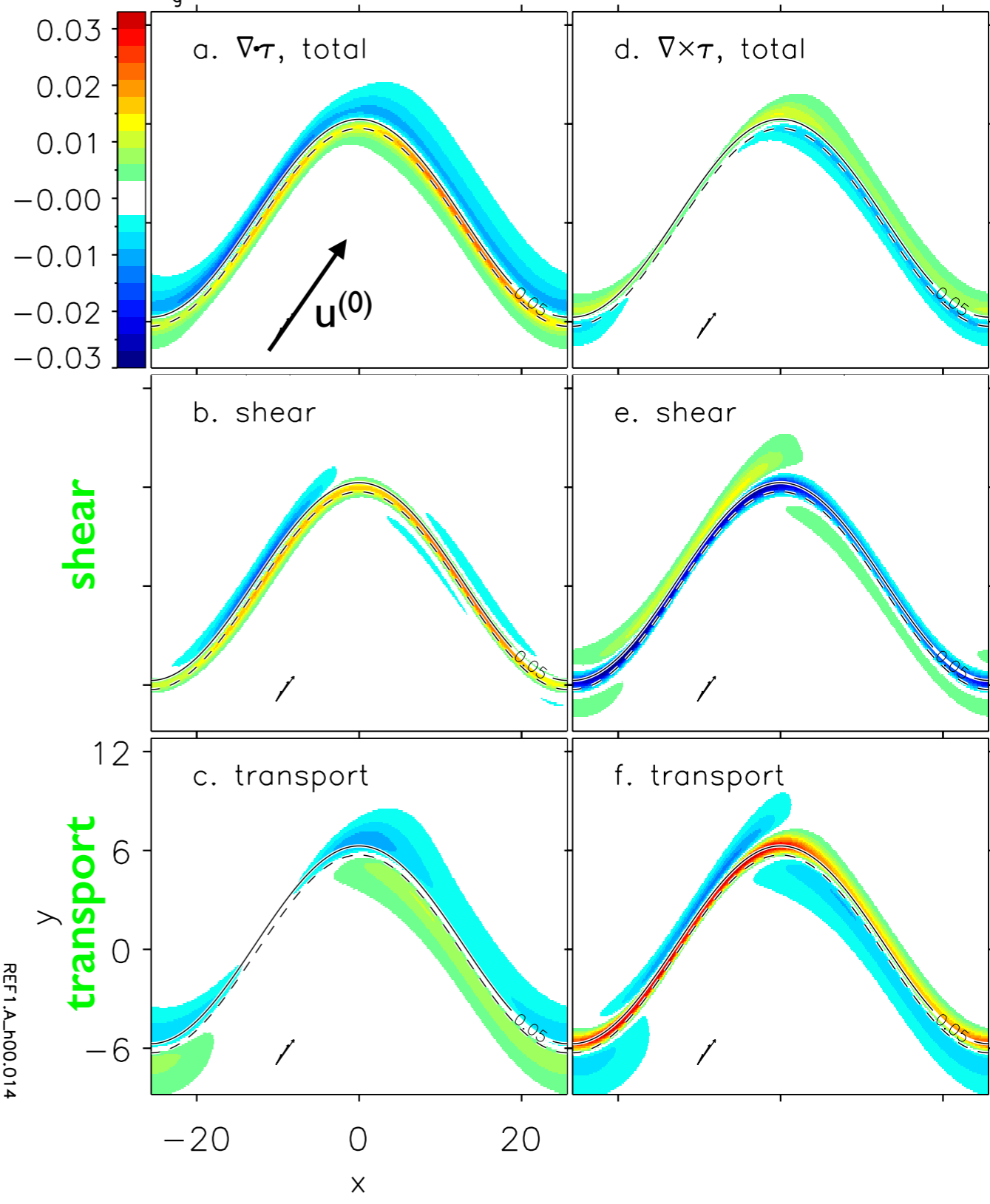
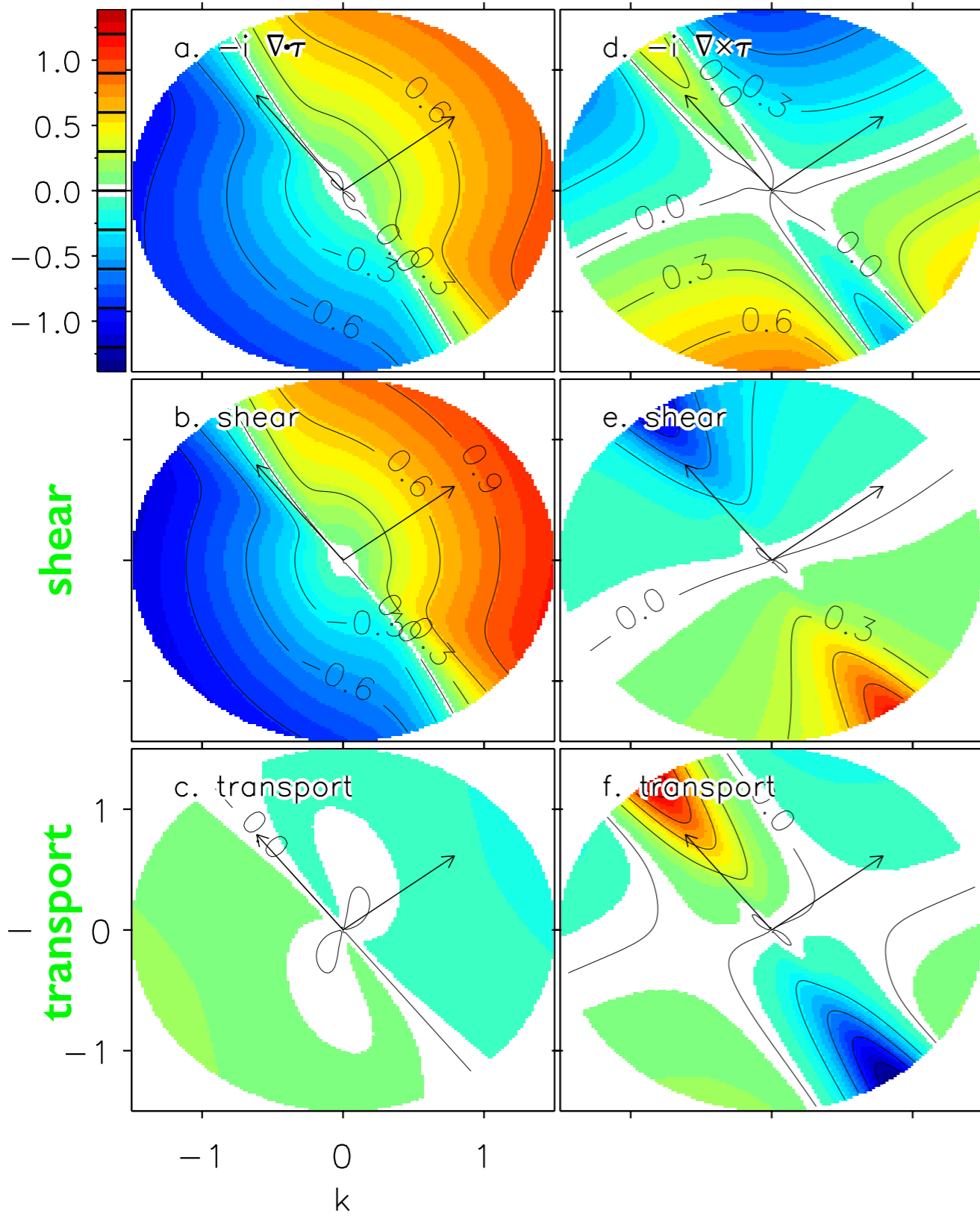
curl

divergence

curl

$R=0.50$

$U_g=0.50$



REF1.ALH00.014

# Conclusions

- Linear model captures observed modulation by SST front of
  - ✓ **lower/higher** wind speed over **cold/warm** water, wind stress curl and wind stress divergence
  - ✓ wind stress *divergence*/**curl** aligned with *down*/**cross**-wind gradient of SST
  - ✓ stronger coupling coefficients for wind stress *divergence* than for **curl**
- Dynamics governed by Rossby adjustment, gravity wave response for strong, geostrophic/spin-down for weak cross-frontal winds
- Distinct coupling coefficients of *divergence*/**curl** result from *gravitational sea breeze*/**spin-down**