

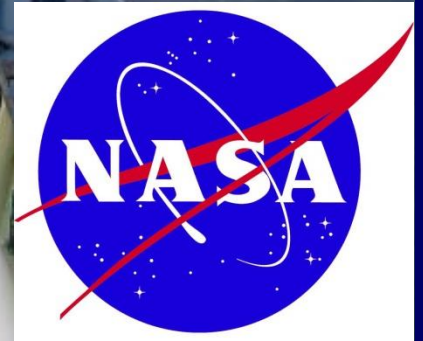
# The Response of Surface Winds, Wind Stress and Pressure to SST Variability on Monthly to Seasonal Time Scales

**Jim Edson**

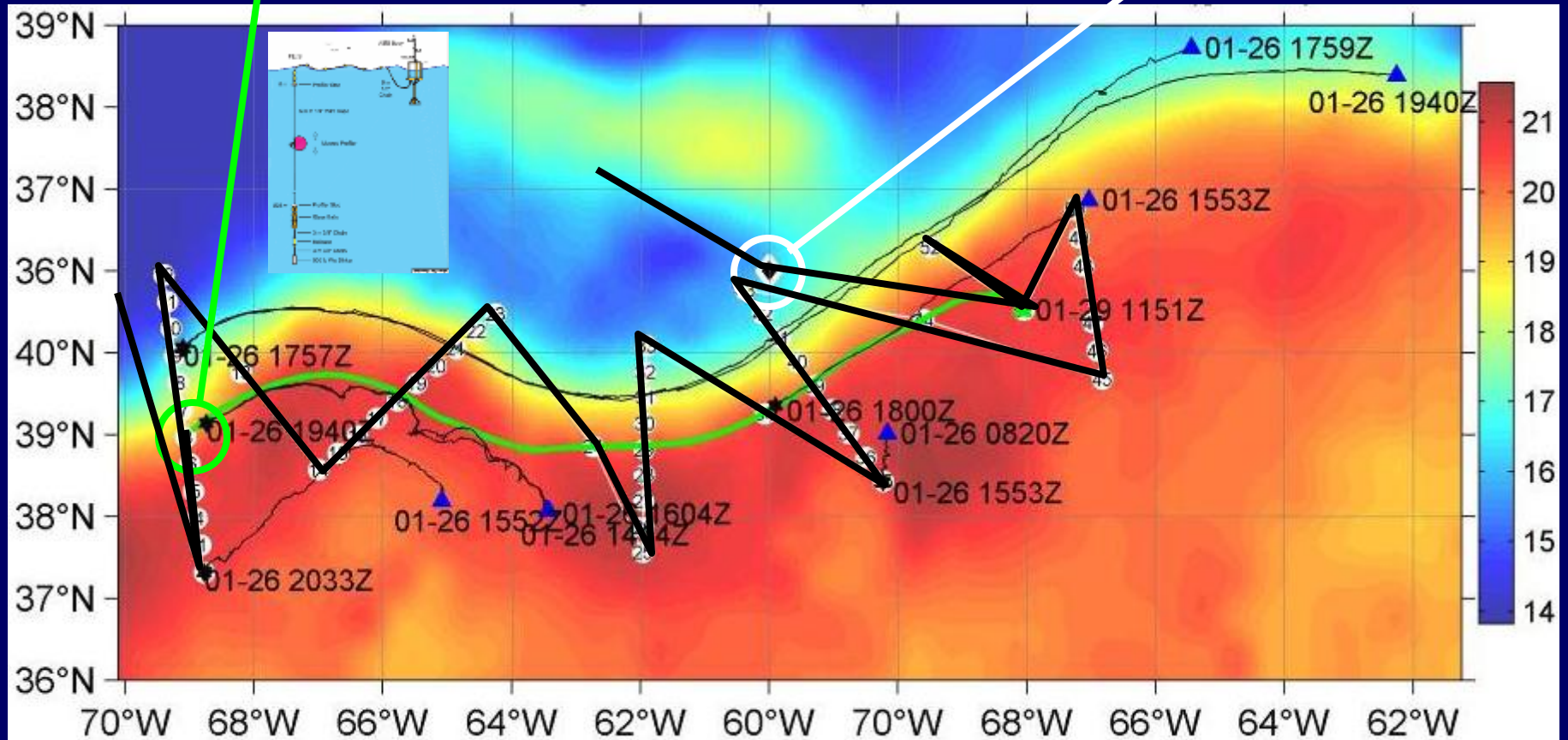
*University of Connecticut*

**Doug Vandemark & Amanda Plagge**

*University of New Hampshire*

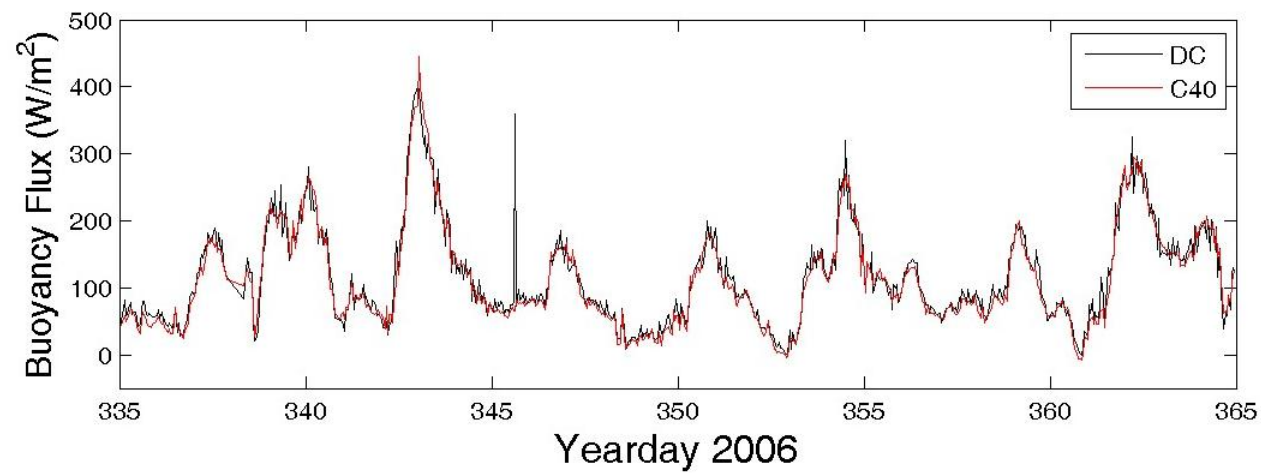
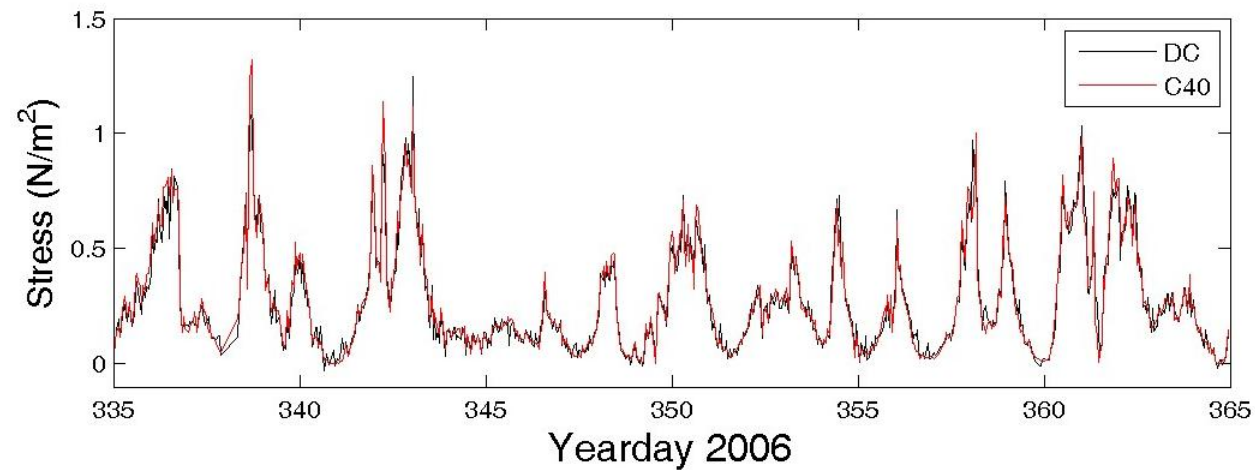


## A photograph showing a yellow offshore platform being lowered into the water from a ship's deck. A green arrow points to the platform's base in the water.





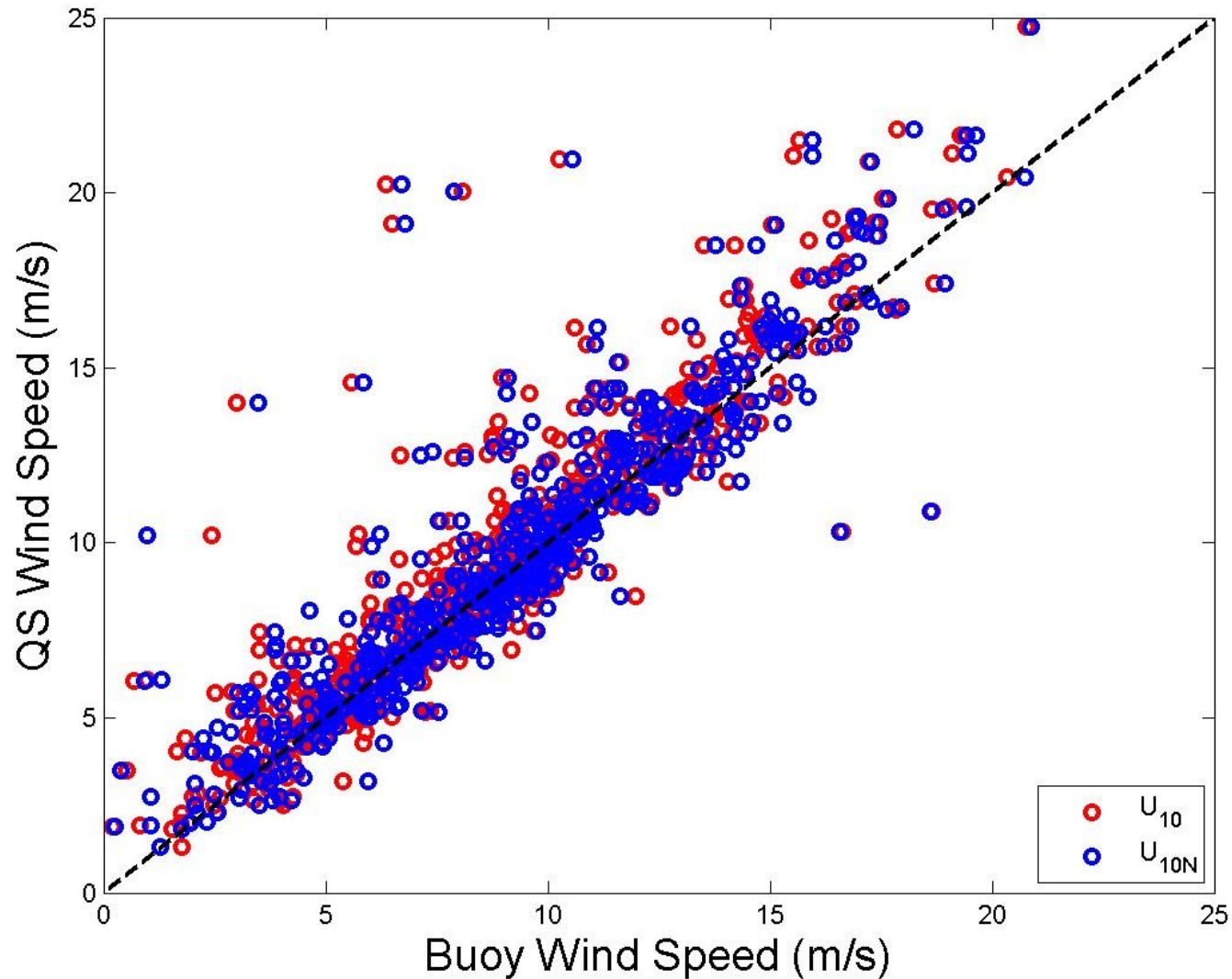
# Flux Time Series



# Stability/Baroclinic Effects Near SST Fronts

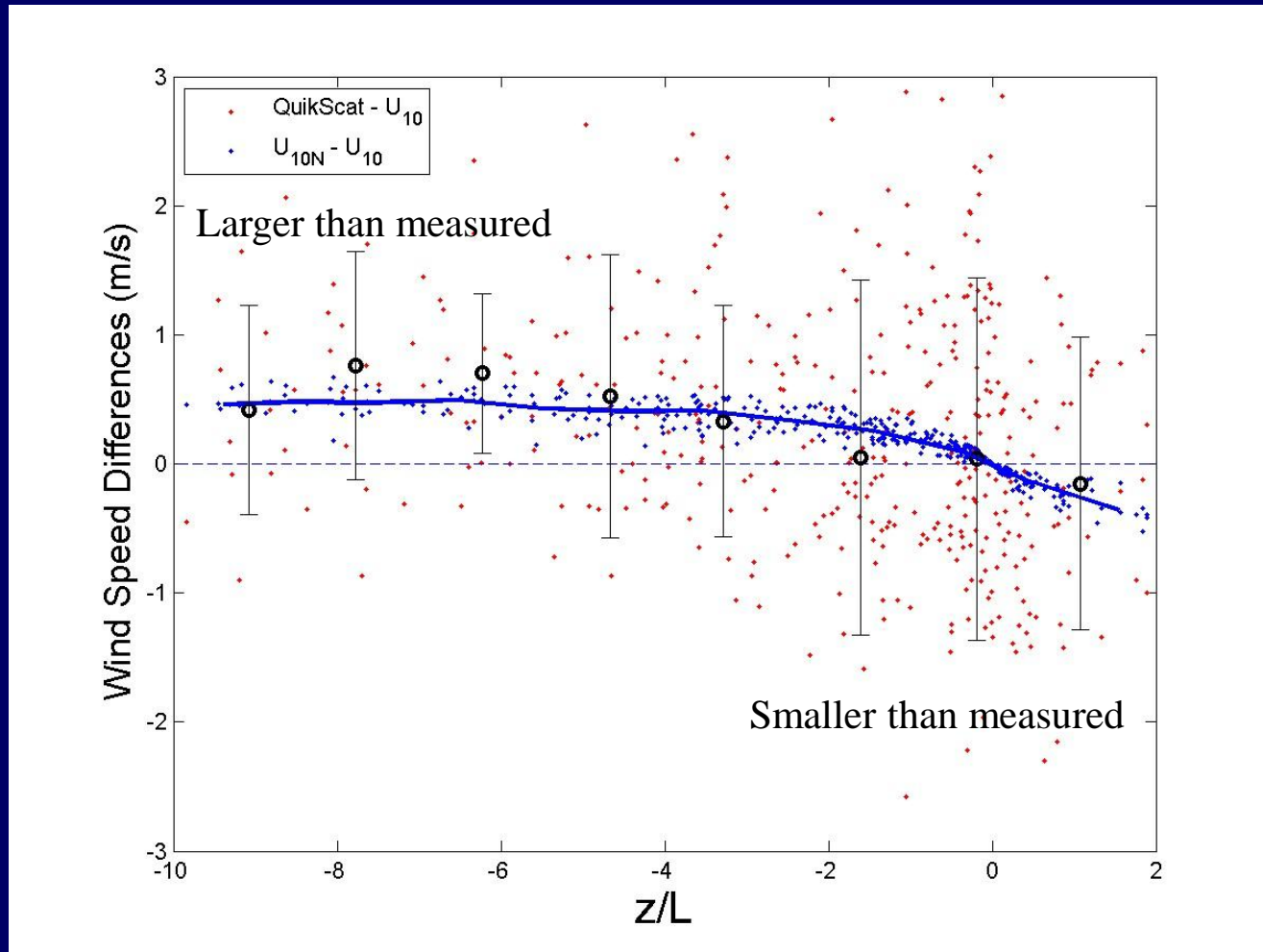
- Coupling Coefficient  $\Delta U = \alpha_v \Delta T_{sea}$  with higher winds over warmer and lower winds over cooler SST.
- Surface Layer Adjustment
  - QuikSCAT measures surface roughness/stress
  - Surface stress is proportional to neutral winds,  $U_N$
  - MO similarity predicts:
    - $U_N < U$  in unstable conditions
    - $U_N > U$  in stable conditions
- Boundary Layer Adjustment
  - Acceleration/deceleration of surface winds.
  - Enhancement of vertical mixing due to cool-air advection over warmer water that mixes down larger momentum from aloft.
  - Pressure perturbations driven by the adjustment of air temperature and humidity to the underlying SST.
  - Both!

# QuikSCAT vs. Buoy Wind Speeds



# QuikSCAT vs. Buoy Wind Speeds

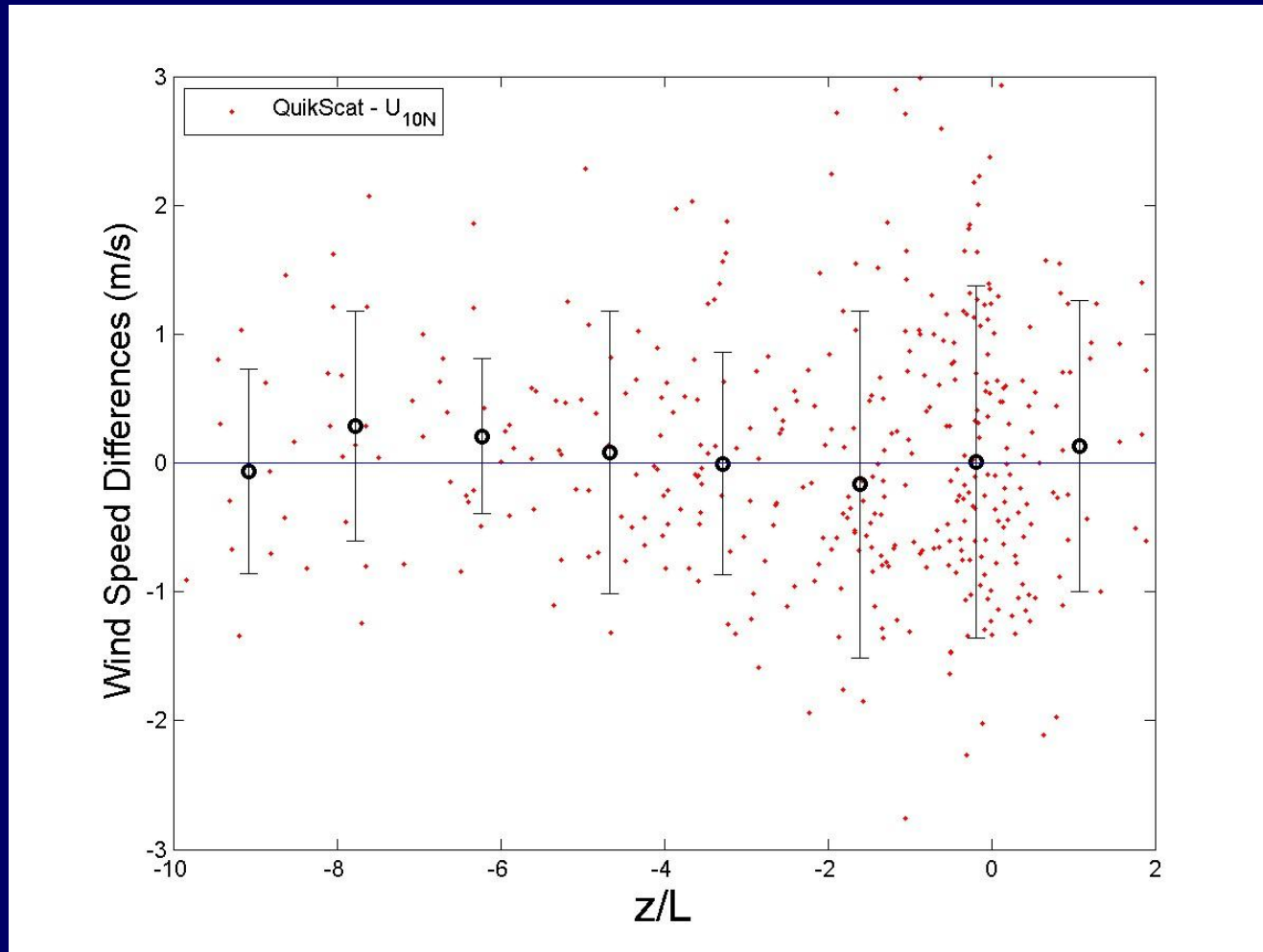
## “Surface Layer Adjustment”



$$U(z) = u_*/\kappa[\ln(z/z_0) - \psi_m(z/L)] \quad U_N(z) = u_*/\kappa[\ln(z/z_0)]$$

# QuikSCAT vs. Buoy Wind Speeds

## “Surface Layer Adjustment”



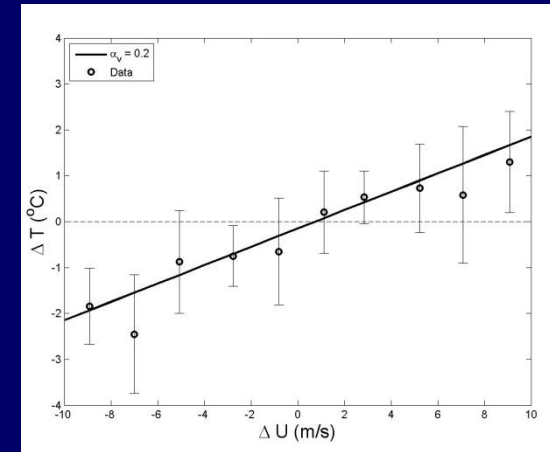
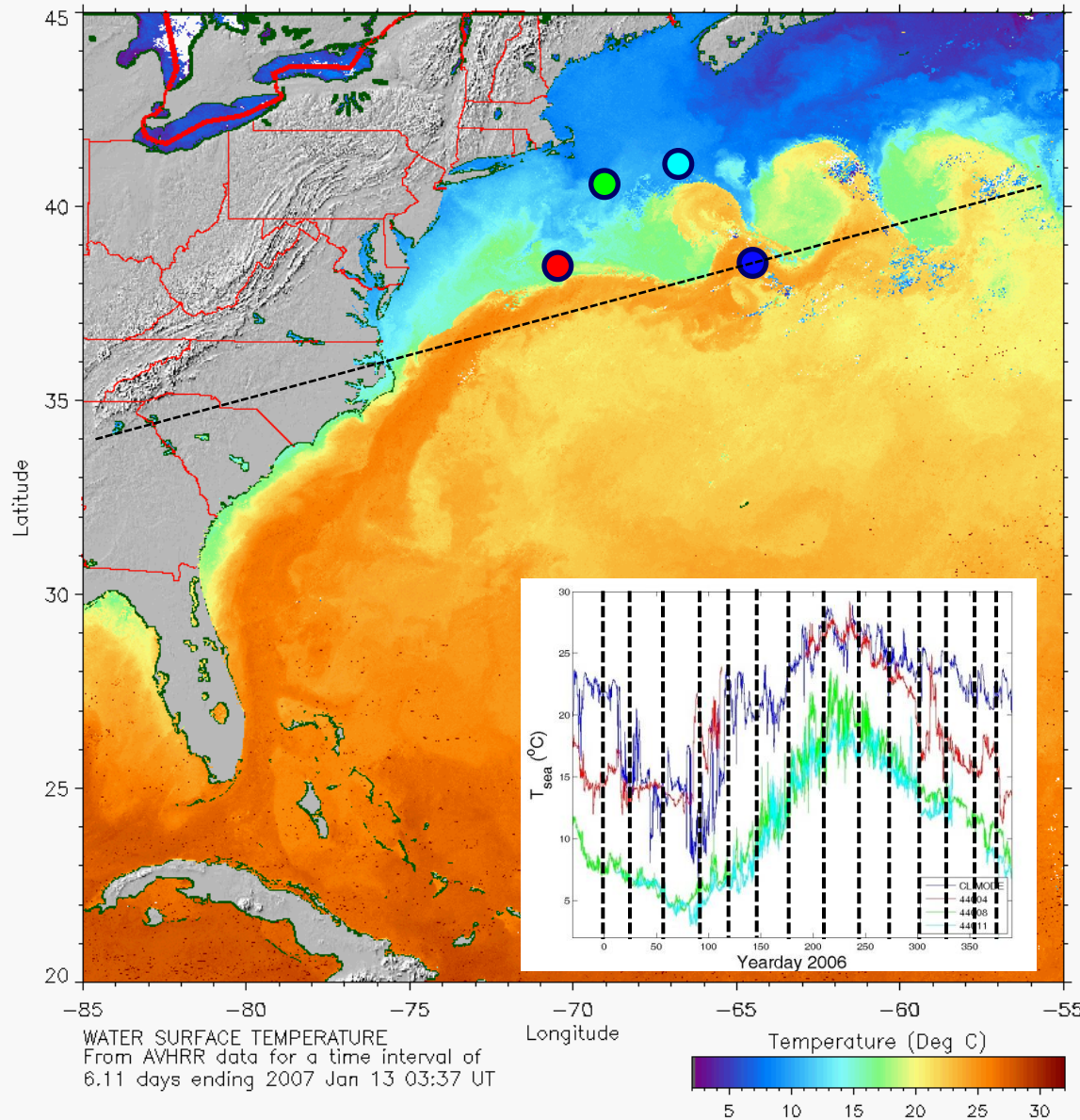
$$U(z) = u_*/\kappa[\ln(z/z_o) - \psi_m(z/L)] \quad U_N(z) = u_*/\kappa[\ln(z/z_o)]$$



# Buoy-Pair Approach

$$\Delta U = \alpha_v \Delta T_{sea}$$

O'Neill (2012)





# Single Buoy Approach

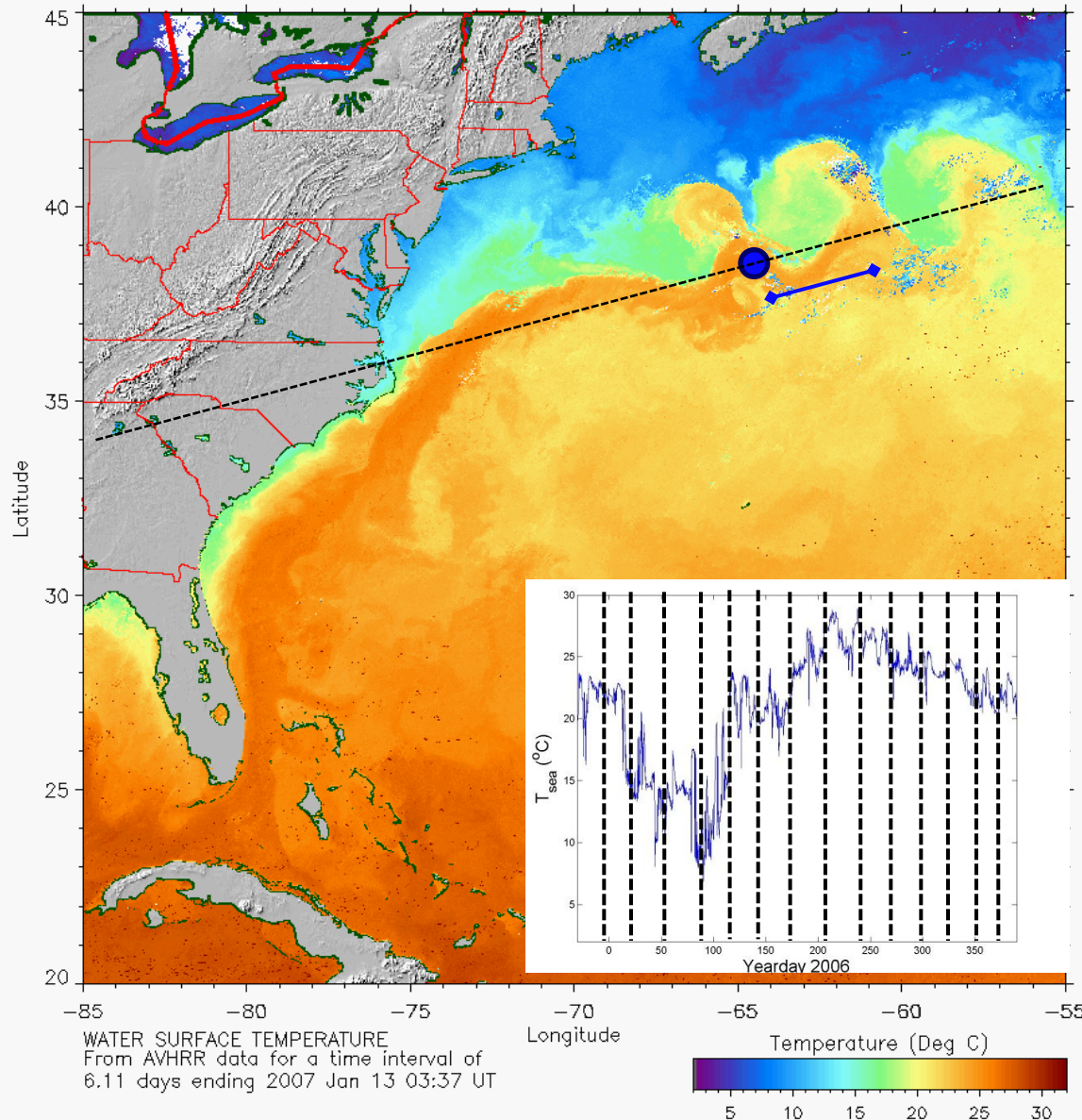
$$\Delta U = \alpha_v \Delta T_{sea}$$

Halliwell &  
Mooers (1983)

Gulf Stream Meanders

Wavelength 200-400 km

Phase speed 5-10 cm/s



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Image courtesy of JHU/APL

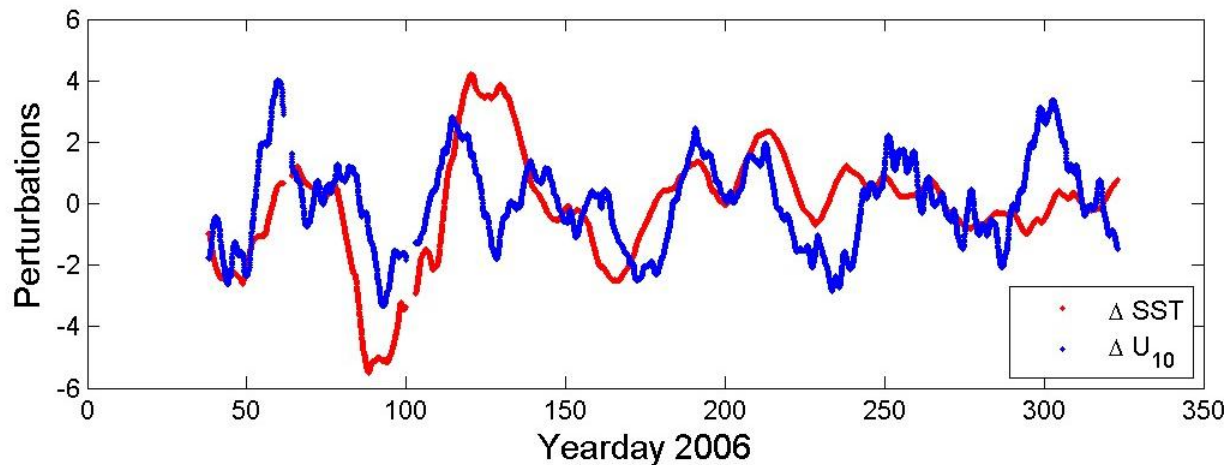
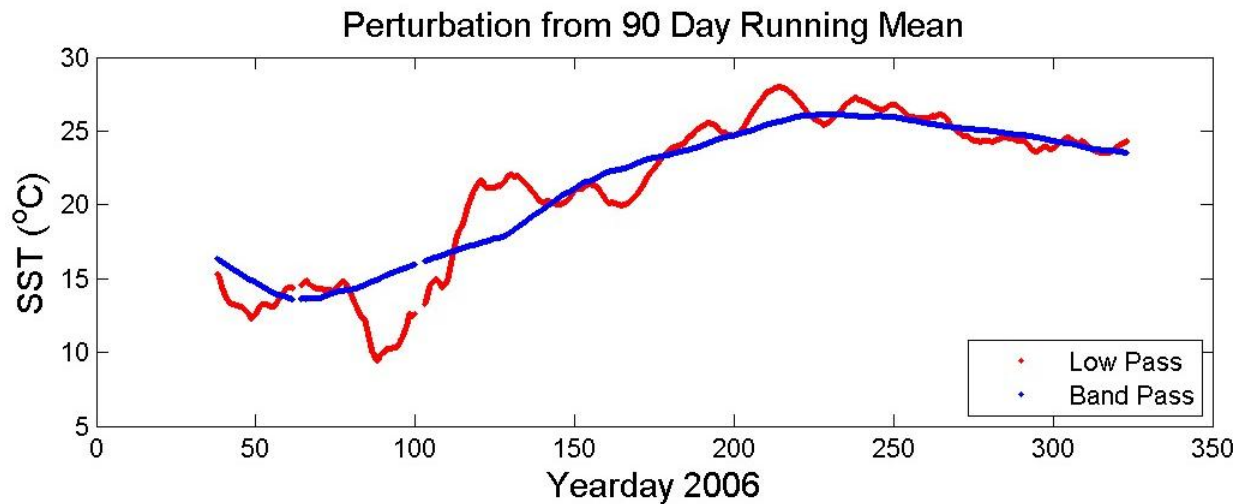
# Single Buoy Approach

Halliwel &  
Mooers (1983)

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# Single Buoy Approach

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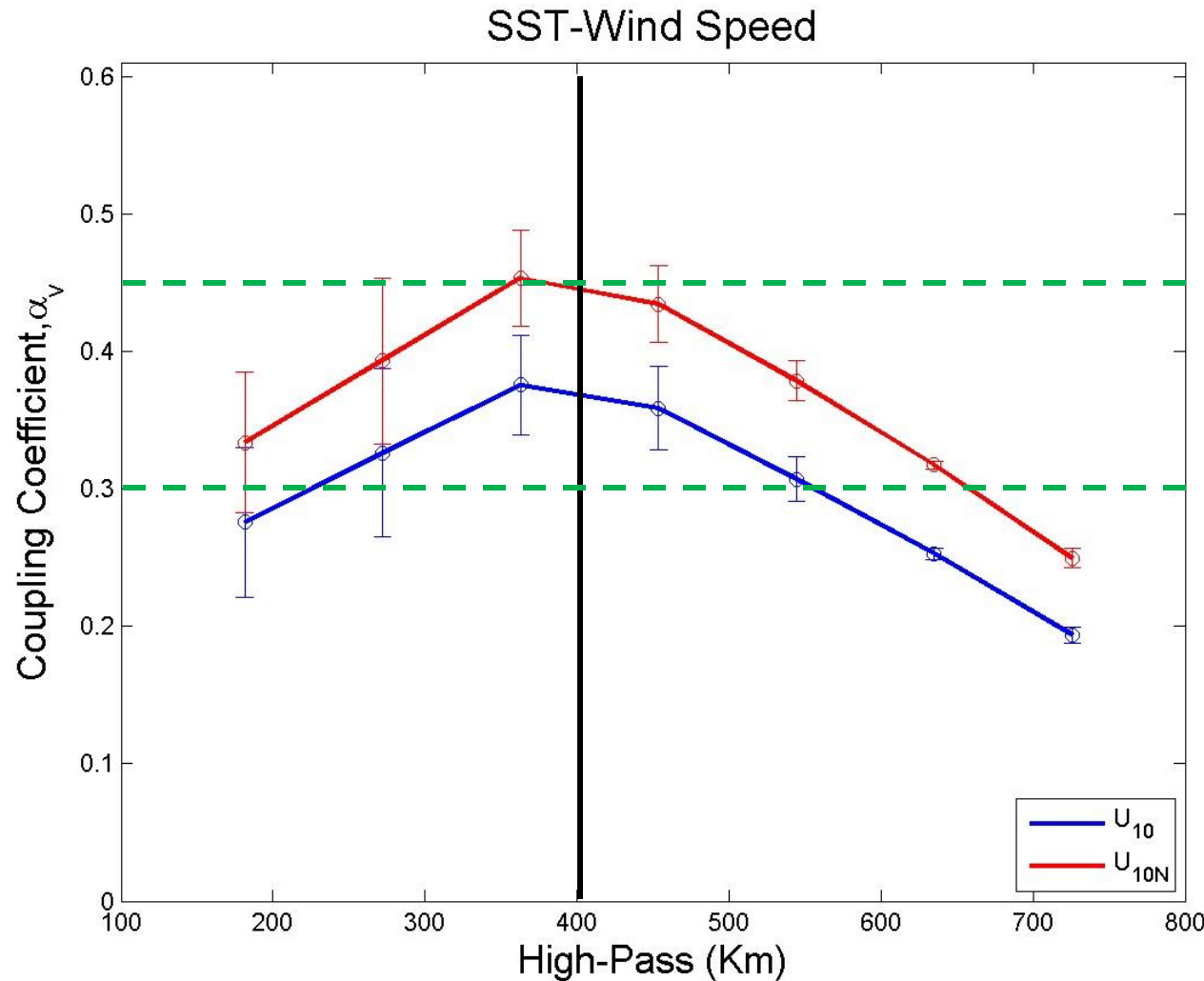
Phase speed 5-10 cm/s

O'Neill et al. (2010a)

Agulhas Current

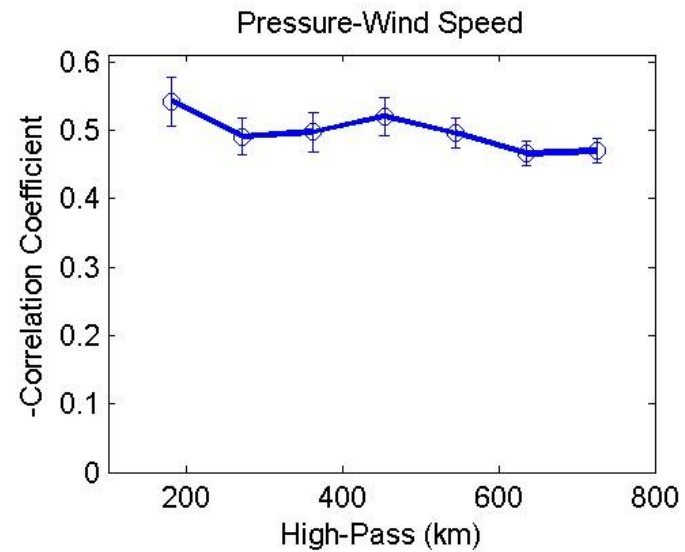
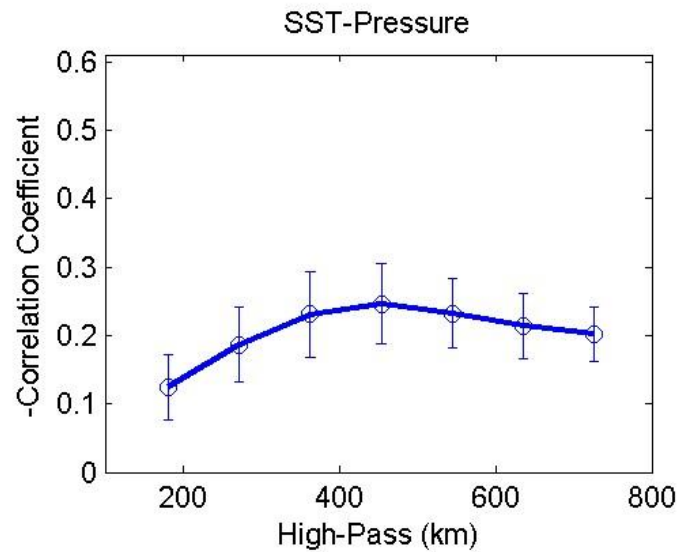
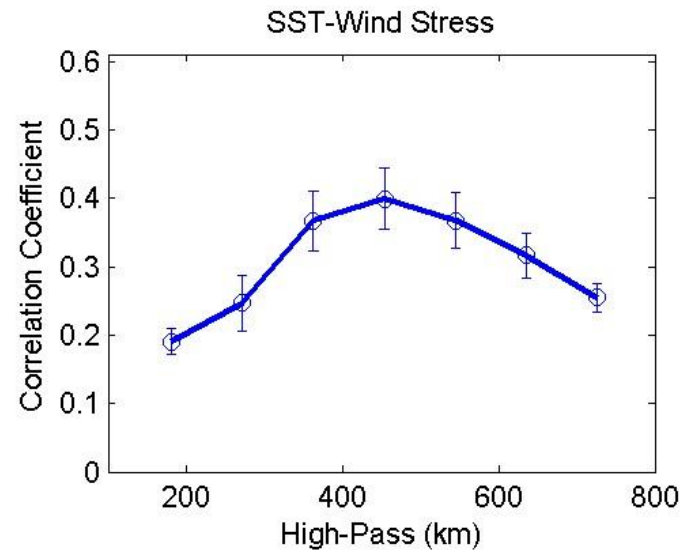
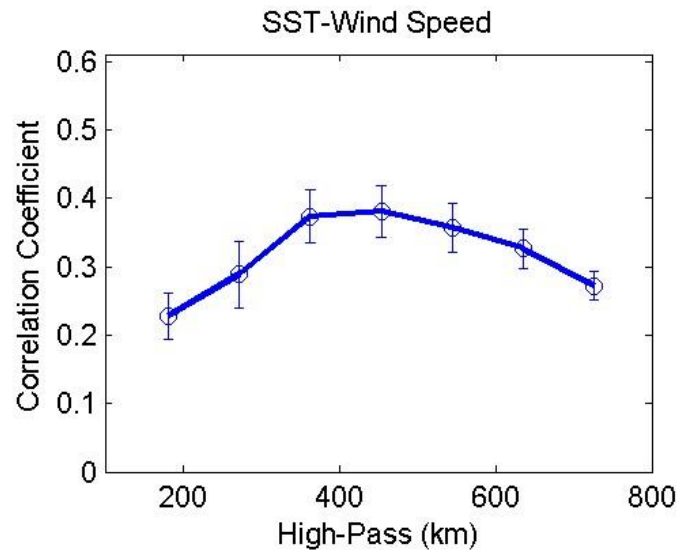
O'Neill et al. (2010b)

Gulf Stream

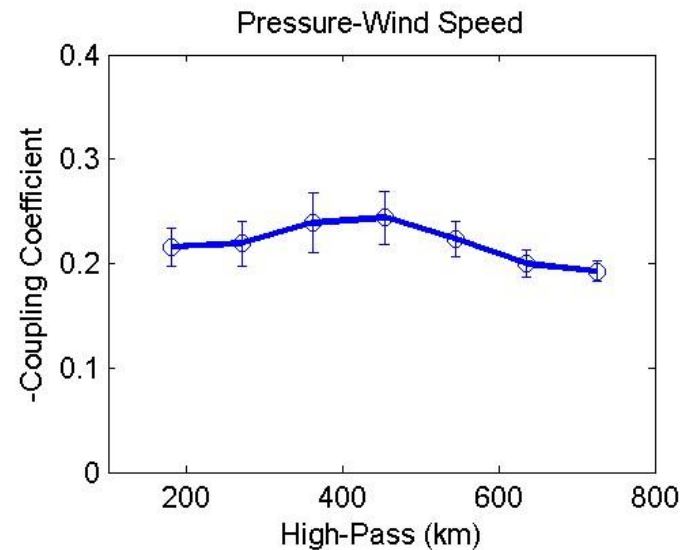
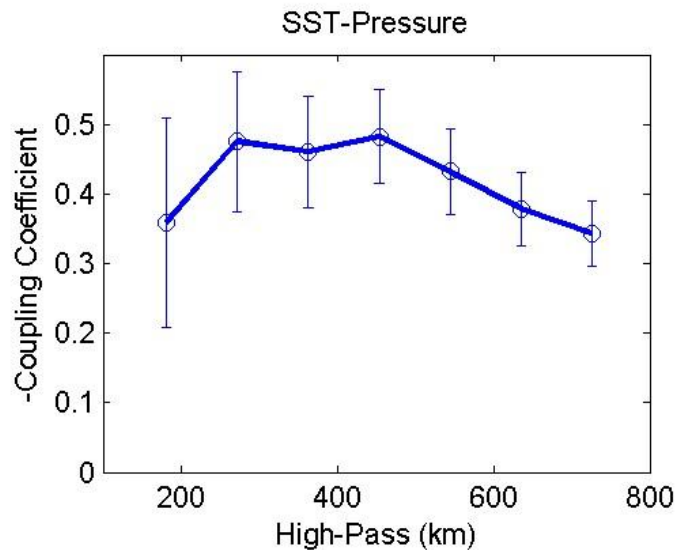
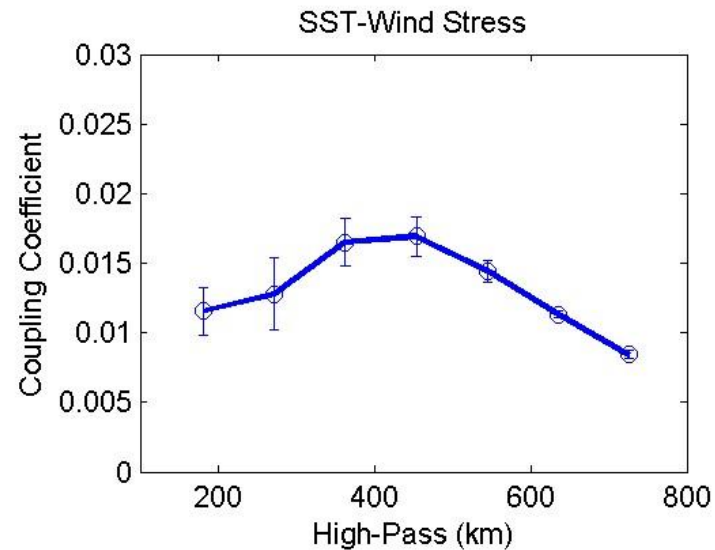
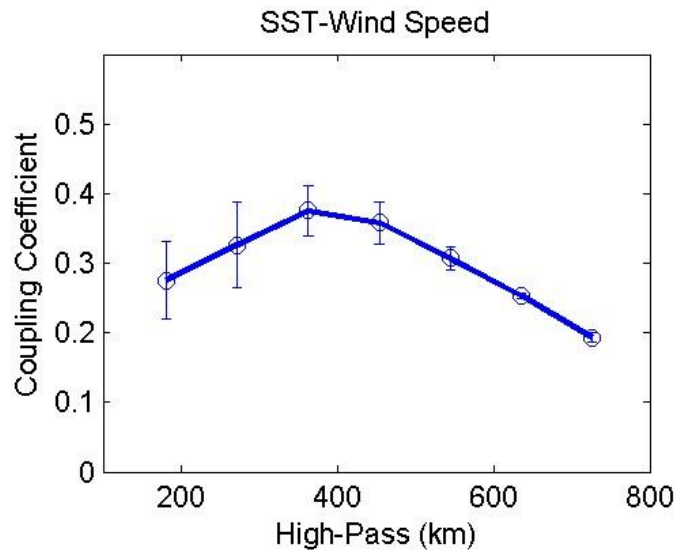




# Correlations



# Coupling Coefficients



Cool SST  
Low Wind  
High Pressure

Warm SST  
High Wind  
Low Pressure

# Summary

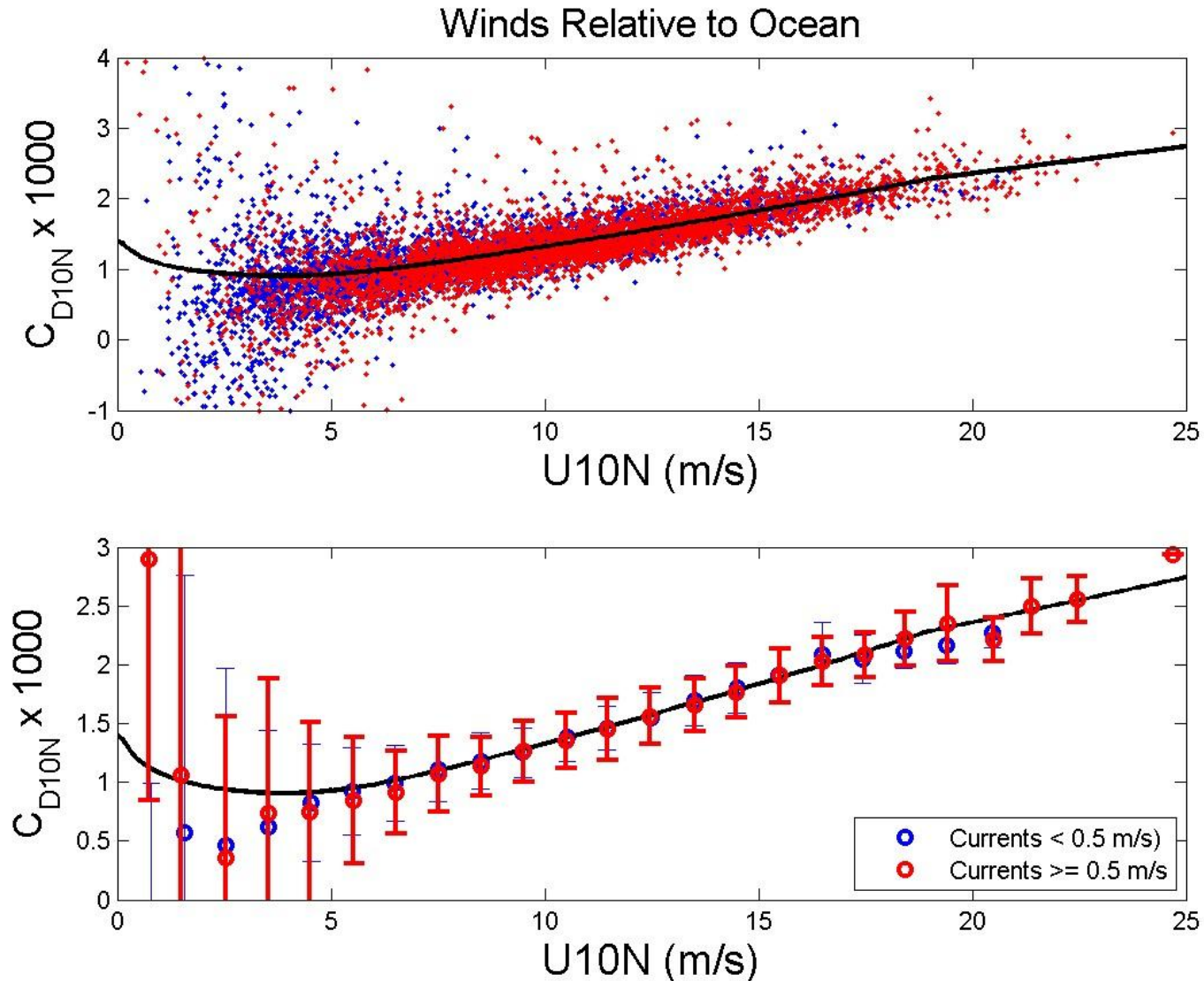
- Some of the variability in the QuikSCAT winds is due to adjustment of the neutral wind to changes in stratification and not changes in the actual wind speeds.
  - This variability obeys MO-Similarity in the mean.
  - Scatterometer winds represent ENW.
  - This effect enhances the gradient in neutral winds but not actual.
  - This accounts for about 20% of the coupling coefficient, which should be considered artificial, i.e., an artifact of measuring ENW.
  - However, significant variability in the QuikSCAT winds is not explained by this effect
- The one-buoy approximation of the coupling coefficients is in good agreement with previous studies.
  - The coupling coefficient depends on the scale of the SST variability.
  - In the Gulf Stream regions, the variability is driven by Gulf Stream meanders.
  - Results suggest that pressure adjustment is mainly responsible for observed coupling between wind speed/wind stress and SST at these scales.
- Next: Scatterometers as stressmeters





# Relative Velocity

$$C_{DN}(z/z_o) = \frac{-\overline{uw}}{\Delta U_N G} = \left( \frac{\kappa}{\ln(z/z_o)} \right)^2$$



# Future Research w/ Surface Stress Measurements

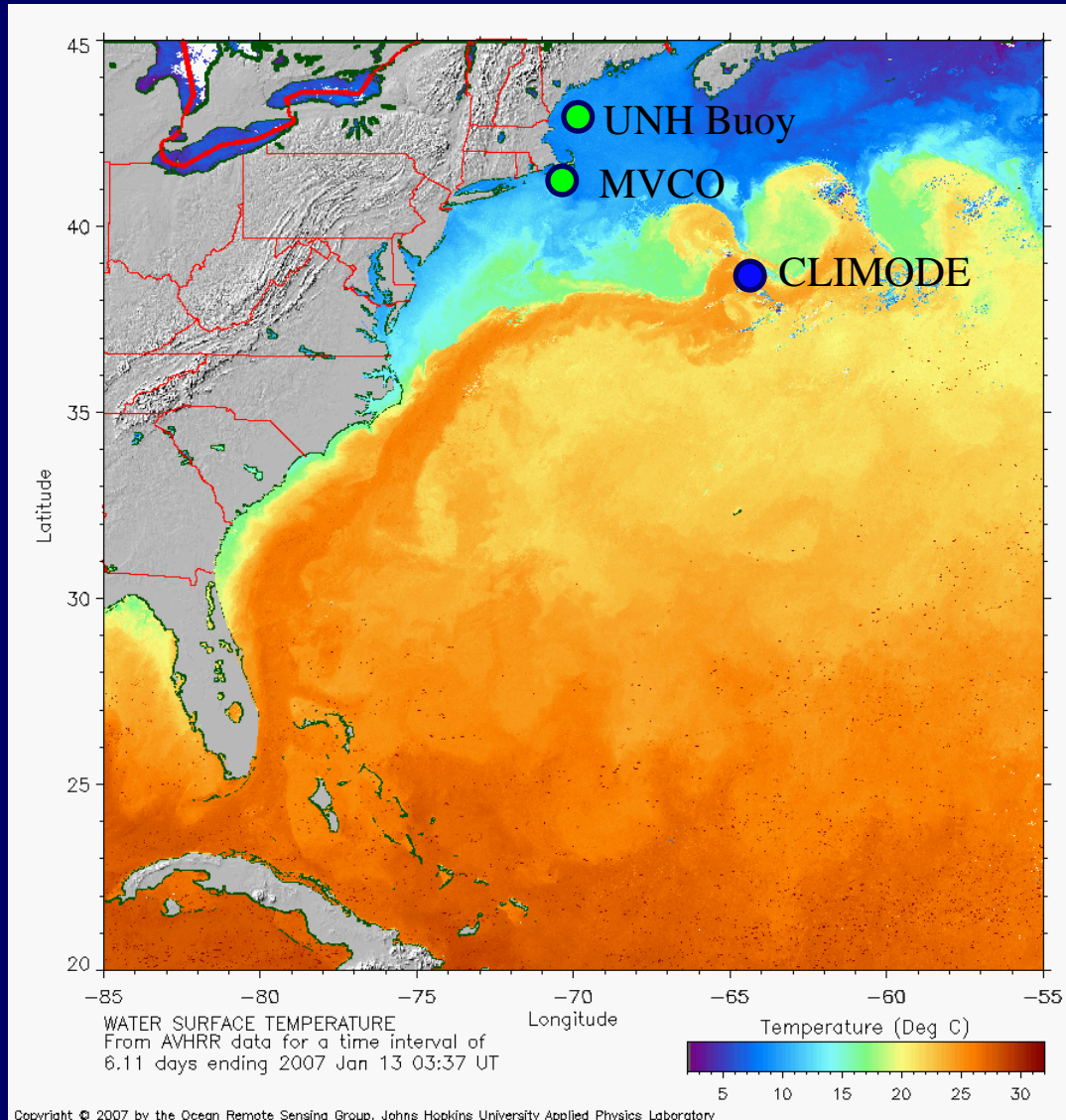
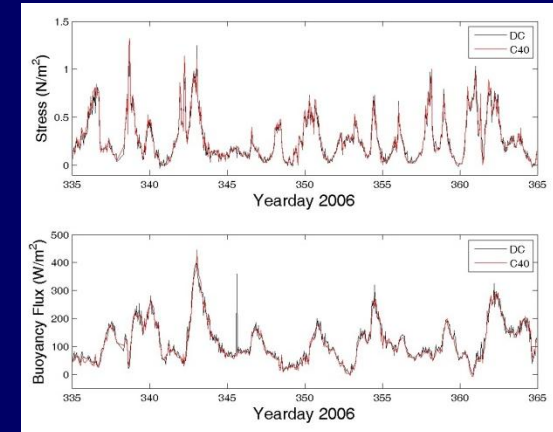


Image courtesy of JHU/APL



MVCO: 2002-2013

UNH Buoy: 2009-2012

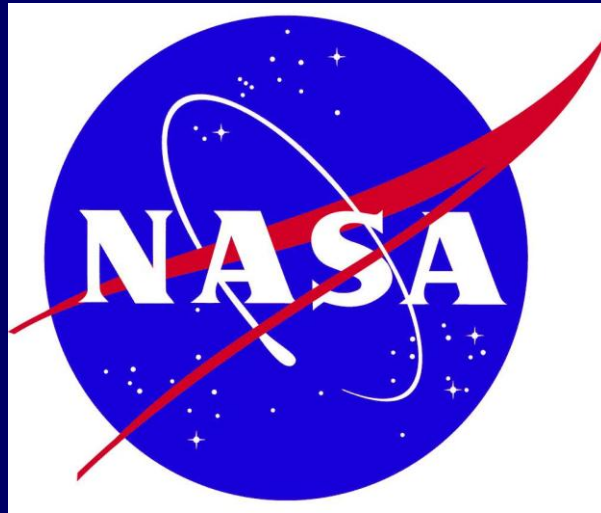
CLIMODE: 2005-2007

SPURS: 2012-2013

SPURS ●●

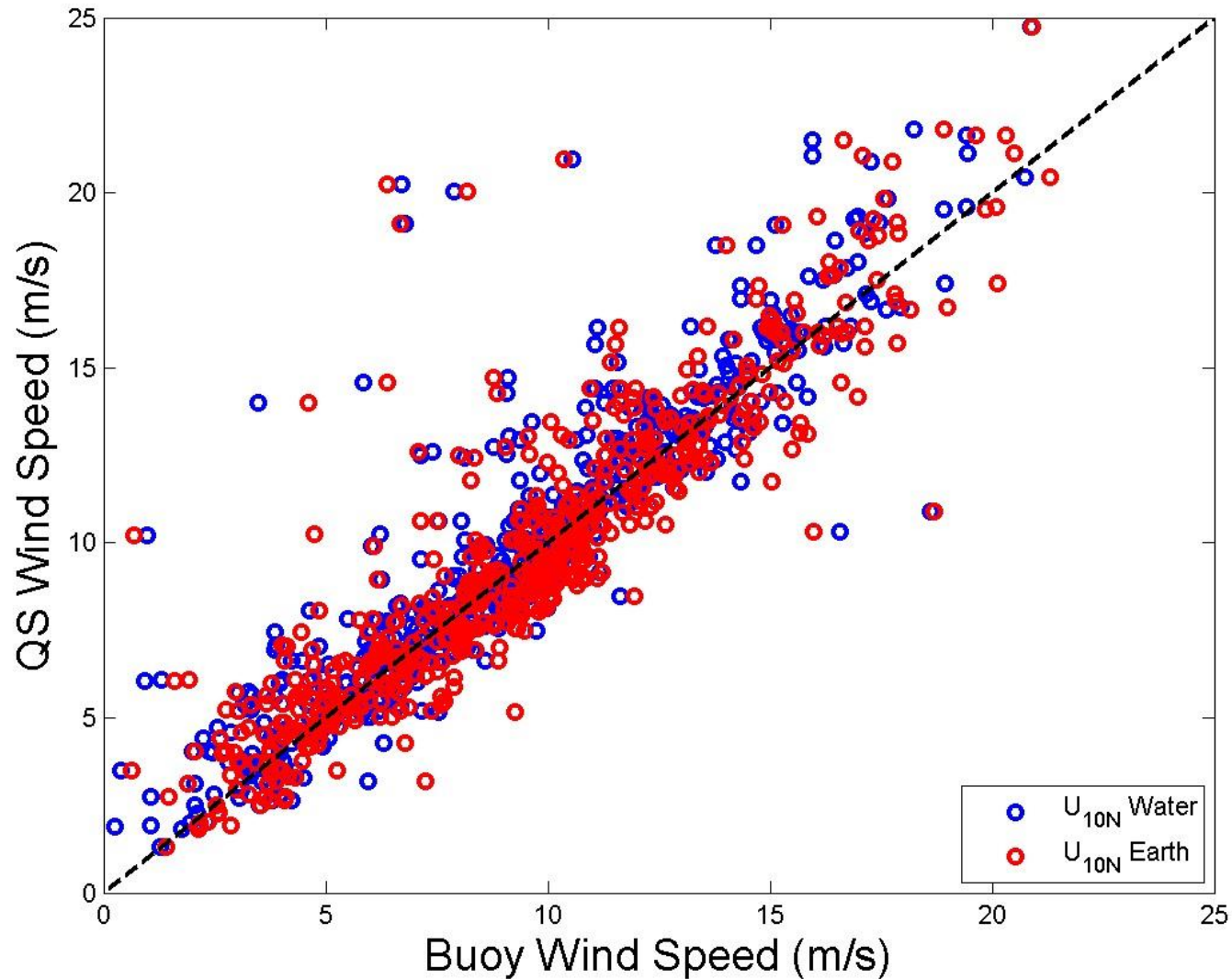
- Exploit scatterometers as stressmeters.
- Develop model functions using direct covariance stress with ASCAT and OSCAT.
- Explore, e.g., sensitivity to currents, SST and long waves.

Thanks to NASA for supporting this research.

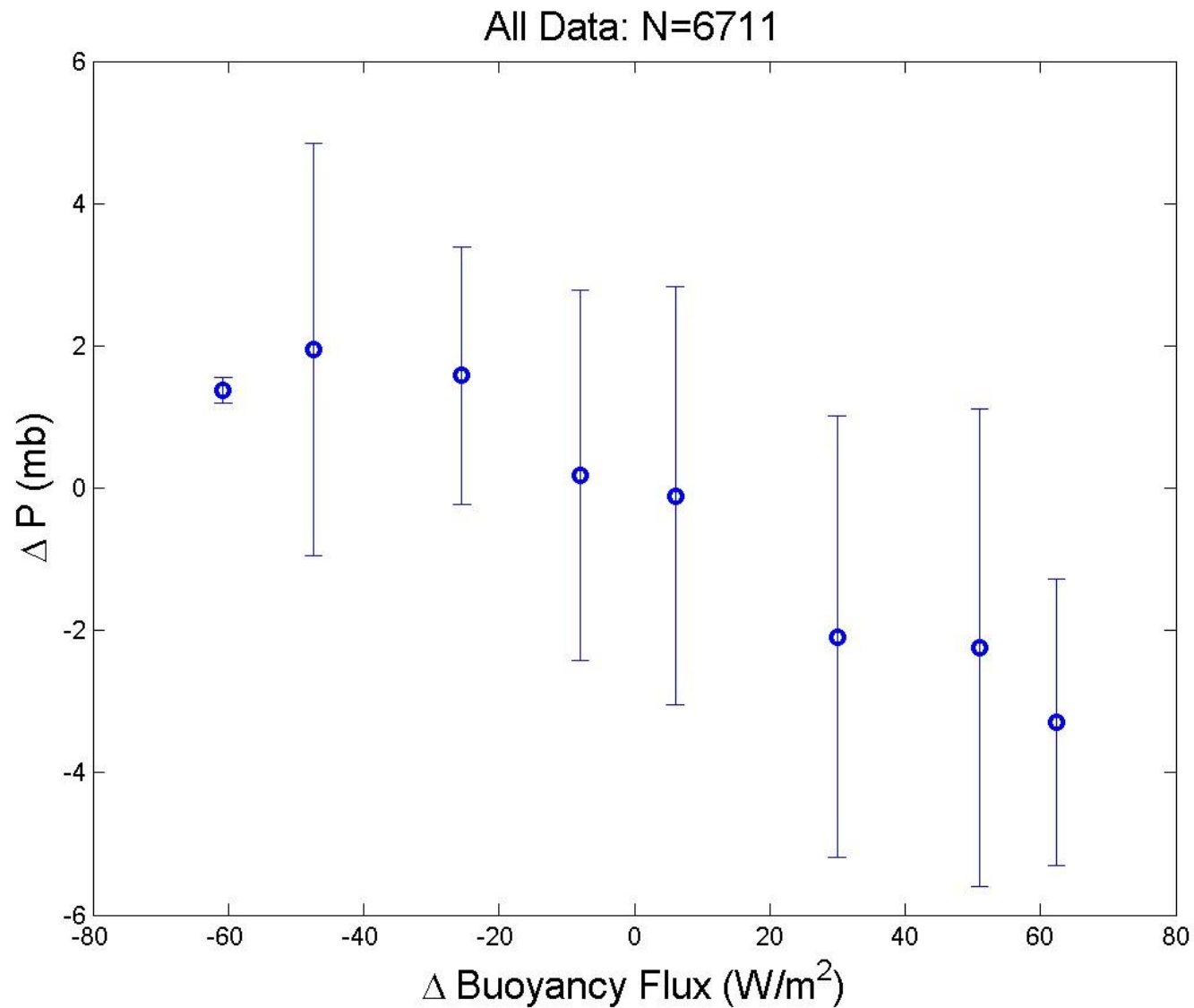




# QuikSCAT vs. Buoy Wind Speeds



# Correlations with Buoyancy Flux

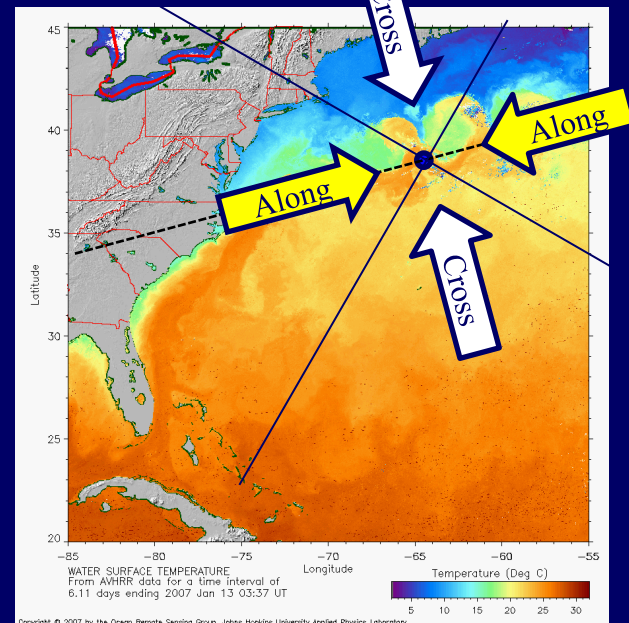
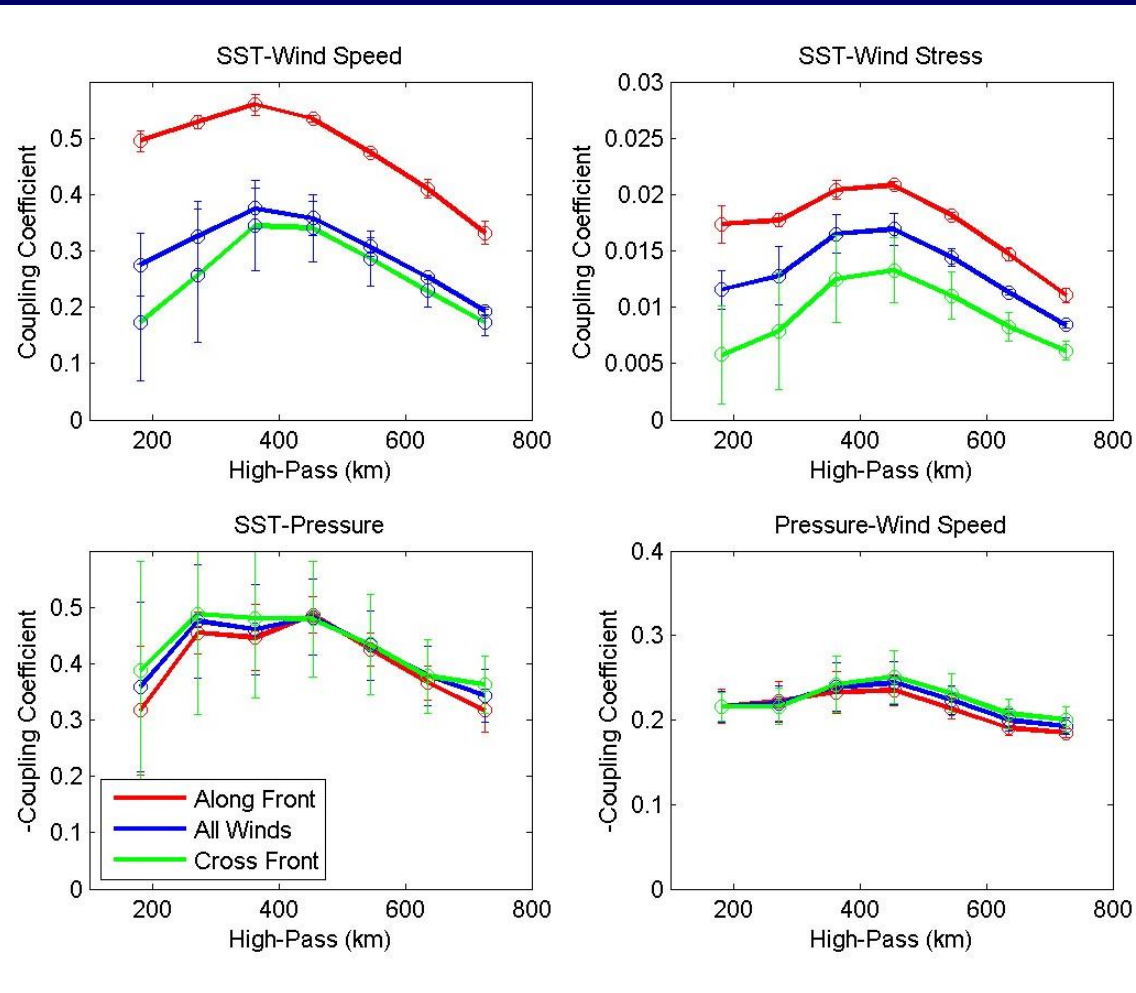


# Cross Frontal Velocity & Advective Adjustment

Length scale over which the PG changes (Spall 2007):

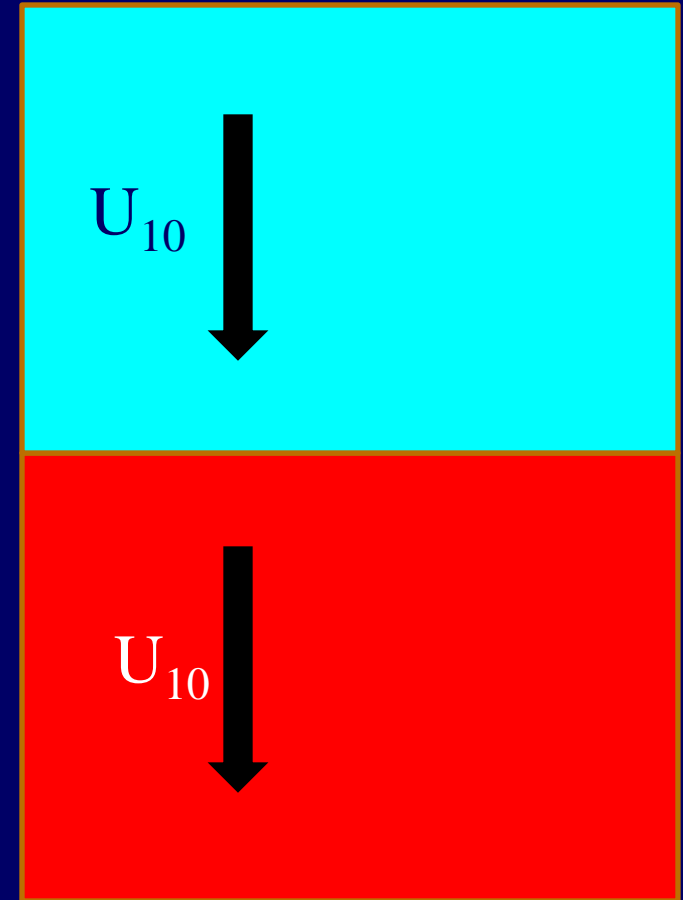
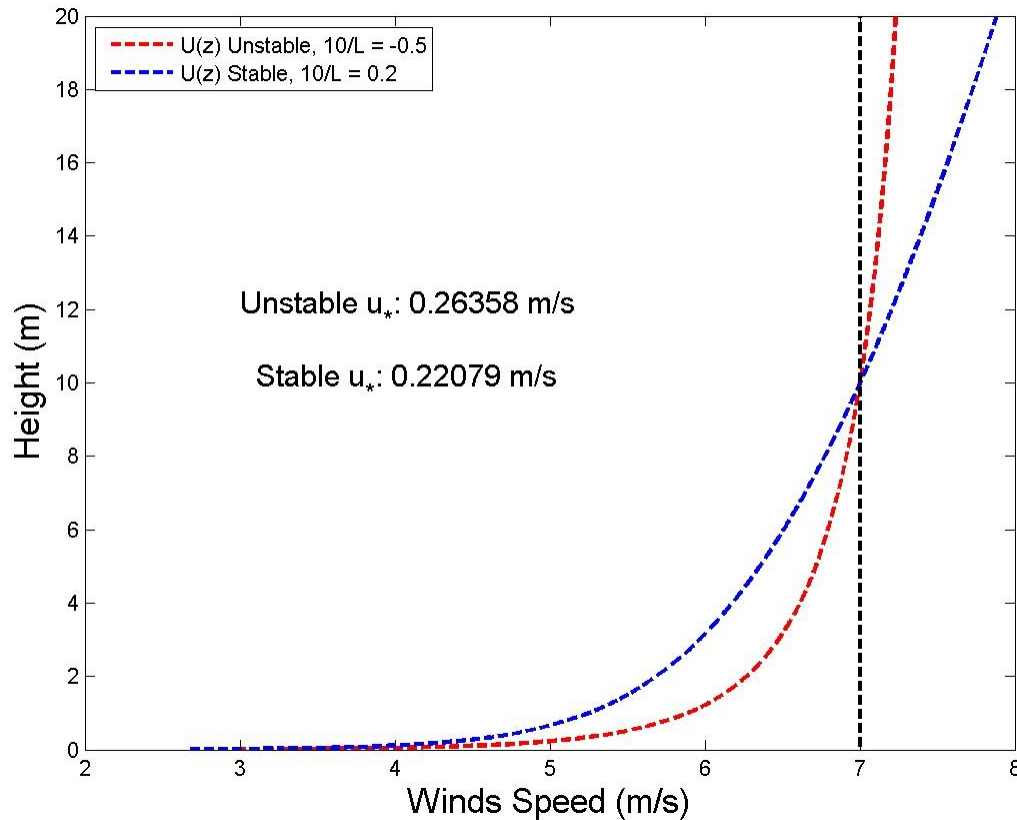
$$L_p = \frac{U_{cross} H^2}{K_m} \propto \frac{U_{cross}}{f}$$

The shorter this scale, the quicker the PG can adjust to the SST gradient and the more important it becomes in driving the winds.



# QuikSCAT vs. Buoy Wind Speeds

## “Surface Layer Adjustment”

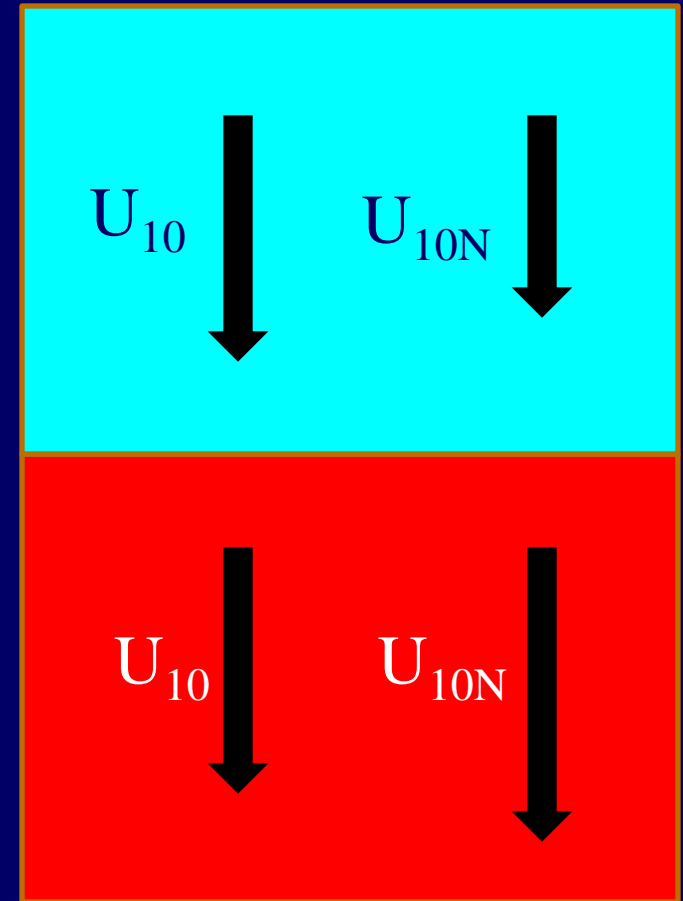
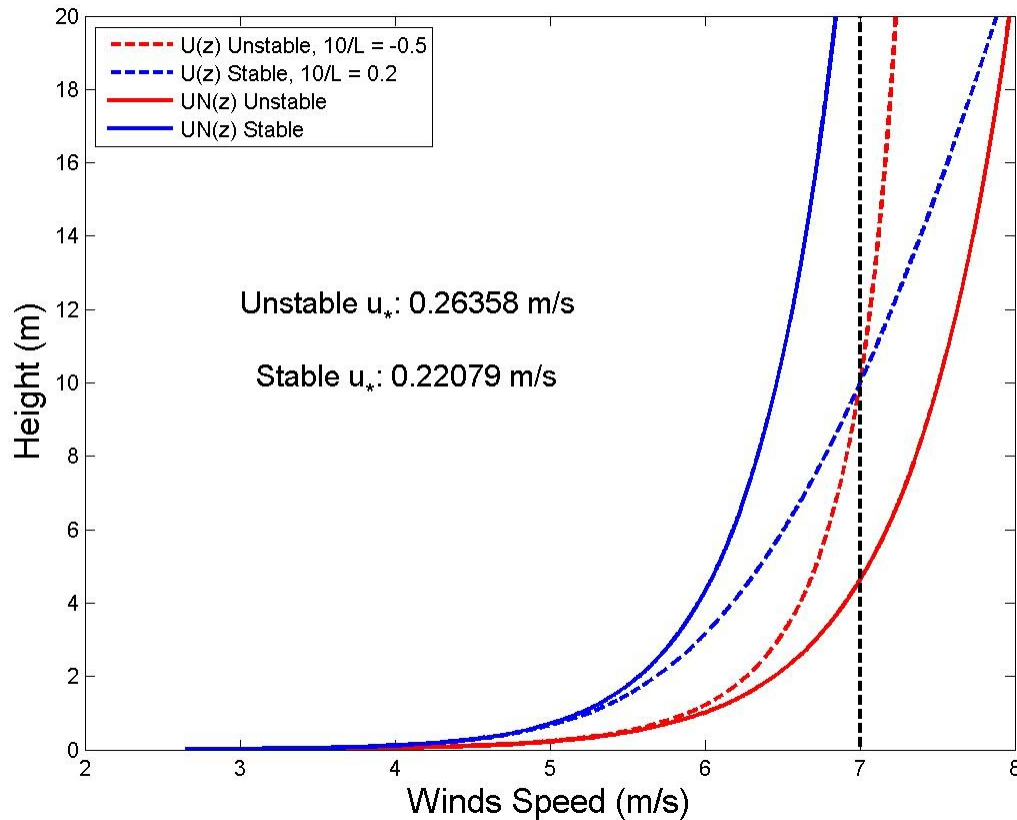


$$U(z) = u_* / \kappa [\ln(z/z_o) - \psi_m(z/L)]$$



# QuikSCAT vs. Buoy Wind Speeds

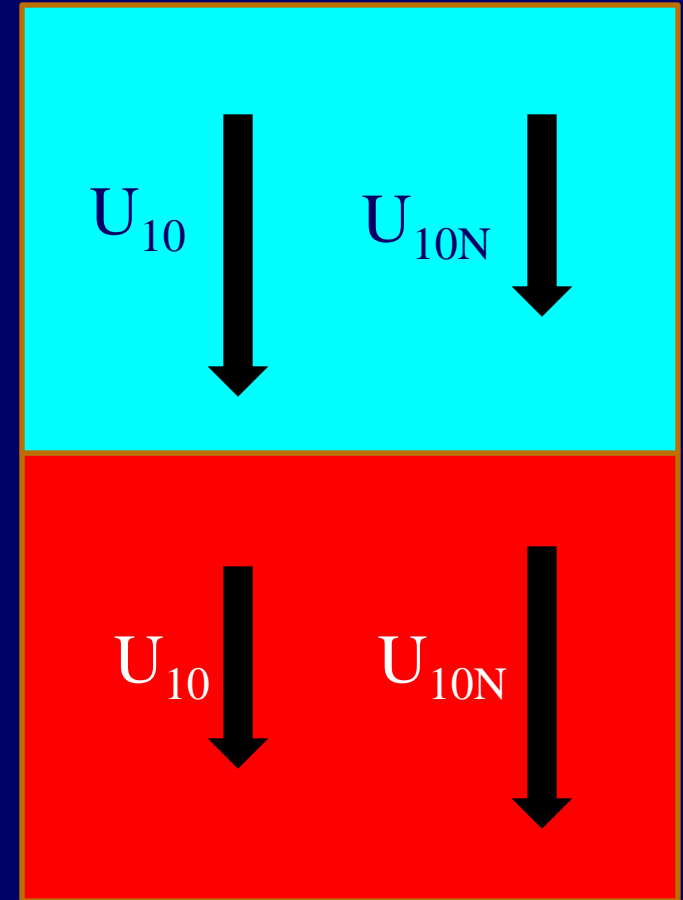
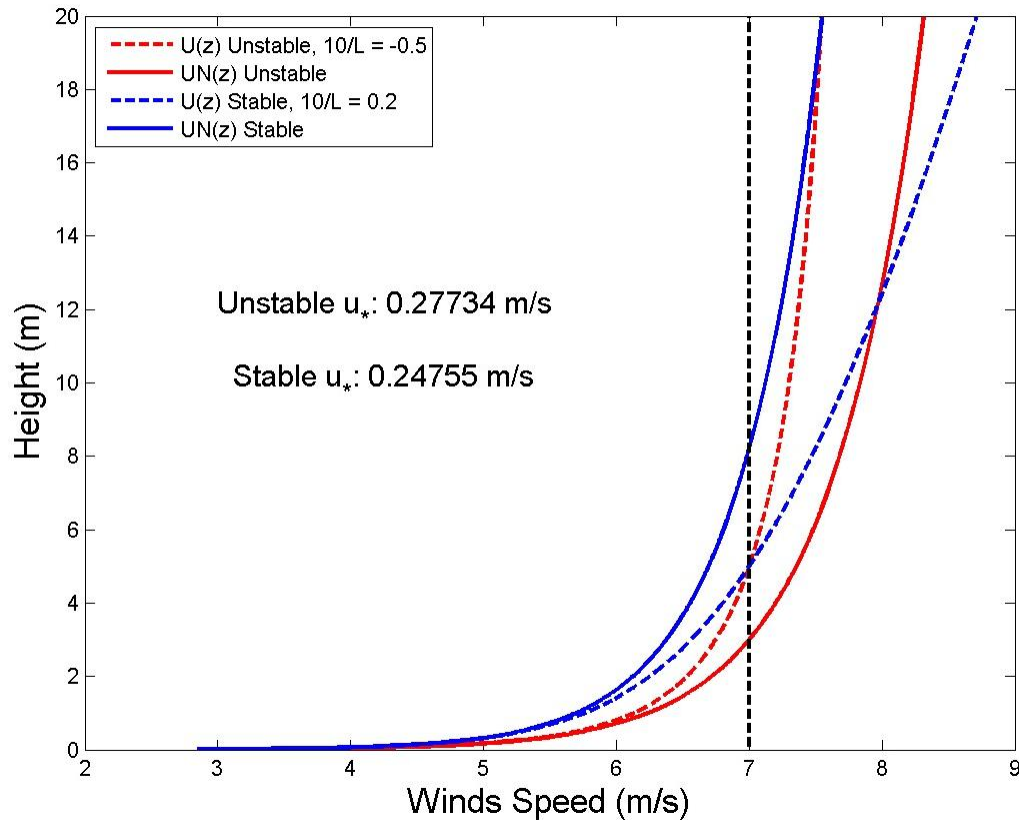
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