

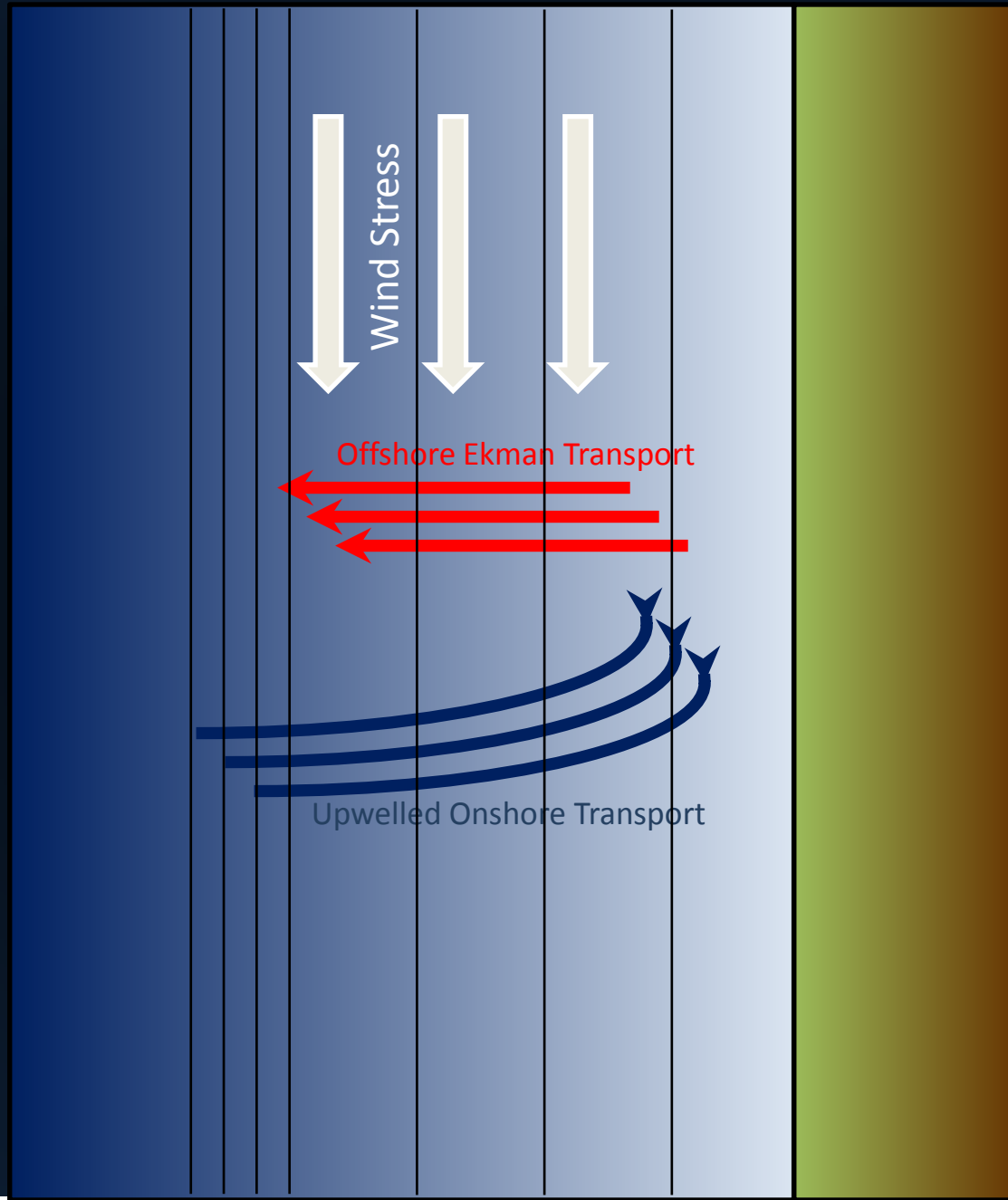
Spatial Variability of Upwelling on a Shelf with Complex Geometry inferred from Scatterometry

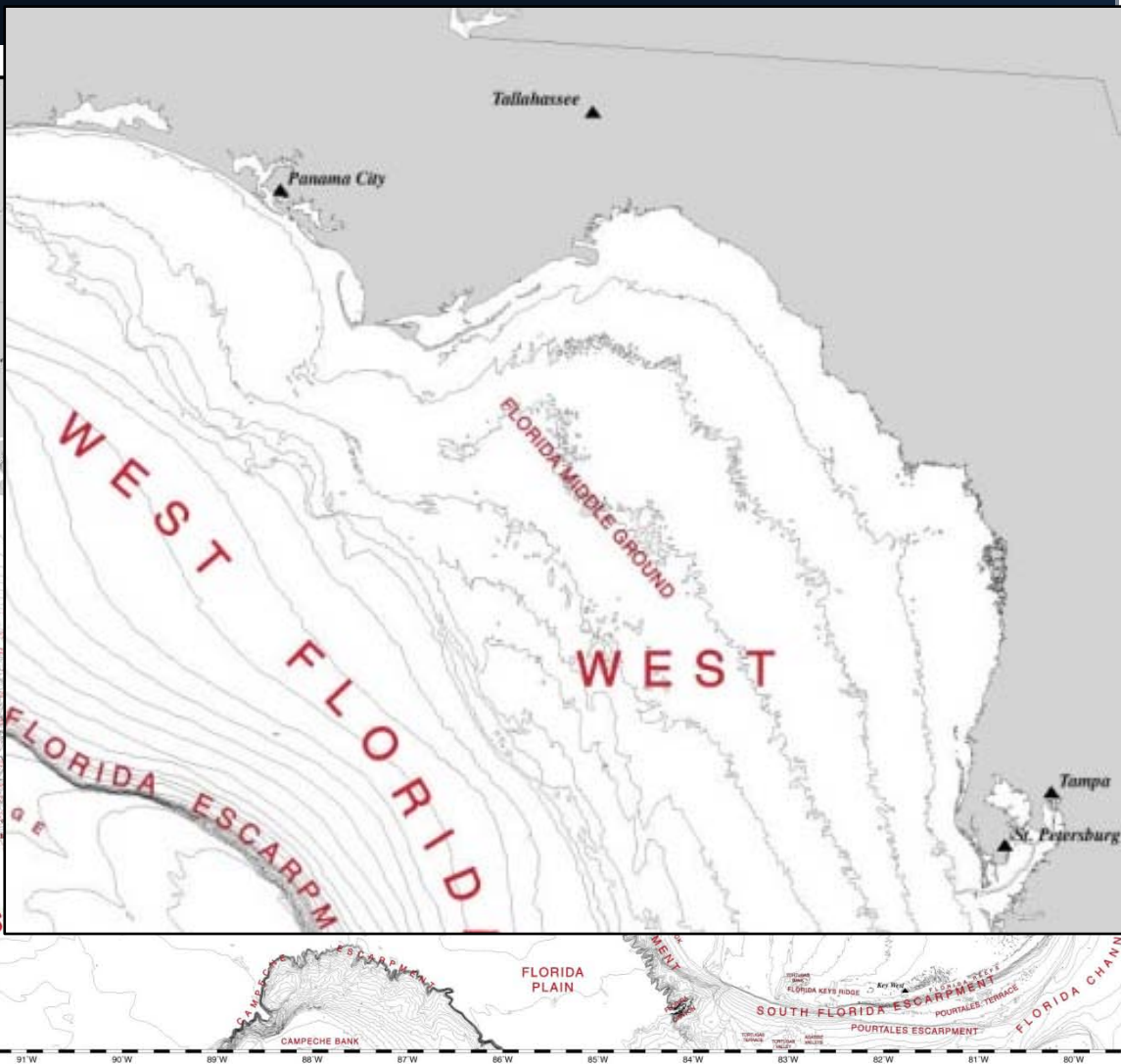
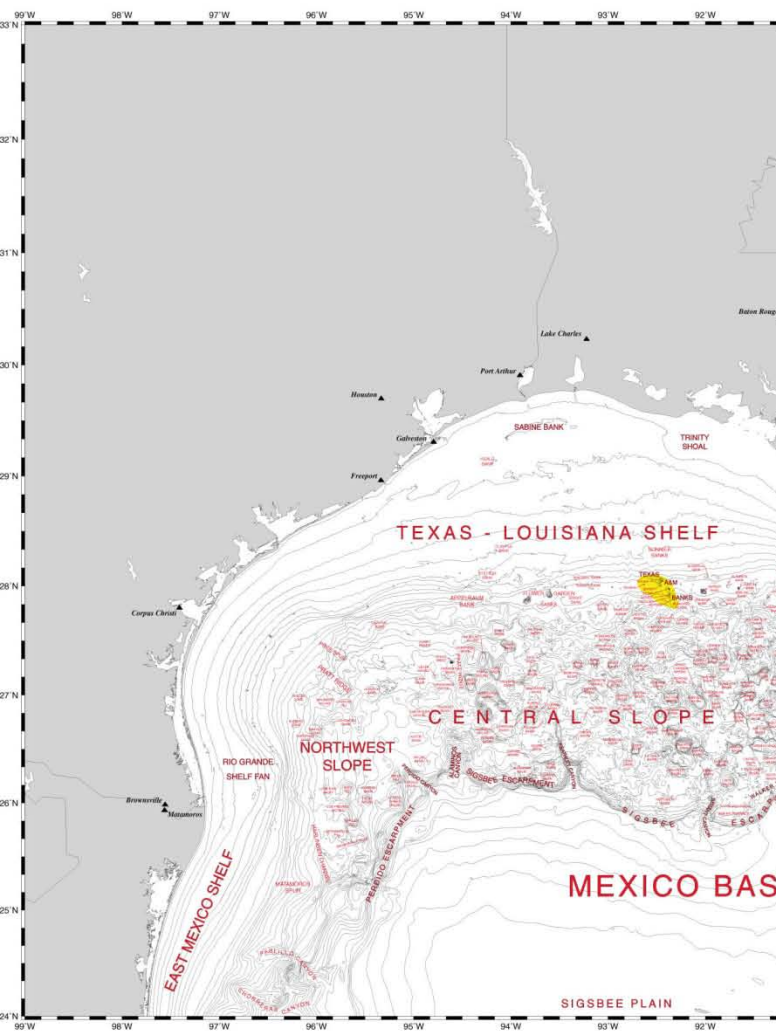


Steve Morey

Center for Ocean – Atmospheric Prediction Studies
The Florida State University

Idealized Coastal Upwelling





0 100 200 300
km

Mercator Projection Scale 1:1,800,000 at the equator

Contour interval 100 meters, with supplemental contours at 10 meters
on the continental shelves and the Hatteras Abyssal Plain

This map is intended to show ocean floor topography
and is not to be used for navigation.

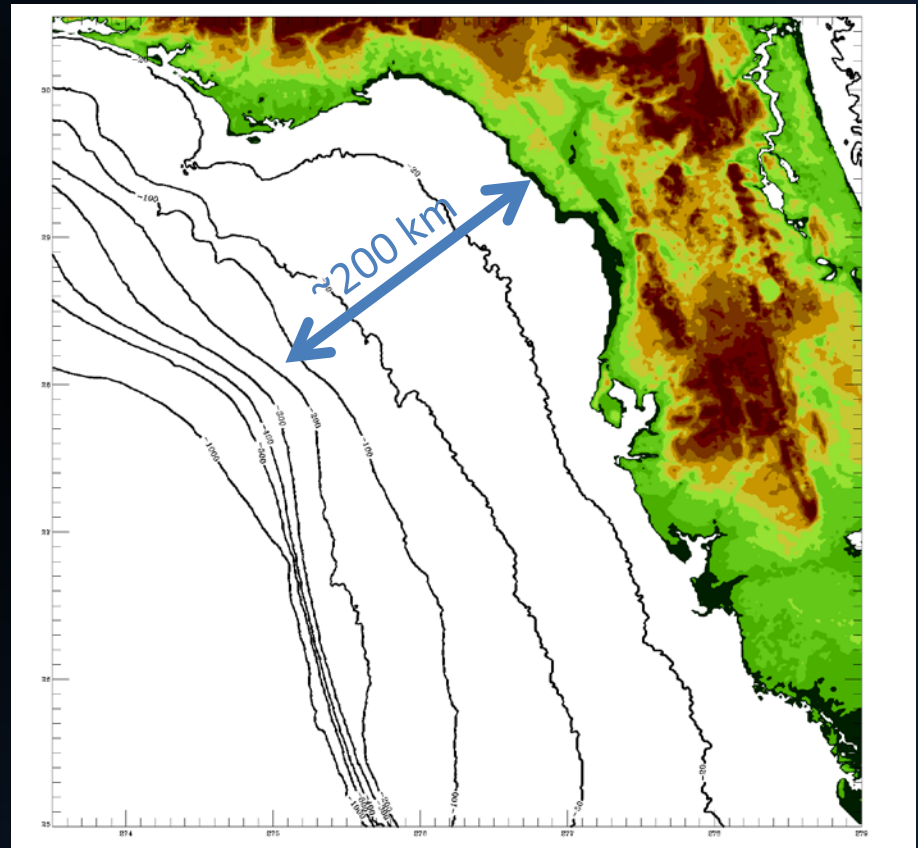
Upwelling on the northern West Florida Shelf

QuikSCAT wind pseudostress is projected onto the along-isobath direction

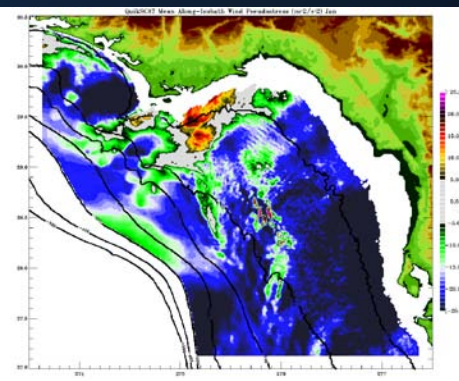
- Positive in the topographic wave propagation direction → downwelling

1. The topographic gradient is computed over a $\sim 30\text{km}$ length scale (comparable to R_d)
2. QuikSCAT Level 2B winds are converted to pseudostress $|\mathbf{u}|\mathbf{u}$
3. Pseudostresses are bin-averaged and projected onto the local along-isobath direction

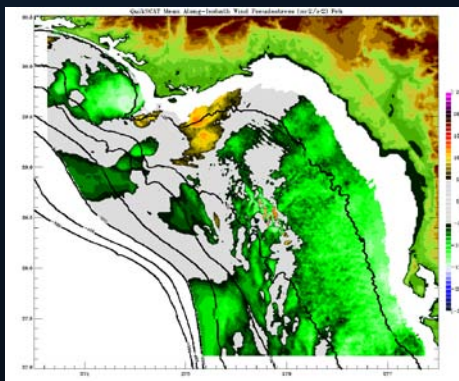
(Computed between 10m and 300m isobaths)



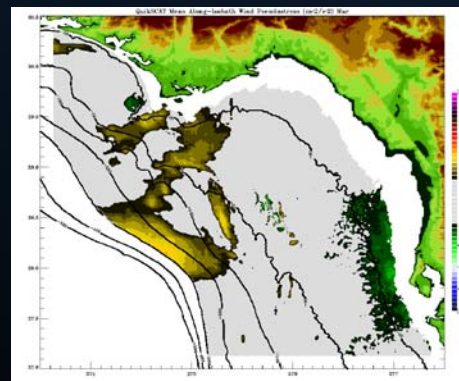
Jan



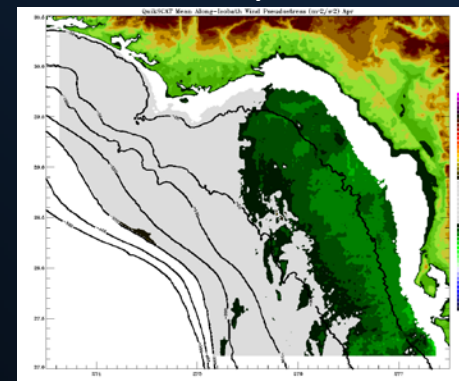
Feb



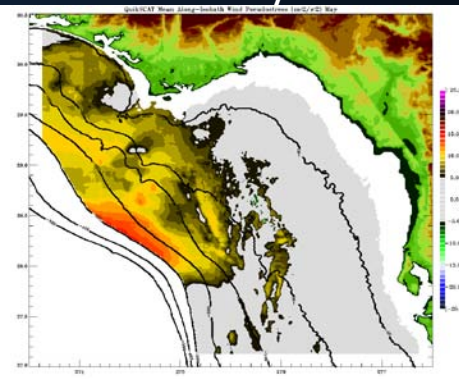
Mar



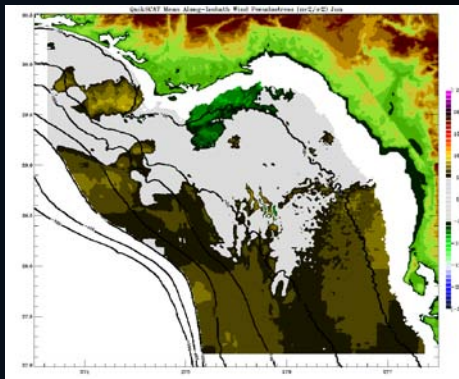
Apr



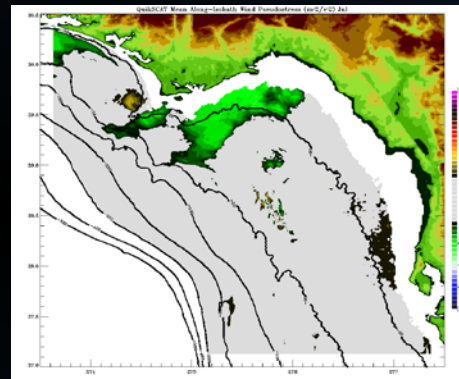
May



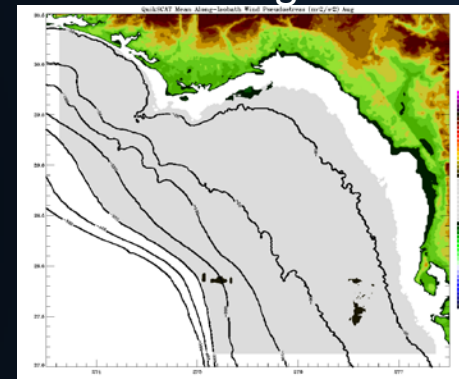
Jun



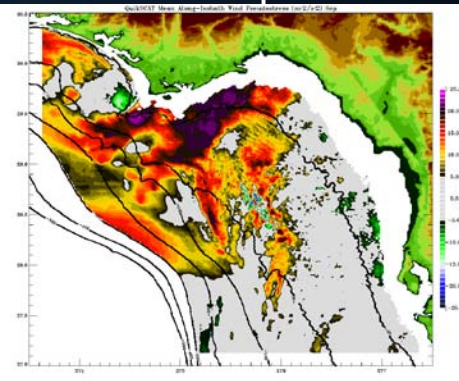
Jul



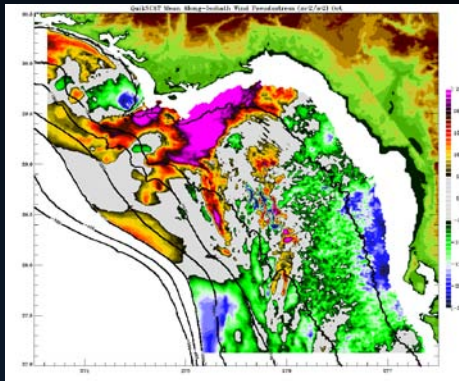
Aug



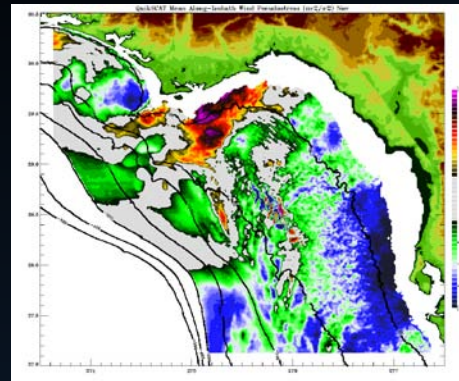
Sep



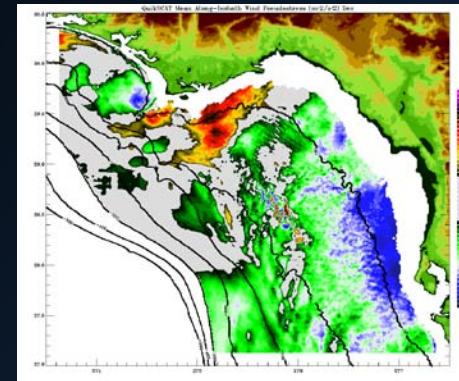
Oct



Nov

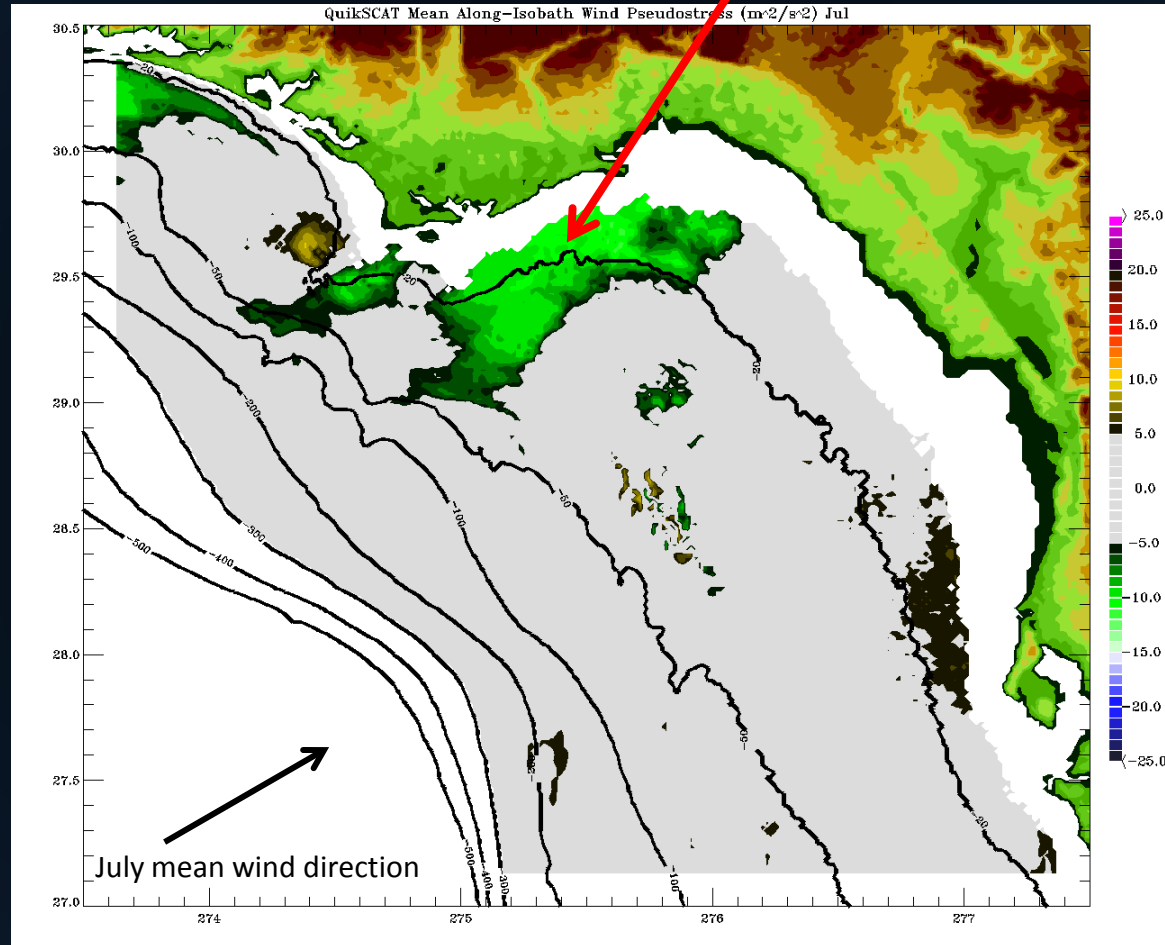


Dec



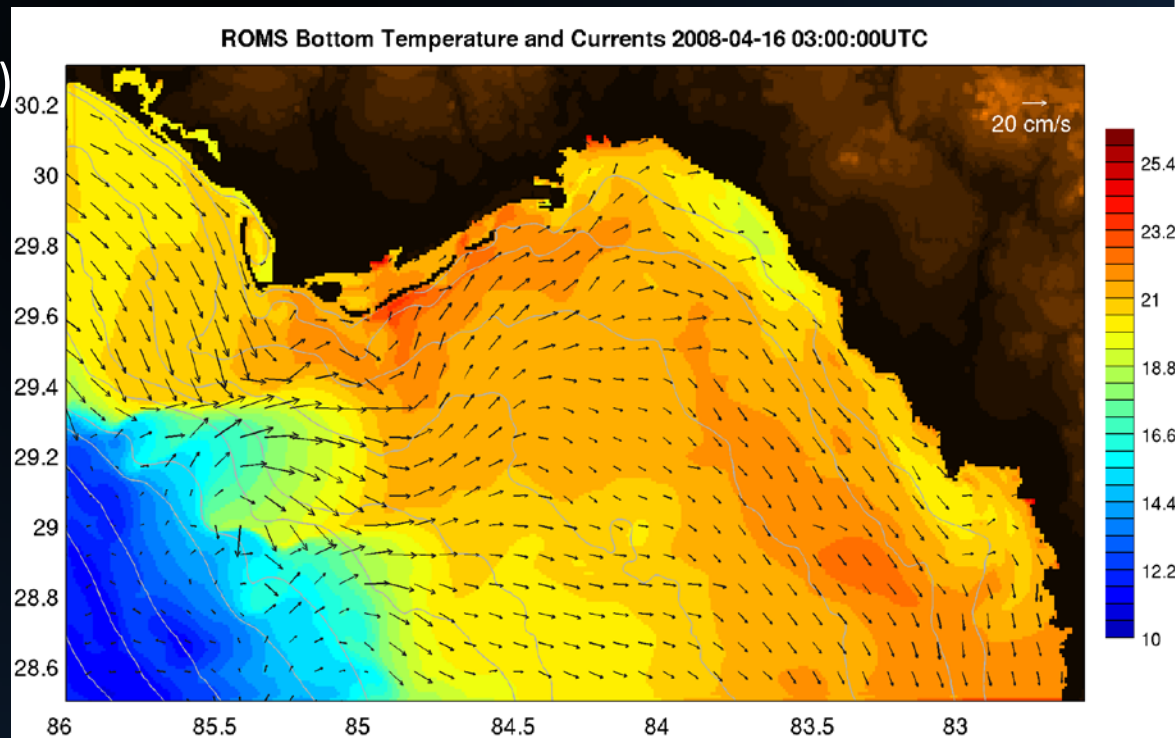
July Climatology

Upwelling-favorable winds only over the northern and western Apalachee Bay region



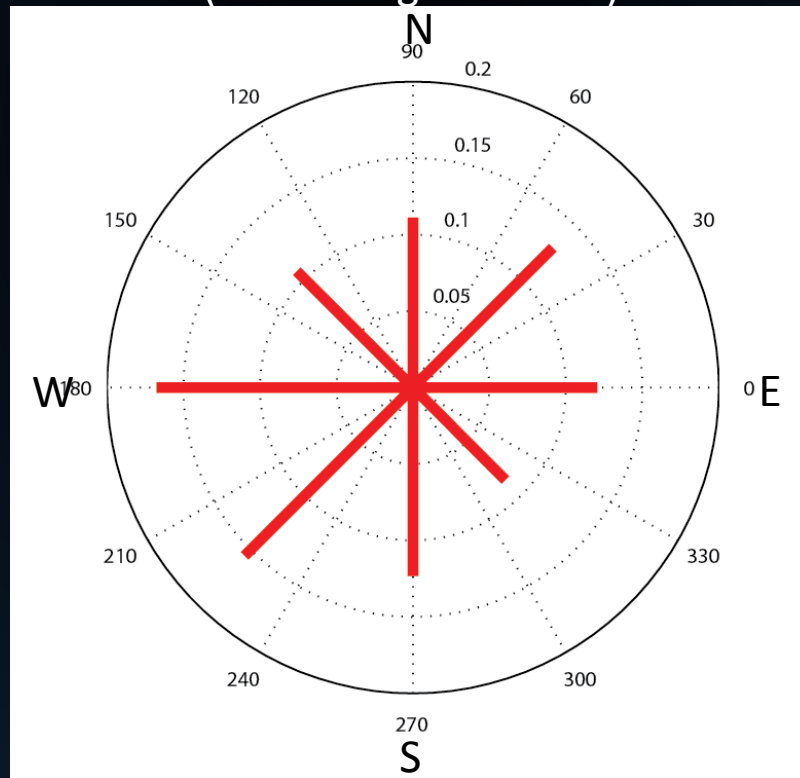
Numerical Simulation

- ROMS configured for northeastern Gulf of Mexico
 - 30 arcsec (~ 800 m) horizontal resolution
 - 30 vertical stretched terrain-following layers
- Nested within $1/25^\circ$ HYCOM Gulf of Mexico Hindcast
- Surface forcing – COARE 3.0 fluxes with CFSR atmospheric variables
- 5-year simulation (2005-2009)
- 3-hourly hourly output

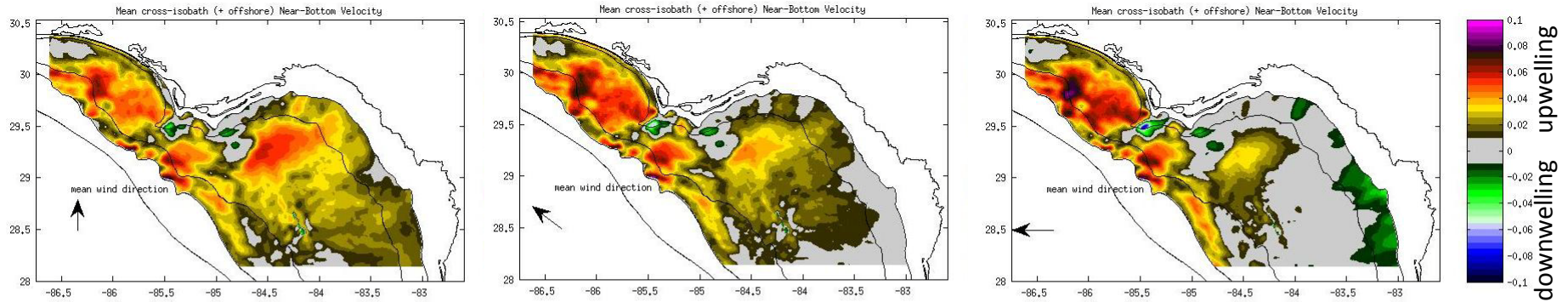


- Near-bottom currents are subset in time based on the spatially-averaged wind direction (in 45° bins)
- Near-bottom velocity and wind fields are temporally averaged for each directional bin
- Near-bottom velocity is projected onto the topographic gradient
- Winds are projected along isobaths

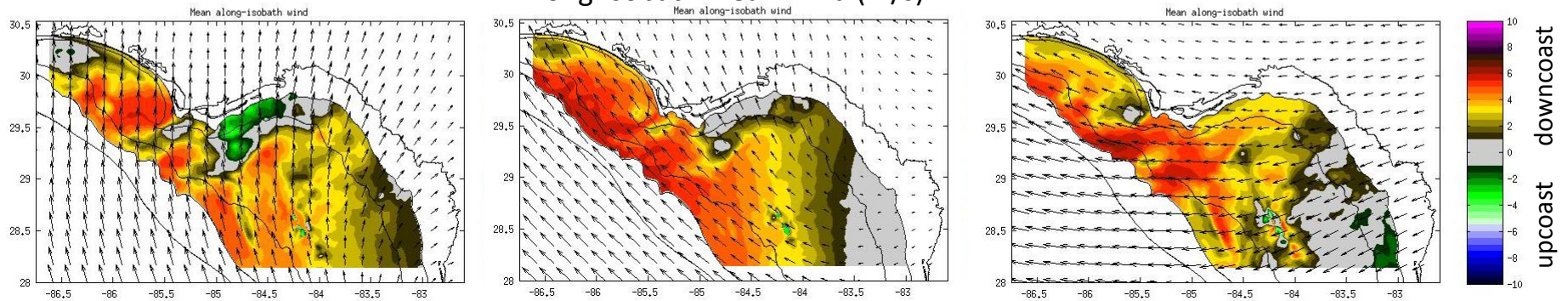
Wind Directional Histogram
(meteorological sense)



Near-bottom cross-isobath mean velocity (m/s)

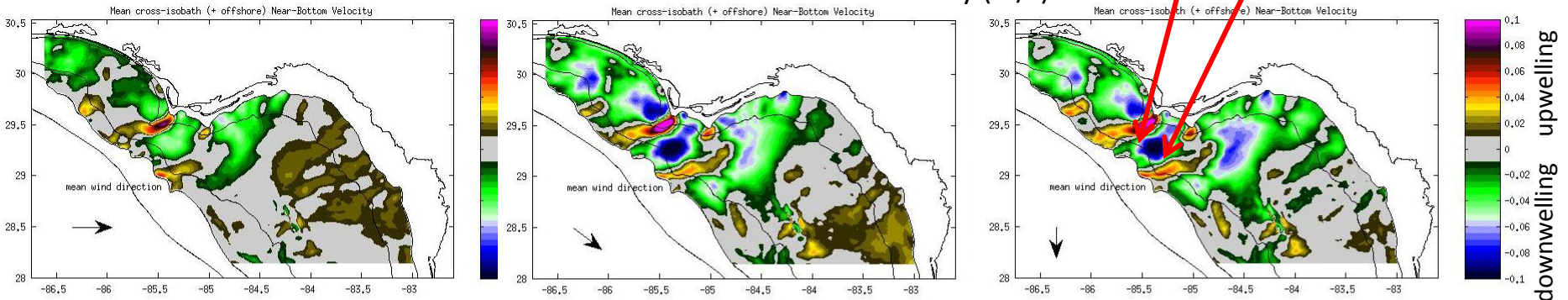


Along-isobath mean wind (m/s)

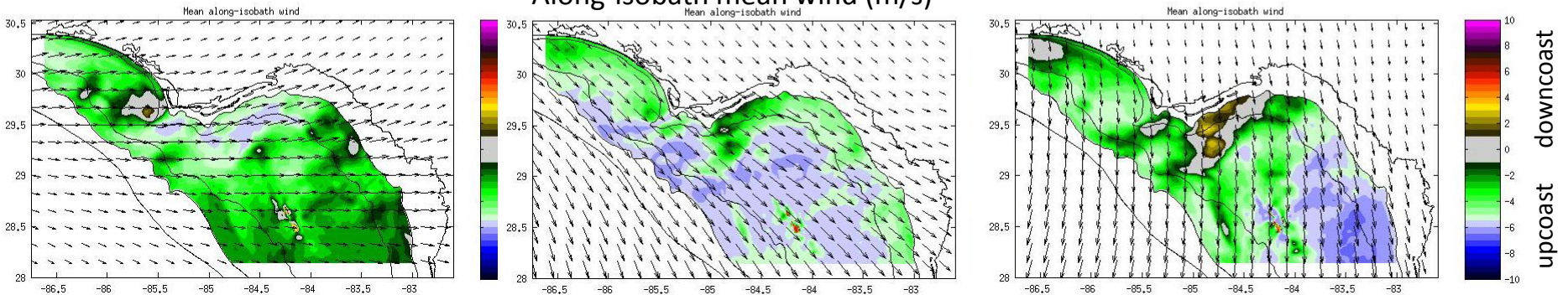


Regions of alternating
up-isobath and down-
isobath transport

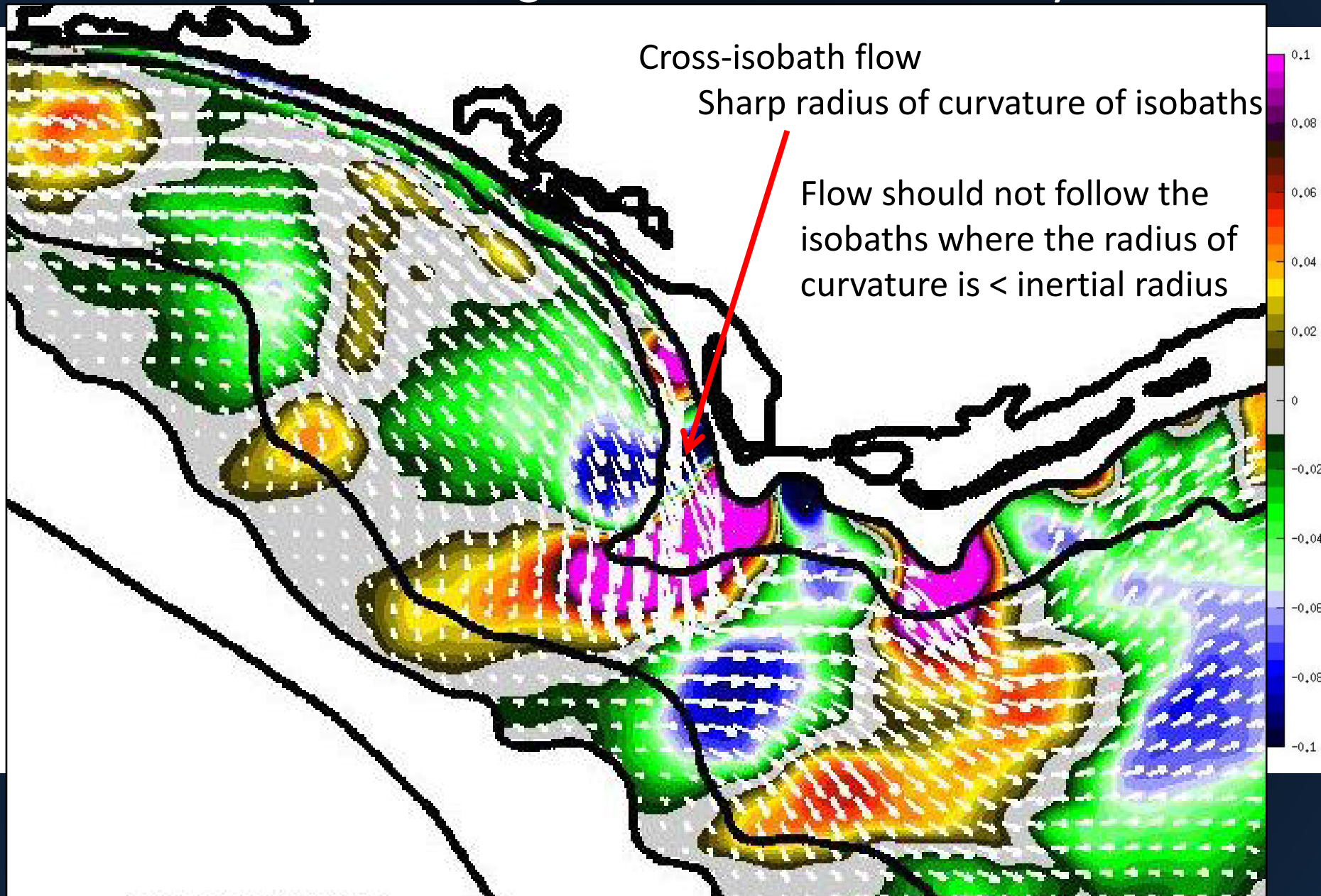
Near-bottom cross-isobath mean velocity (m/s)

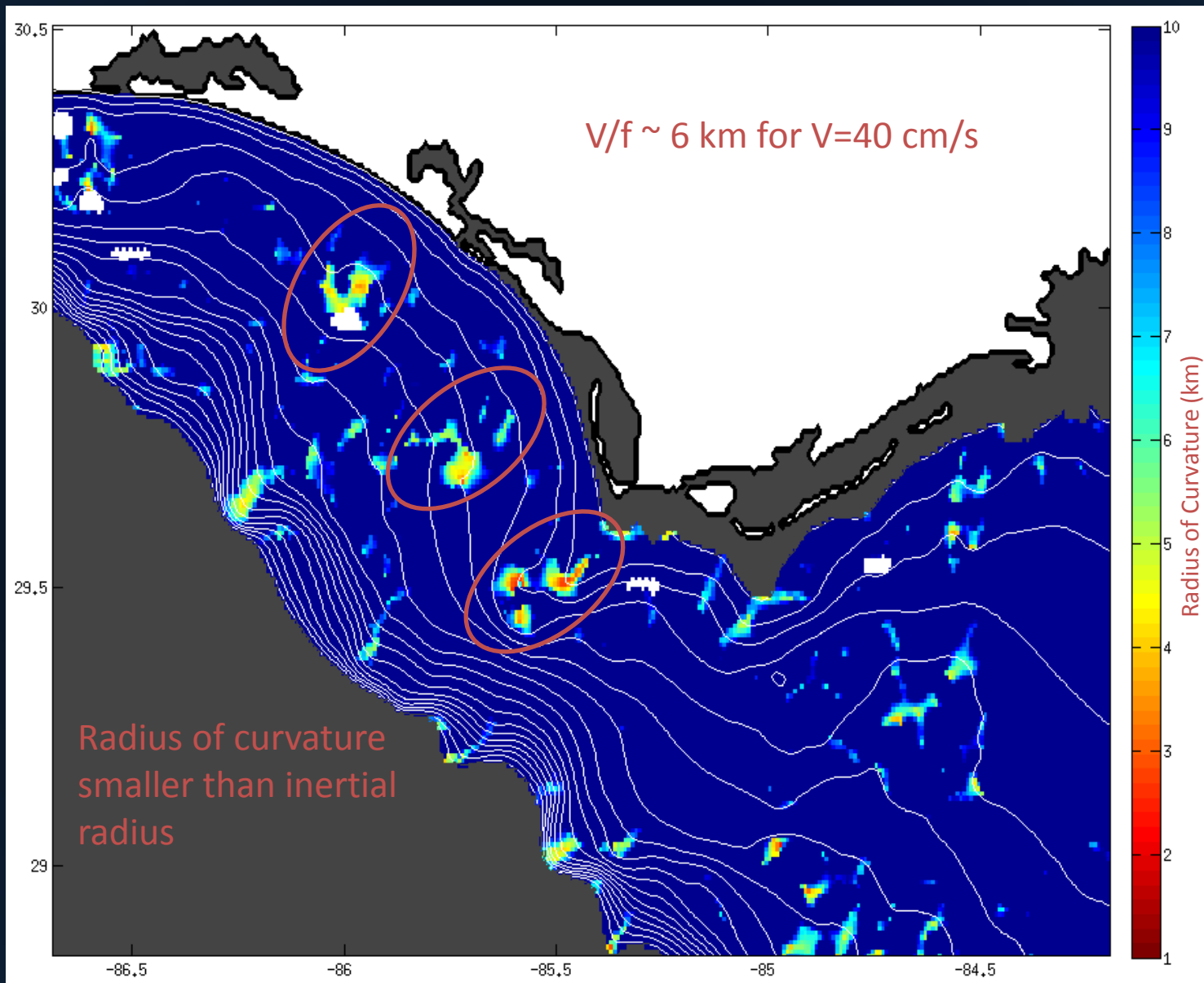


Along-isobath mean wind (m/s)

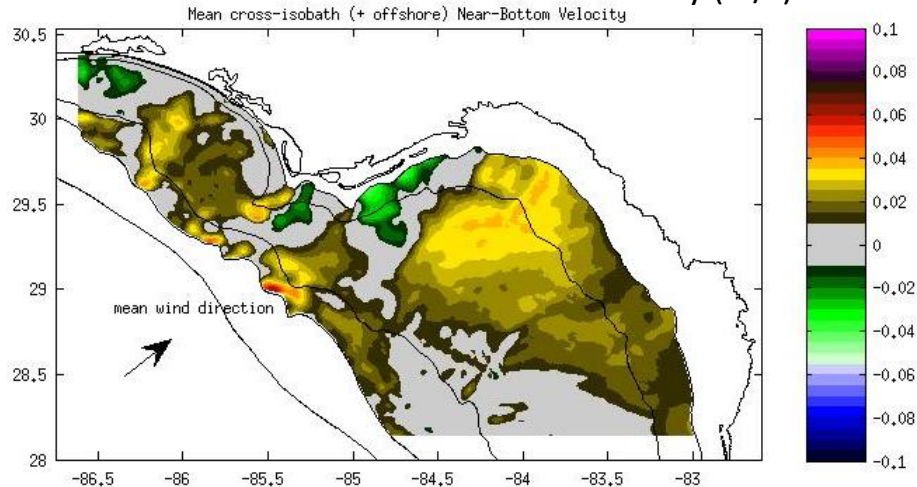


Depth-Averaged Near-Bottom Velocity



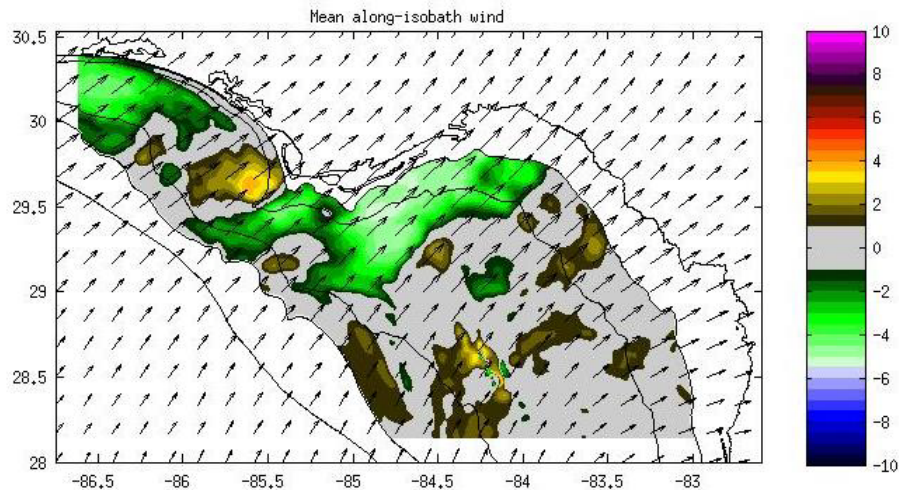


Near-bottom cross-isobath mean velocity (m/s)

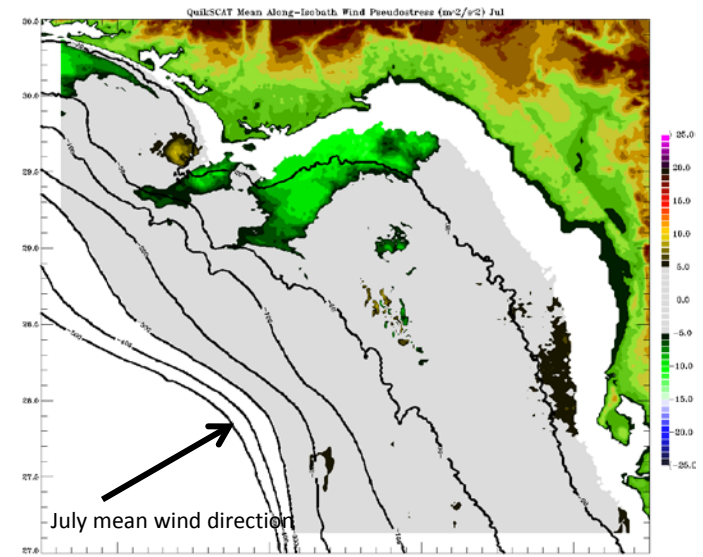


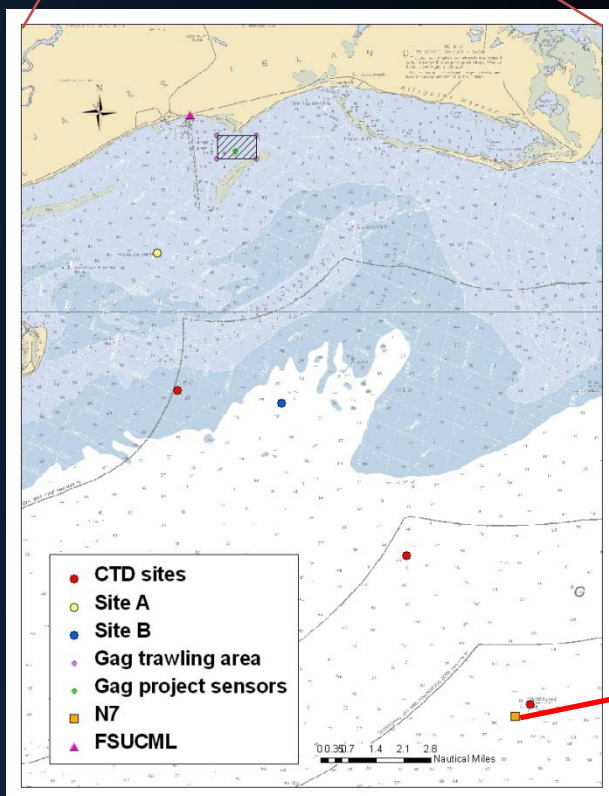
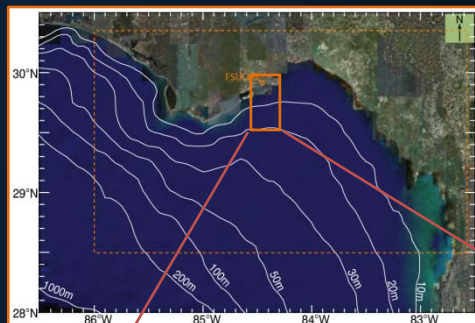
The model simulates the localized upwelling patch inferred by the scatterometer winds.

Along-isobath mean wind (m/s)

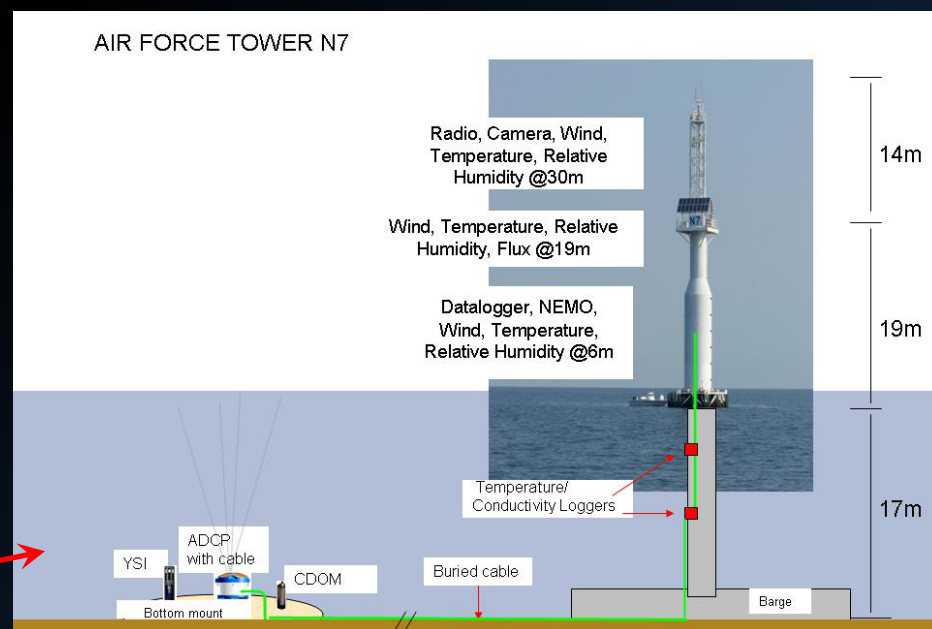
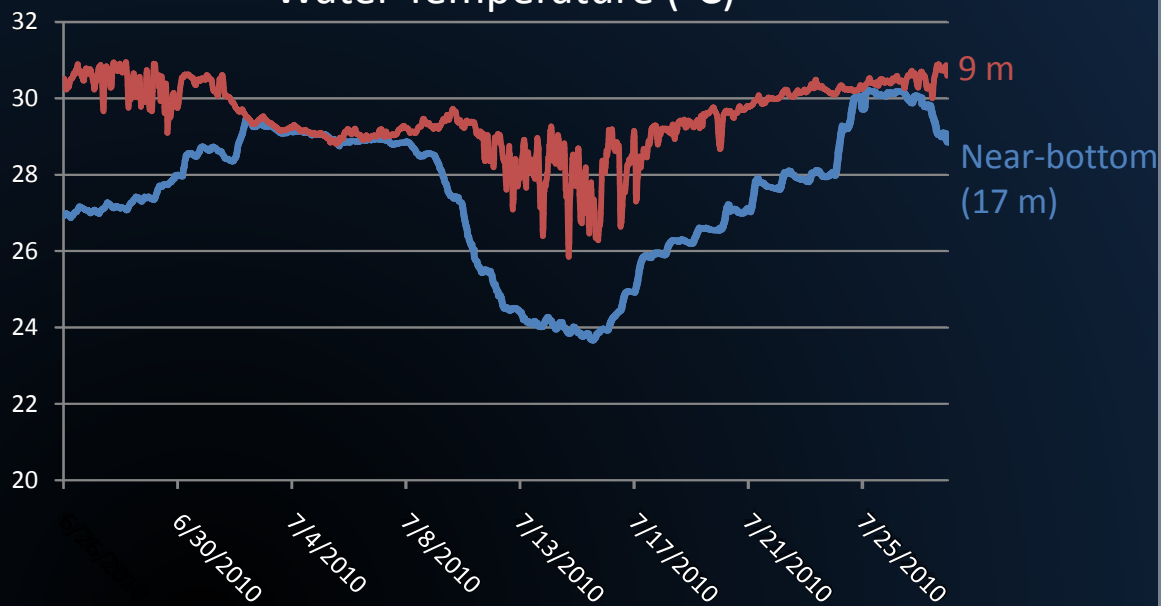


July Climatology QuikSCAT pseudostress





Water Temperature (°C)



Conclusions

- A numerical model shows that spatial inhomogeneities in near-bottom cross-isobath transport can occur due to complex shelf geometries
- Drastic curves (90°) in bathymetry contours can support localized upwelling regions where the radius of curvature is sufficiently large
- Surface winds can be reasonably used to infer the existence of localized upwelling patches given knowledge of the bathymetry.
- Current knowledge of coastal upwelling zones around the globe may be enhanced by analyzing the scatterometer-derived wind fields together with bathymetric data sets.