# Indian Ocean Intraseasonal SST Variability During Boreal Summer MJO Versus Submonthly Forcing and Processes JGR - Oceans, under revision



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### Outline:

- 1. Motivation/Importance
- 2. Background:
  - Goals & Procedure
  - Model
  - -Data
- 3. Results
- 4. Summary & Conclusion

# Motivation & Importance:

- Lack of current research (submonthly)/consensus (MJO)
- More knowledge about these processes is important:



- Atmospheric feedback
- Benefit society in India
  - 80% rainfall during summer monsoon
    → crops, flooding

Webster & Hoyos, 2004)



BBC News, 2005: http://news.bbc.co.uk/2/hi/south\_asia/4733897.stm

### Goals & Procedure:

Use a model/observation experiment to:

- 1. Assess model validity compare model SST with data
- 2. Study impact of submonthly & MJO events on SST

Use model runs to isolate processes that cause SST variability

Look at specific events & general timescales

If a process consistently dominates, then how & why?



# The Model:

- HYbrid Coordinate Ocean Model (HYCOM)
  - OGCM with hybrid vertical coordinates, 18 layers
  - Vertical Mixing: K-Profile Parameterization (Large et al. 1994)
  - Configured for Indian Ocean, horiz resolution 0.5°x0.5°
  - Model runs from 1 January 1998-November 2004
- Use a hierarchy of model experiments to isolate processes

Experiment Number:	Forcings used:	Difference Solution:	Isolates Ocean Response To:
MR	All 3-day mean fields: winds, air temp, humidity, fluxes, precipitation	MR-EXP1	Subseasonal ISOs
		MR-EXP2	Submonthly ISOs
EXP 1	Lowpassed 105 days	EXP2-EXP1	30-90 day MJO Events
EXP 2	Lowpassed 30 days	MR-EXP3	ISO wind stress
EXP 3	Lowpassed wind stress	MR-EXP4	ISO wind stress and speed
EXP 4	Lowpassed wind stress and speed	EXP3-EXP4	ISO wind speed
EXP 5	Lowpassed shortwave flux	MR-EXP5	ISO shortwave flux
EXP 6	Lowpassed precipitation	MR-EXP6	ISO precipitation



### Data Used:

- For model forcing:
  - QuickSCAT 3-day winds & derived wind stress (*Tang & Liu*)
  - ISCCP net shortwave and longwave fluxes
  - CMAP precipitation
  - ERA-40 air temperature & specific humidity
  - Surface and latent heat fluxes from bulk formulas (Kara et al. 2000)
- For model-data comparisons & further study:
  - QuickSCAT 3-day 10 meter winds
  - NOAA Satellite-observed OLR
  - TRMM SST
  - ARGO & TRITON float data





### Results: Model Validity

Time series of 30-90 & 10-30 day SST averaged over region 2 (the BOB) from 2000-2003

Thick line observed SST Dotted line modeled SST



### **Results: General Processes**

*Time series of 30-90d SST from forcings in BOB from 1999-2004 All forcings (top) and wind only (bottom).* 



### Results: Composite vs. Specific MJO Event

Composite event (left) underestimates the MJO SST extremes seen in the specific strong event (right). Color scales for composite event are ½ those for the specific event.





### Why is wind speed stronger than wind stress in the BOB?

Hmix due to Wind Speed

### **Turbulent Heat Flux!**



#### Heat Flux due to Wind Speed (Turbulent Heat Flux)

c) Heat Flux due to Wind Speed



d) Heat Flux due to Wind Speed



-100-50, -25, -5.0 5.00 25.0 50.0 100, Wm<sup>-2</sup>

# *Why* turbulent fluxes in the BOB? **Barrier layer & Thin mixed layer**





### Summary & Conclusions:

- MJO composite can average out details
- Effects of Wind >> SW, Precipitation
  - Agrees with winter events (*Han et al* 2007) and with *Waliser et al* (2004) study of canonical summer ISOs
- Effects of Wind Speed ~ Wind Stress in East Eq IO
- Effects of Wind Speed > Wind Stress in the BOB & AS
  - Less so during submonthly events
- Turbulent Heat Flux > Entrainment in the BOB
  - Barrier Layer, Shallow mixed layer

# Thank you! Any Questions?

# Begin Bonus Slides...

### Spectral Analysis: Qscat vs. ERA40



Figure 2: (a) Variance spectra of zonal wind stress  $\tau^x$ , for QuickSCAT observations (thick solid line) and scaled ERA-40 reanalysis (thick dashed line) from 1 August 1999 – 31 December 2001, averaged over the western BOB (80-90°E, 4-15°N). The thin solid/dashed lines show the 90% significance level for each. (b) Same as Figure 2a, but for meridional wind stress  $\tau^y$ . Units are Nm<sup>-2</sup>.

### **Results - Model Validity:**



Spectral Analysis of TRMM (solid line) vs. HYCOM MR (dashed line) SSTs, averaged over summers from 1999-2003, in the 3 regions of interest. Thin line is 90% confidence level for each.

### Results – Submonthly Event:

- Wind speed & stress effects are closer in magnitude in all regions
  - Speed still slightly stronger than stress in the Arabian Sea and BOB



### Hmix & Turb Flux Due to Wind Speed, Submonthly Event



c) Heot Flux due to Wind Speed

d) Heat Flux due to Wind Speed



0 25.0 m -100 -50, -25, -5.0 5.00 25.0 50.0 100, Wm<sup>-2</sup>

### Wind stress and horizontal advection equations:

- QuickSCAT-derived wind stress:
  - Have wind speed  $|\boldsymbol{V}|$  and direction  $\boldsymbol{\theta}$
  - Wind vectors  $\mathbf{V}=(u,v)$  are  $u=|\mathbf{V}|\cos\theta \& v=|\mathbf{V}|\sin\theta$
  - Then, wind stress ( $\tau$ ) is:  $\vec{\tau} = \rho_a C_D |\mathbf{V}| \mathbf{V}$  $\vec{\tau} = \tau_x + \tau_y$
- Wind stress curl is:  $curl(\vec{\tau}) = \frac{\partial}{\partial x}(\tau_y) \frac{\partial}{\partial y}(\tau_x)$
- And Ekman pumping velocity (w<sub>E</sub>) is:  $w_E = \frac{\partial}{\partial x} (\frac{\tau^y}{f}) \frac{\partial}{\partial y} (\frac{\tau^x}{f})$
- Horizontal advection due to wind stress is calculated with:

$$\Delta T = -\left[u\frac{\partial T}{\partial x} + v\frac{\partial T}{\partial y} - \left(\overline{u}\frac{\partial \overline{T}}{\partial x} + \overline{v}\frac{\partial \overline{T}}{\partial y}\right)\right]\Delta t$$

 Where u, v, & T are zonal and merid currents & SST from MR and overbars indicate fields from EXP3