

Dependence of OSCAR Surface Currents on Scatterometer Winds

Kathleen Dohan, Gary S. E. Lagerloef, and John T. Gunn

Earth & Space Research Seattle, Washington

OVWST Workshop Aug 19-21, 2009 – p. 1

Introduction

Ocean Surface Currents Analyses Realtime processing system (OSCAR) is a satellite-derived surface current database based on a combination of quasi-steady geostrophic and locally wind-driven dynamics (Bonjean and Lagerloef, 2002).

- The geostrophic term is computed from the gradient of surface topography fields (AVISO/CLS).
- Wind-driven velocity components are computed from an Ekman/Stommel formulation with variable viscosity using QuikSCAT winds (FSU/COAPS)
- with a thermal wind adjustment using Reynolds SST data.
- Data available at http://www.oscar.noaa.gov and http://www.aoml.noaa.gov/phod/dac/drifter_climatology.html.







- Climate application of OSCAR
- Examine the performance of the wind-driven component of OSCAR with:
 - Wind speed
 - Temporal wind resolution
- Main area for improvement
 - Time dependent term: near inertial oscillations

Surface currents and ENSO indices





- First principal EOF of surface current (SC) and SST anomaly variations in the Tropical Pacific.
- Top: Amplitude of the EOFs normalized by their respective standard deviations. Bottom: corresponding EOF structures.
- Surface current anomalies in the tropical Pacific typically lead SST anomalies by several months.

High Wind Speed Performance





High Wind Speed Performance





High Wind Speed Performance





Relative Error





1-Day vs 5-Day Winds





Relative Error





- Relative difference between drifters and 1-day OSCAR as a function of wind speed.
- Median of points in black.

OVWST Workshop Aug 19-21, 2009 - p. 10

Missing Physics: Near-Inertial Oscillations

- OSCAR is a quasi-steady solution.
- Ekman currents are generated by vertical gradients in the wind-generated surface stress.

$$\frac{\partial \mathbf{u}_{\mathbf{E}}}{\partial t} + if\mathbf{u}_{\mathbf{E}} = \frac{\partial \tau}{\partial z}.$$

Integrate over the mixed layer depth (slab layer) gives the solution

 $\mathbf{U}_{\mathbf{E}} = -i\frac{\tau}{\rho}(1 - \exp(-ift)).$

- Steady current to the right of the wind and an anticyclonic inertial oscillation.
- Pure inertial oscillations have zero vertical group velocity and infinite horizontal extent. Propagation via:
 - β effect = variation in Coriolis with latitude causes de-correlation and finite horizontal extent
 - interactions with eddies
 - small-scale storms.

Example: Near-Inertial Oscillations





- Mooring data (Ocean Storms; Dohan and Davis 2009, JPO). Hourly winds. Two storms similar magnitude.
- Strong near-inertial oscillations in second storm due to resonant turning of wind direction, lasting well after strong winds.

OVWST Workshop Aug 19-21, 2009 – p. 12

Significance to Climate





Mixing

- Shear-driven mixing within the upper thermocline.
- The major source of internal waves in the deep ocean from the surface mixed layer.
- Wind-driven momentum transfer to the interior ocean.

Near-Inertial Oscillations: Wind Resolution









Example of winds on a daily grid.

OVWST Workshop Aug 19-21, 2009 - p. 14





High Winds and OSCAR

- OSCAR performance improves with higher winds.
- Inconclusive whether there is a very high-wind degradation seen in OSCAR.

Resolution in time

- 1-day winds show improved performance over 5-day
- High enough temporal resolution together with more dynamics than steady Ekman will be necessary to capture near-inertial oscillations (NIO).

Resolution in space

Currents will improve with higher spatial resolution. Also necessary component of NIO.

Other sources of drifter/OSCAR divergence

- Geostrophic velocity subtracted from drifters does not include filamentary features.
- Drifter slip in high winds.
- Eddy viscosity parameterization, constant in vertical.

Energy Budgets



- Munk and Wunsch (1998) estimate 2.1 TW of mixing is necessary to maintain the abyssal stratification. Wunsch and Ferrari (2004): 0.6 TW from winds, 0.9 from tides.
- Estimates of the power input from the wind from NCEP winds using a slab model is around 0.5 TW (Alford, 2003).
- GCM studies indicate smaller values of 0.1 TW (Furuichi et al., 2008,Zhai et al., 2009).
- Primitive equation models find seven times more dissipation when moving from daily wind forcing to 3 hour wind forcing (Klein et al. 2004).
- A big missing link is a detailed knowledge of the wind structure
 - ocean circulation models
 - abyssal stratification
 - climate predictions.